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# The Raman spectrum of CaCO<sub>3</sub> polymorphs calcite and aragonite: A combined experimental and computational study

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Powder and single crystal Raman spectra of the two most common phases of calcium carbonate are calculated with *ab initio* techniques (using a “hybrid” functional and a Gaussian-type basis set) and measured both at 80 K and room temperature. Frequencies of the Raman modes are in very good agreement between calculations and experiments: the mean absolute deviation at 80 K is 4 and 8 cm<sup>-1</sup> for calcite and aragonite, respectively. As regards intensities, the agreement is in general good, although the computed values overestimate the measured ones in many cases. The combined analysis permits to identify almost all the fundamental experimental Raman peaks of the two compounds, with the exception of either modes with zero computed intensity or modes overlapping with more intense peaks. Additional peaks have been identified in both calcite and aragonite, which have been assigned to <sup>18</sup>O satellite modes or overtones. The agreement between the computed and measured spectra is quite satisfactory; in particular, simulation permits to clearly distinguish between calcite and aragonite in the case of powder spectra, and among different polarization directions of each compound in the case of single crystal spectra. © 2014 AIP Publishing LLC. [<http://dx.doi.org/10.1063/1.4871900>]

## I. INTRODUCTION

Anhydrous calcium carbonates can crystallize into three crystalline forms: calcite, aragonite, and vaterite. The nucleation and crystal growth of calcium carbonate minerals is a process widely occurring in nature, mostly as a result of biomineralization.<sup>1</sup> In fact, calcium carbonate is the major constituent of coral reefs, shells, pearls, and other biominerals, where it grows preferentially at ambient conditions due to the effect of organic templates.<sup>2</sup> This mineral also occurs in many other environments as deposition of hot springs, in stalactite and stalagmite cave formations.<sup>3</sup> From an industrial point of view, aside of the many applications as an additive, e.g., building materials, medicines, food, paper, plastics, printing inks,<sup>4</sup> it also represents a main problem in devices and desalination membranes due to the formation of scale.

While vaterite is a metastable polymorph that is mostly present in biological environments, calcite and aragonite are the main constituents of the majority of anhydrous calcium carbonate materials, and single crystals of good quality can be isolated for both these phases. Due to the complexity of vaterite, which forms extremely small single crystals having multiple structures and exhibiting chirality,<sup>5-7</sup> in this study we focus on aragonite and calcite. In particular, we present a comprehensive experimental and theoretical analysis of the Raman vibrational properties of these two phases, which includes the full classification of modes, wavenumbers, and Raman intensities. Full range Raman spectra are here obtained

from high quality single crystals, whose orientations were determined by X-ray diffraction, and polycrystalline powders. Accurate quantum mechanical *ab initio* simulations are also performed, capable of providing an extremely useful support in the processing of experimental measurements.

The first polarized Raman spectra of calcite and aragonite were measured by Couture.<sup>8</sup> This pioneering work was completed by Frech *et al.*<sup>9</sup> and Carteret *et al.*<sup>10</sup> for aragonite, and Porto *et al.*<sup>11</sup> for calcite, who discussed in detail the spectrum and polarization of Raman scattering in these calcium carbonate polymorphs. In recent times, Raman spectra have also been used as an analytical tool by chemists and mineralogists to detect the different polymorphs.<sup>12-16</sup> However, a complete characterization of all spectral peaks, in terms of both frequency and intensity, is rarely performed; the attribution of spectral features to specific modes remains a hard task, too. Previous simulations on the vibrational properties of carbonates have mainly focused on methodological aspects,<sup>17</sup> on the effect of cation substitutions on the infrared spectra,<sup>18</sup> and on the role of the vibrational contributions to the relative stability of polymorphs.<sup>19</sup> Some of the present authors have recently presented an accurate study on the infrared and Raman properties of aragonite,<sup>10</sup> and a similar one had been devoted 8 years ago to calcite;<sup>20</sup> in both cases Raman intensities were not available. The recent implementation of the calculation of Raman intensities in the CRYSTAL code<sup>21,22</sup> permits the present study to achieve the main aims of reconstructing the full spectrum of the two compounds, and using the simulated data in the interpretation of the features of the experimental spectrum.

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The paper is organized as follows: Sec. II is devoted to the description of the Raman experimental set up and of the examined samples. Section III presents the computational schemes adopted to perform the simulations. Results from theory and experiment are discussed and compared in Secs. IV–VI. Section VII presents the main conclusions.

Additional information is available in the supplementary material;<sup>23</sup> figures and tables in this section are labelled with the “S” prefix.

## II. EXPERIMENTAL DETAILS

The dimensions of the aragonite single crystal are about  $5 \times 4 \times 2 \text{ mm}^3$  along the  $a$ ,  $b$ , and  $c$  axes, respectively. The crystal was characterized by X-ray diffraction (XRD), proving a single  $\text{CaCO}_3$  aragonite phase with lattice parameters  $a = 4.9633$ ,  $b = 7.9703$ ,  $c = 5.7441 \text{ \AA}$ , very close to the published refined structure.<sup>24</sup> The crystal was oriented by XRD, cut and polished to reveal (100), (010), (001) faces. The calcite single crystal, provided by MTI corporation, is a polished plate with dimensions  $10 \times 10 \times 0.5 \text{ mm}^3$  and (100) orientation. The lattice parameters are  $a = 4.989$  and  $c = 17.062 \text{ \AA}$ .

Raman spectra were collected on a Jobin-Yvon T64000 spectrometer coupled to a  $\text{N}_2$ -cooled CCD detector and an Olympus optical microscope. Samples were excited by an argon ion laser beam at 514.5 nm using a Spectra Physics Stabilite 2017 laser. The laser beam was focused on the sample via  $\times 16$  microscope objective (long working distance and 0.20 numerical aperture) with approximately 40 mW excitation power. The back-scattered Raman spectra were collected in confocal mode with a low numerical aperture objective to avoid optical artifacts. The Rayleigh scattering was removed by an edge filter and the Raman spectra were recorded between 90 and  $1600 \text{ cm}^{-1}$  with acquisition times of 5 min. The spectral resolution was around  $1.3 \text{ cm}^{-1}$ . The spectrometer was wavenumber calibrated using Raman Shift Frequency Standards (ASTM E 1840), so that the reported wavenumbers are accurate to less than  $1 \text{ cm}^{-1}$  of actual values. The Raman spectra were also intensity corrected for the effects of the beam splitter, the edge filter, the grating, and the CCD detector response over the wavenumber range measured. The calibration was carried out with a calibrated QTH Source (Newport, Inc). This calibration permits quantitative comparison of spectral intensities obtained in different scattering orientations and allows measurements to be compared with theoretical predictions using a single scaling parameter.

The crystals were oriented along either  $a$  or  $b$  or  $c$  axes on the heating-cooling block of a Linkam THMS600 variable temperature cell deposited at the center of a rotation stage. The spectra were collected at 300 and 80 K. We made careful adjustment of the rotation stage to ensure the focused laser overlapping with the rotation stage center. The polarized Raman measurements were performed using a polarizer for both cross and parallel configurations (i.e., with respect to the incident laser polarization vector) and referred to for the (100) orientation as  $x(zz)\bar{x}$  and  $x(yz)\bar{x}$ , respectively, according to Porto's notation.<sup>25</sup> The notation of the spectrum is described by four symbols, two inside parentheses and two outside. The inside symbols are, left to right, the polarizations of the in-

cident and scattered beams, while the letters preceding and following the parentheses indicate the respective propagation directions of the incident and scattered beams. The  $x$ ,  $y$ , and  $z$  cartesian axes, according to the current laboratory reference, were selected as the  $a$ ,  $b$ , and  $c$  axes of the orthorhombic crystallographic structure, respectively. Extreme cares (orientation of the crystals, polarization of the incident laser source) were taken to avoid leakage of unexpected modes.

The Origin 8.5 program (OriginLab, Northampton, MA) was used for spectral data manipulation and curve fitting. The bands were fitted using a pseudo-Voigt line shape (mixed Gaussian and Lorentzian) with a polynomial baseline correction and were optimized by a nonlinear least-squares fitting algorithm consisting of four adjustable parameters per band (position, FWHM, intensity, and percent Lorentzian character).

## III. COMPUTATIONAL METHODS

Calculations were performed by using an all-electron Gaussian-type basis set, the hybrid Hartree-Fock/Density Functional B3LYP functional,<sup>26,27</sup> and the 2014 release of the CRYSTAL code.<sup>28,29</sup> The chosen computational setup has been shown to reproduce structure and vibrational properties of ionic and semi-ionic compounds in excellent agreement with experimental data.<sup>17,30–38</sup> Supplementary calculations were performed by means of the Hartree-Fock Hamiltonian; corresponding data for single crystal calcite at 80 K are reported in the supplementary material (Table S3). As regards intensities, critical qualitative aspects such as relative order remain unaltered when compared to B3LYP, even if quantitative agreement with experiment is slightly improved. On the other hand, three out of five frequencies show overestimations larger than  $70 \text{ cm}^{-1}$ .

A split-valence basis set of triple zeta quality with double set of  $d$  polarization functions was adopted. Oxygen, calcium, and carbon are described by  $(8s) - (411sp) - (1d1d)$ ,  $(8s) - (6511sp) - (2d1d)$ , and  $(6s) - (311sp) - (1d1d)$  contractions, respectively, which have already been used in previous investigations of calcite (basis set referred to as BSD in Ref. 20) and aragonite.<sup>10</sup> The complete basis sets and details on the computational parameters can be found on the CRYSTAL Web site,<sup>39</sup> where input and output files used for the calculations are reported. Supplementary calculations were performed upon addition of a set of  $f$  polarization functions; the corresponding exponents were optimized in the present study, by taking calcite as the reference system. For sake of comparison, data obtained with this extended basis set for single crystals at 80 K are reported in the supplementary material for both calcite and aragonite (Tables S4 and S5). Frequencies show slight variations in general; the only modes subject to larger changes are the ones below  $100 \text{ cm}^{-1}$  and above  $1000 \text{ cm}^{-1}$ , whose agreement with experiment worsens using the second basis set. More relevantly, intensities show only minor variations when adopting the second basis set; indeed, the corresponding statistical indices remain essentially unaltered.

The level of accuracy in evaluating the Coulomb and Hartree-Fock exchange series is controlled by five parameters<sup>28</sup> that have been set to 7 7 7 8 18. The threshold

on the self-consistent field (SCF) energy was set to  $10^{-10}$  Ha for both geometry optimization and construction of the Hessian. The reciprocal space was sampled according to a regular sublattice with shrinking factor<sup>28</sup>  $IS = 6$ , corresponding to 32 (calcite) and 64 (aragonite) independent  $\mathbf{k}$  vectors in the irreducible part of the Brillouin zone. The exchange-correlation contribution to the Fock matrix was evaluated by numerical integration over the unit cell volume. Radial and angular points of the grid are generated through Gauss-Legendre radial quadrature and Lebedev two-dimensional angular points distributions. In the present work, a pruned grid with 75 radial and 974 angular points was used (XLGRID keyword in the CRYSTAL14 manual<sup>28</sup>).

Structures were optimized by using the total energy analytical energy gradients with respect to atomic coordinates and unit cell parameters,<sup>40–42</sup> within a quasi-Newton scheme combined with the BFGS algorithm for Hessian updating.<sup>43–46</sup> Convergence was checked on both gradient components and nuclear displacements, for which the default values<sup>28</sup> were chosen.

The calculation of the vibrational frequencies  $\nu_0$  at the  $\Gamma$  point was performed within the harmonic approximation. The key quantity is the matrix  $H$  of the second derivatives of the total energy  $V$  with respect to the atomic cartesian coordinates  $u$ , which was constructed by numerical differentiation of the analytical gradient vector  $v$ :

$$\begin{aligned} H_{\alpha i, \beta j} &= \frac{1}{2} \left[ \frac{\partial^2 V}{\partial u_{\alpha i} \partial u_{\beta j}} \right]_0 = \frac{1}{2} \left[ \frac{\partial v_{\alpha i}}{\partial u_{\beta j}} \right]_0 \\ &\approx \frac{1}{2} \frac{v_{\alpha i}(0, \dots, u_{\beta j}, 0, \dots) - v_{\alpha i}(0, \dots, -u_{\beta j}, 0, \dots)}{2u_{\beta j}}, \end{aligned} \quad (1)$$

where Greek and Latin indices refer to atoms and atomic cartesian coordinates, respectively. Note that the calculated (optimized) equilibrium geometry is taken as a reference. The frequencies were then obtained as the eigenvectors of the mass-weighted Hessian matrix  $W$  at the  $\Gamma$  point:

$$W_{\alpha i, \beta j} = \frac{H_{\alpha i, \beta j}}{\sqrt{M_{\alpha} M_{\beta}}}, \quad (2)$$

where  $M_{\alpha}$  and  $M_{\beta}$  are the masses of atoms  $\alpha$  and  $\beta$ . Once the Hessian matrix  $H$  is calculated, frequency shifts due to isotopic substitutions can be obtained at negligible computational cost (the expensive part of the calculation being the construction of the Hessian matrix, not its diagonalization), by changing the masses in Eq. (2). Further details on the calculation of vibrational frequencies can be found in Refs. 30, 47, and 48.

Raman tensors were determined by means of a fully analytical approach<sup>21,22</sup> recently formulated by some of the present authors and implemented in the 2014 release of the CRYSTAL code. The formalism – that is an extension of the one developed in the context of infrared intensities<sup>49</sup> – is based on the combination of energy gradients of the integrals<sup>40,41</sup> with a Coupled Perturbed Hartree-Fock/Kohn Sham (CPHF/KS) scheme<sup>50,51</sup> for the response of the den-

sity with respect to the electric field. The scheme is computationally efficient, since it avoids numerical differentiation with respect to either wave vectors or atomic coordinates. No perturbation equations for the atomic displacements need to be solved.

For an oriented single crystal, the Raman integrated intensity  $I$  associated with the  $ij$  component of the Raman tensor  $\chi = \partial\alpha/\partial Q$  of the  $n$ th mode can be calculated as<sup>22,52</sup>

$$II_n^{ij} \propto C \cdot V \cdot \left( \frac{\partial \alpha_n^{ij}}{\partial Q_n} \right)^2, \quad (3)$$

where  $V$  is the unit cell volume,  $\alpha_n^{ij}$  is the polarizability component,  $Q_n$  is the normal mode coordinate, and the prefactor  $C$  gives the dependence on temperature  $T$  and frequency of laser  $\nu_L$ :<sup>22,52</sup>

$$C \sim (\nu_L - \nu_n)^4 \frac{1 + n_B(\nu_n)}{30\nu_n}, \quad (4)$$

with the Bose occupancy factor being  $1 + n_B(\nu_n) = [1 - \exp(-\hbar\nu_n/k_B T)]^{-1}$ .

For a powder polycrystalline sample, the Raman integrated intensity of the  $n$ th mode can be obtained using the following formula:<sup>52</sup>

$$II_n^{tot} \propto C \cdot V \cdot [10G_n^{(0)} + 5G_n^{(1)} + 7G_n^{(2)}], \quad (5)$$

where the terms  $G_n^{(k)}$  are the rotational invariants defined through

$$G_n^{(0)} = \frac{1}{3} (\chi_n^{xx} + \chi_n^{yy} + \chi_n^{zz})^2, \quad (6)$$

$$G_n^{(1)} = \frac{1}{2} [(\chi_n^{xy} - \chi_n^{yx})^2 + (\chi_n^{xz} - \chi_n^{zx})^2 + (\chi_n^{yz} - \chi_n^{zy})^2], \quad (7)$$

$$\begin{aligned} G_n^{(2)} &= \frac{1}{2} [(\chi_n^{xy} + \chi_n^{yx})^2 + (\chi_n^{xz} + \chi_n^{zx})^2 + (\chi_n^{yz} + \chi_n^{zy})^2] \\ &+ \frac{1}{3} [(\chi_n^{xx} - \chi_n^{yy})^2 + (\chi_n^{xx} - \chi_n^{zz})^2 + (\chi_n^{yy} - \chi_n^{zz})^2]. \end{aligned} \quad (8)$$

Conventions for the Raman tensor<sup>53–55</sup> in the case of calcite (space group (s.g.) 167) and aragonite (s.g. 62) are reported in Table I. For each computed dataset, integrated intensities are normalized so that the value for a selected reference mode is set to 1000.

A graphical representation of the calculated Raman spectra  $S(\nu)$  was obtained as a superposition of pseudo-Voigt

TABLE I. Convention for Raman tensor elements in the case of calcite (space group 167) and aragonite (s.g. 62).<sup>53–55</sup>

s.g. 167			s.g. 62			
$A_{1g}$	$E_{g,1}$	$E_{g,2}$	$A_g$	$B_{1g}$	$B_{2g}$	$B_{3g}$
$a \cdot \cdot c$	$\cdot \cdot \cdot$	$-c-d$	$a \cdot \cdot \cdot d$	$\cdot \cdot \cdot e$	$\cdot \cdot \cdot$	$\cdot \cdot \cdot$
$\cdot a \cdot \cdot$	$-c-d$	$-c \cdot \cdot$	$\cdot b \cdot d$	$\cdot \cdot \cdot$	$\cdot \cdot \cdot$	$\cdot \cdot \cdot f$
$\cdot \cdot b \cdot d$	$-d \cdot \cdot$	$\cdot \cdot \cdot$	$\cdot \cdot c \cdot \cdot \cdot e$	$\cdot \cdot \cdot$	$\cdot \cdot \cdot f$	$\cdot \cdot \cdot$

functions, one for each mode:

$$S(\nu) = \sum_n [\eta_n \cdot L(\nu; \nu_{0,n}, II_n, \gamma_n) + (1 - \eta_n) \cdot G(\nu; \nu_{0,n}, II_n, \gamma_n)] \quad (9)$$

with

$$L(\nu; \nu_{0,n}, II_n, \gamma_n) = \frac{II_n}{\pi} \left[ \frac{\gamma_n/2}{(\nu - \nu_{0,n})^2 + \gamma_n^2/4} \right], \quad (10)$$

$$G(\nu; \nu_{0,n}, II_n, \gamma_n) = 2\sqrt{\frac{\ln 2}{\pi}} \frac{II_n}{\gamma_n} \exp \left[ -\frac{4 \ln 2 (\nu - \nu_{0,n})^2}{\gamma_n^2} \right], \quad (11)$$

where  $\gamma_n$  is the full width at half maximum of the  $n$ th mode peak and  $\eta_n$  is the Lorentz factor.  $L$  (the Lorentzian function) represents the response of the system itself to light scattering, while  $G$  (the Gaussian function) reproduces the response of the experimental apparatus. The combination of the two profiles, described by the  $\gamma$  and  $\eta$  parameters, is an almost unique feature of a given experiment. For  $\gamma$ , values obtained from the experimental spectra were adopted, for better readability of the figures; they are reported in the following along with measured frequencies and integrated intensities. For  $\eta$ , a unique average value of 0.7 has been adopted for all spectra.  $S(\nu)$  curves were evaluated in the range of 50–1650  $\text{cm}^{-1}$ , in steps of 0.5  $\text{cm}^{-1}$ .

Frequencies and Raman integrated intensities (here both generically indicated as  $x$ ) have been analyzed according to the following statistical indices ( $M$  is the number of data in the set): mean absolute difference between the calculated and the experimental values  $|\overline{\Delta}| = M^{-1} \sum_{i=1}^M |x_i^{calc} - x_i^{exp}|$ , mean difference  $\overline{\Delta} = M^{-1} \sum_{i=1}^M (x_i^{calc} - x_i^{exp})$ , percentage mean absolute difference relative to the average of the experimental data  $|\overline{\Delta}| \% = |\overline{\Delta}| / \overline{x^{exp}} \cdot 100 = \sum_{i=1}^M |x_i^{calc} - x_i^{exp}| / \sum_{i=1}^M x_i^{exp} \cdot 100$ , maximum absolute difference  $|\Delta|_{max} = \max_{i=1}^M |x_i^{calc} - x_i^{exp}|$ .

Graphical animations of the normal modes are available at Refs. 56 and 57 so that the reader can directly interpret the “nature” of the mode (stretching, bending, rotation, translation, etc.) in a simple and intuitive manner.

Manipulation and visualization of structures were performed with the Jmol 3D engine.<sup>58,59</sup> Data analysis was performed using the LibreOffice suite<sup>60</sup> and the Octave environment.<sup>61</sup> Graphs were realized with the Gnuplot utility.<sup>62</sup>

#### IV. STRUCTURE AND SYMMETRY

The computed cell parameters for calcite are  $a = 5.038$  and  $c = 17.325$  Å, with a 1.0% and 1.5% overestimation, respectively, compared with the experimental data. In the case of aragonite, calculations yield 5.008, 8.029, and 5.861 Å for the  $a$ ,  $b$ , and  $c$  axes, the differences with respect to the experiment being +0.9%, +0.7%, and +2.0%. These discrepancies are typical for the adopted functional (see, for example, Ref. 35).

The rhombohedral primitive cell of calcite contains 2  $\text{CaCO}_3$  formula units, for a total of 10 atoms; its 27 vibrational modes can be classified according to the irreducible representations of the  $\bar{3}m$  point group as follows:

$$\Gamma_{total} = 1A_{1g} \oplus 2A_{1u} \oplus 3A_{2g} \oplus 3A_{2u} \oplus 4E_g \oplus 5E_u.$$

$A_{1g}$  and  $E_g$  (double degenerate) modes are Raman active,  $A_{2u}$  and  $E_u$  (double degenerate) are IR active,  $A_{1u}$  and  $A_{2g}$  are spectroscopically inactive (silent modes).

The orthorhombic cell of aragonite contains 4 formula units, i.e., 20 atoms; its 57 vibrational modes can be classified according to the irreducible representations of the  $mmm$  point group as follows:

$$\Gamma_{total} = 9A_g \oplus 6A_u \oplus 6B_{1g} \oplus 8B_{1u} \oplus 9B_{2g} \oplus 5B_{2u} \oplus 6B_{3g} \oplus 8B_{3u}.$$

$A_g$ ,  $B_{1g}$ ,  $B_{2g}$ , and  $B_{3g}$  modes are Raman active,  $B_{1u}$ ,  $B_{2u}$ , and  $B_{3u}$  are IR active,  $A_u$  modes are silent.

#### V. POWDER POLYCRYSTALLINE DATA

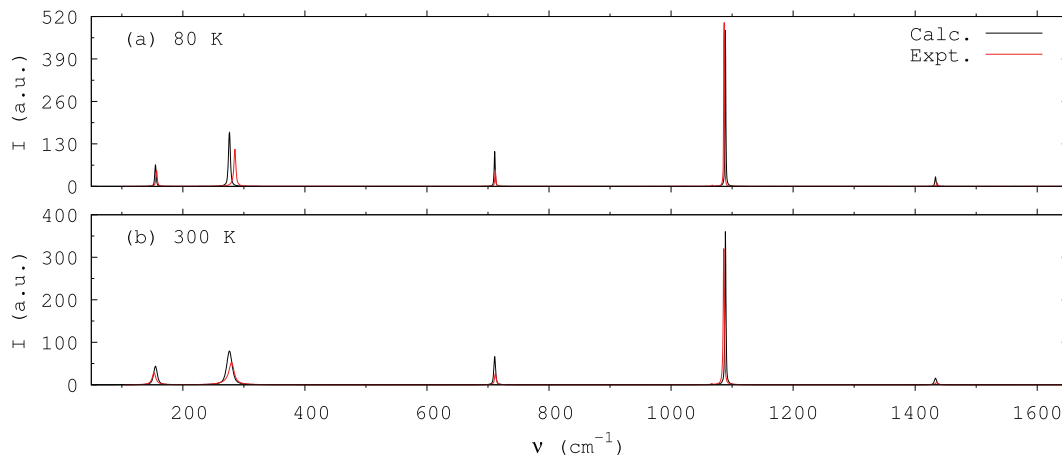
Data for calcite and aragonite (at both 80 and 300 K) are reported in Tables II and III, respectively; the corresponding spectra are shown in Figures 1 and 2. Note that on the

TABLE II. Calculated and experimental Raman properties of powder polycrystalline calcite. In the calculations the temperature effect was taken into account for integrated intensities only. Frequencies  $\nu$  and peak widths  $\gamma_{exp}$  are in  $\text{cm}^{-1}$ ; for each temperature Raman integrated intensities were re-normalized so that value for  $A_{1g}$  mode at 1089.1  $\text{cm}^{-1}$  is equal to 1000.  $\Delta$  are the differences between calculated and experimental values.

	Symmetry	$\nu$					$II_{tot}$							
		80 K		300 K			80 K			300 K			$\gamma_{exp}$	
	Calc.	Expt.	$\Delta$	Expt.	$\Delta$	Calc.	Expt.	$\Delta$	Calc.	Expt.	$\Delta$	80 K	300 K	
1	$A_{1g}$	1089.1	1087.1	2.0	1086.2	2.9	1000	1000		1000	1000		1.5	2.0
2	$E_g$	155.1	159.0	-3.9	154.9	0.2	239	184	55	425	295	130	2.6	7.1
3		276.3	286.9	-10.6	281.2	-4.9	843	633	210	1134	851	283	3.7	10.4
4		711.0	712.4	-1.4	712.4	-1.4	294	136	158	302	133	169	2.0	3.3
5		1433.3	1436.6	-3.3	1435.8	-2.5	96	34	62	95	33	62	2.4	4.5
$ \overline{\Delta} $				4.2		2.4			121			161		
$\overline{\Delta}$				-3.4		-1.1			121			161		
$ \overline{\Delta}  \%$									49			49		
$ \Delta _{max}$				10.6		4.9			210			283		

TABLE III. Calculated and experimental Raman properties of powder polycrystalline aragonite. For each temperature Raman integrated intensities were re-normalized so that value for  $A_g$  mode at  $1095.3 \text{ cm}^{-1}$  is equal to 1000. See caption to Table II for more details.

	Symmetry	$\nu$					$I_{tot}$						$\gamma_{exp}$	
		80 K			300 K		80 K			300 K			80 K	300 K
		Calc.	Expt.	$\Delta$	Expt.	$\Delta$	Calc.	Expt.	$\Delta$	Calc.	Expt.	$\Delta$		
1	$A_g$	148.7	144.4	4.3	142.9	5.8	31	27	4	57	60	-3	2.2	4.0
2		161.9	165.7	-3.8	162.2	-0.3	34	20	14	59	24	35	2.2	4.0
3		195.8	199.0	-3.2			28	5	23	44			3.0	
4		205.0	219.4	-14.4	214.7	-9.7	97	63	34	150	99	51	4.1	8.7
5		280.2	291.0	-10.8	283.6	-3.4	42	30	12	57	48	9	5.2	16.3
6		704.2	705.0	-0.8	706.1	-1.9	146	96	50	151	102	49	1.8	2.6
7		862.8	853.8	9.0	853.3	9.5	10	5	5	10	4	6	1.7	2.2
8		1095.3	1087.2	8.1	1085.3	10.0	1000	1000		1000	1000		1.5	1.9
9		1473.9					6			6				
10	$B_{1g}$	97.4	117.6	-20.2	113.5	-16.1	59	30	29	130	33	97	4.0	7.3
11		152.1	156.2	-4.1	153.0	-0.9	627	660	-33	1127	950	177	3.1	7.9
12		199.0					1			2				
13		213.4	226.5	-13.1	225.5	-12.1	25	9	16	38	9	29	4.5	10.0
14		705.5					11			11				
15		1463.9	1464.1	-0.2	1462.2	1.7	35	29	6	34	25	9	3.4	5.6
16	$B_{2g}$	182.5	183.9	-1.4	180.0	2.5	141	78	63	232	103	129	2.9	6.2
17		207.2	212.3	-5.1	206.3	0.9	339	370	-31	523	430	93	3.0	6.2
18		249.2	252.1	-2.9	248.3	0.9	82	68	14	115	52	63	5.6	9.7
19		260.7	266.1	-5.4	260.1	0.6	4	45	-41	5	45	-40	3.2	9.4
20		278.7					23			31				
21		714.6	716.7	-2.1	717.1	-2.5	7	10	-3	8	7	1	2.0	2.8
22		911.8	918.7	-6.9	918.2	-6.4	1	0	1	1	1	0	1.7	2.0
23		1091.6					0			0				
24		1591.8	1579.1	12.7	1573.9	17.9	31	11	20	31	10	21	4.6	7.8
25	$B_{3g}$	101.3	126.6	-25.3	125.0	-23.7	14	4	10	29	5	24	4.0	6.0
26		167.6	181.0	-13.4			94	11	83	160			3.7	
27		177.8	193.4	-15.6	190.7	-12.9	18	41	-23	29	80	-51	3.5	8.7
28		271.4	278.4	-7.0	272.0	-0.6	60	41	19	82	46	36	5.2	9.5
29		701.2	700.0	1.2	701.8	-0.6	138	64	74	142	75	67	1.6	2.2
30		1415.0					0			0				
$\overline{ \Delta }$				8.0		6.4			26			47		
$\overline{\Delta}$				-5.0		-1.9			15			38		
$\overline{ \Delta }\%$									23			32		
$ \Delta _{max}$				25.3		23.7			82			177		

FIG. 1. Calculated and experimental Raman spectra of powder polycrystalline calcite. In the calculations the temperature effect was taken into account for integrated intensities only. For each temperature Raman integrated intensities were re-normalized so that the value for  $A_{1g}$  mode at  $1089.1 \text{ cm}^{-1}$  is equal to 1000.

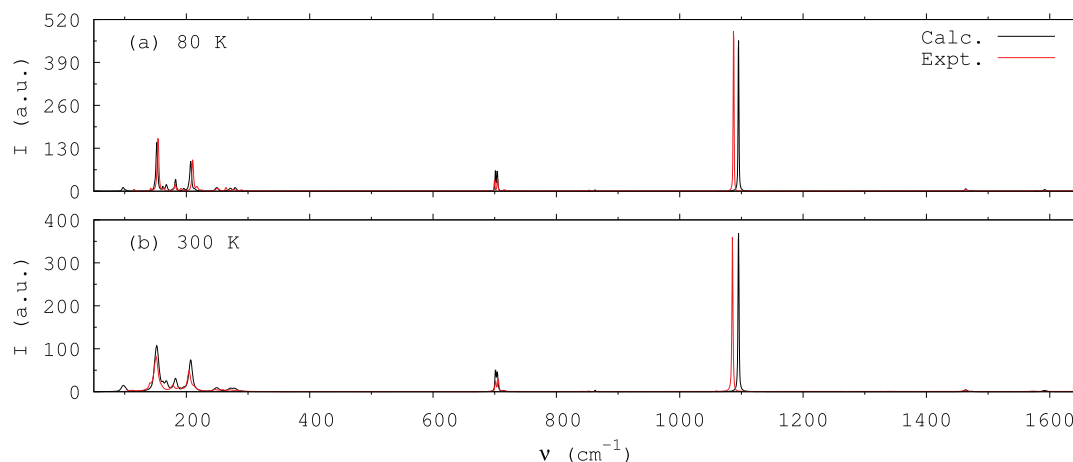


FIG. 2. Calculated and experimental Raman spectra of powder polycrystalline aragonite. In the calculations the temperature effect was taken into account for integrated intensities only. For each temperature Raman integrated intensities were re-normalized so that the value for  $A_g$  mode at  $1095.3 \text{ cm}^{-1}$  is equal to 1000.

computational side the effect of temperature has been taken into account for the Raman tensors (i.e., for the integrated intensities), not for the vibrational frequencies. For each dataset Raman integrated intensities are re-normalized so that the value for a selected reference peak is set to 1000. Calculations at present do not provide values for the peak widths  $\gamma$ , so only the experimental values are given in the tables, for sake of reference. As calculated and experimental Raman frequencies for these compounds have already been compared and discussed in previous studies by some of the present authors,<sup>10,20</sup> the present discussion will focus on Raman intensities and spectra. Finally, in the following and for sake of brevity integrated intensities will be simply referred to as intensities.

Calcite has 5 Raman active modes (Table II); for frequencies, the mean absolute and maximum deviations between calculation and experiment at 80 K are 4.2 and  $-10.6 \text{ cm}^{-1}$ , respectively. If we have a look at the data for Raman intensities at 80 K, we note that all deviations are positive, with a mean absolute value of 121 and a maximum value of 210 (corresponding to the  $E_g$  mode computed at  $276.3 \text{ cm}^{-1}$ ). As the intensities span 4 orders of magnitude, it is worth defining a

percentage mean absolute deviation to permit the comparison between different datasets; in this case it is equal to 49%.

Aragonite has 30 active Raman modes (Table III). At 80 K 24 out of the 30 modes can be identified in the experimental spectrum. Of the 6 missing modes, 2 do have null computed intensity (modes 23 and 30), modes 14 and 20 are quite close to other peaks, whereas modes 9 and 12 sum up the two problems. The mean absolute deviation for frequencies is  $8.0 \text{ cm}^{-1}$ , whereas the maximum value is  $-25.3 \text{ cm}^{-1}$ . As regards Raman intensities, the mean absolute deviation is 26; most discrepancies (19 out of 24) are positive, so that the mean deviation is positive and equal to 15. The percentage mean absolute deviation is 23%; the maximum discrepancy is 82.

If we compare results at 300 K with those at 80 K, for both compounds we observe an improved agreement for frequencies, whereas deviations from experimental intensities increase. The fact that computed frequencies are closer to the room temperature rather than to the low temperature experimental values is likely related to the overestimation of the lattice parameters typical for the adopted Hamiltonian, which has been discussed in Sec. IV; such overestimation

TABLE IV. Calculated and experimental Raman properties of single crystal calcite at  $T = 80 \text{ K}$ . Raman tensors were re-normalized so that the  $a^2$  element of  $A_{1g}$  mode at  $1089.1 \text{ cm}^{-1}$  is equal to 1000. The symbols # indicate modes identified in the experiments with no correspondence in the calculations. See caption to Table II for more details.

Symmetry	Calc.	Expt.	$\Delta$	Calc.	Expt.	$\Delta$	Calc.	Expt.	$\Delta$	$\gamma_{\text{exp}}$	
	$\nu$			$a^2$			$b^2$				
#		1067.4			10			2		1.7	
1	$A_{1g}$	1089.1	2.0	1000	1000		90	182	-92	1.4	
#		1749.8			23			13		5.5	
		$\nu$			$c^2$			$d^2$			
2	$E_g$	155.1	-4.5	52	17	35	118	99	19	2.1	
3		276.3	-11.1	73	58	15	526	354	172	3.2	
4		711.0	-1.6	187	56	131	21	23	-2	1.8	
5		1433.3	-3.8	17	13	4	51	7	44	2.1	
			4.6						57		
			-3.8						36		
									64		
			11.1						172		

TABLE V. Calculated and experimental Raman properties of single crystal aragonite at  $T = 80$  K. Raman tensors were re-normalized so that the  $c^2$  element of  $A_g$  mode at  $1095.3 \text{ cm}^{-1}$  is equal to 1000. The symbols # indicate modes identified in the experiments with no correspondence in the calculations. See caption to Table II for more details.

	Symmetry	Calc.	Expt.	$\Delta$	Calc.	Expt.	$\Delta$	Calc.	Expt.	$\Delta$	Calc.	Expt.	$\Delta$	$\gamma_{\text{exp}}$
		$\nu$			$a^2$			$b^2$			$c^2$			
1	$A_g$	148.7	144.7	4.0	3	1	2	1	2	-1	79	51	28	1.5
2		161.9	166.1	-4.2	49	11	38	6	0	6	45	44	1	1.7
3		195.8	199.2	-3.4	4	3	1	55	5	50	20	2	18	2.1
4		205.0	219.9	-14.9	5	1	4	234	120	114	43	27	16	2.9
5		280.2	291.5	-11.3	0	0	0	87	39	48	36	18	18	3.7
6		704.2	705.0	-0.8	1	1	0	217	104	113	216	99	117	1.6
7		862.8	854.0	8.8	24	5	19	0	2	-2	0	4	-4	1.8
#			1061.2			1			3			3		1.6
8		1095.3	1087.2	8.1	169	210	-41	977	980	-3	1000	1000		1.4
9		1473.9	1466.2	7.7	2	0	2	6	0	6	5	12	-7	1.8
			$\nu$			$d^2$			$e^2$			$f^2$		
10	$B_{1g}$	97.4	117.6	-20.2	87	47	40							3.7
11		152.1	156.2	-4.1	931	918	13							2.8
12		199.0	211.7	-12.7	2	5	-3							3.5
#			220.0			3								3.0
13		213.4	227.0	-13.6	37	23	14							4.1
14		705.5	705.9	-0.4	16	14	2							1.6
15		1463.9	1464.1	-0.2	51	29	22							1.8
16	$B_{2g}$	182.5	184.6	-2.1				210	87	123				1.9
17		207.2	212.8	-5.6				503	415	88				2.2
18		249.2	252.5	-3.3				121	75	46				4.7
19		260.7	267.0	-6.3				6	53	-47				2.4
20		278.7	284.1	-5.4				34	6	28				5.0
21		714.6	716.9	-2.3				11	10	1				1.7
22		911.8	911.1	0.7				1	0	1				2.0
23		1091.6						0						
24		1591.8	1579.0	12.8				47	17	30				4.8
25	$B_{3g}$	101.3	126.2	-24.9							20	8	12	2.7
26		167.6	181.0	-13.4							139	17	122	2.3
27		177.8	193.7	-15.9							26	50	-24	2.3
28		271.4	278.5	-7.1							90	47	43	3.7
29		701.2	700.1	1.1							205	96	109	1.5
30		1415.0									0			
				7.7										32
				-4.6										26
														31
				24.9										123

induces a general underestimation of the vibrational frequencies. Note also that, in the case of aragonite, at 300 K there are 2 more missing modes, due to superposition with more intense modes, which is in turn a consequence of the broadening of the peaks associated to the higher temperature.

As regards Raman intensities, a general overestimation of calculation with respect to experiment is observed, probably related to both uncertainties due to the experimental setup and approximations of the simulation. Concerning the latter aspect, basis set is the most likely source of discrepancy. For this study we adopted a high quality basis set for solid state calculations, whose accuracy in describing structure, vibrational frequencies, and infrared intensities of minerals has been widely documented (see Sec. III). It is known from molecular calculations<sup>63</sup> that Raman intensities may re-

quire even richer basis sets, such as aug-cc-pVTZ or Sadlej pVTZ; however, development and assessment of these kinds of basis functions has so far not yet been addressed for solids.

Figures 1 and 2 show the Raman powder spectra for calcite and aragonite at both temperatures. Despite the discrepancies observed for the intensities, the relevant features of the spectra are reproduced satisfactorily in all cases. It is worth comparing the powder spectra of calcite and aragonite (at 80 K for instance); as pointed out in Ref. 13, the two phases show the most relevant differences in the range 0–300  $\text{cm}^{-1}$ : calcite has two peaks around 160 and 290  $\text{cm}^{-1}$ , whereas aragonite has two intense peaks around 160 and 210  $\text{cm}^{-1}$ . Notably, the two figures show that the computational description of this frequency range permits a clear distinction between the two phases.



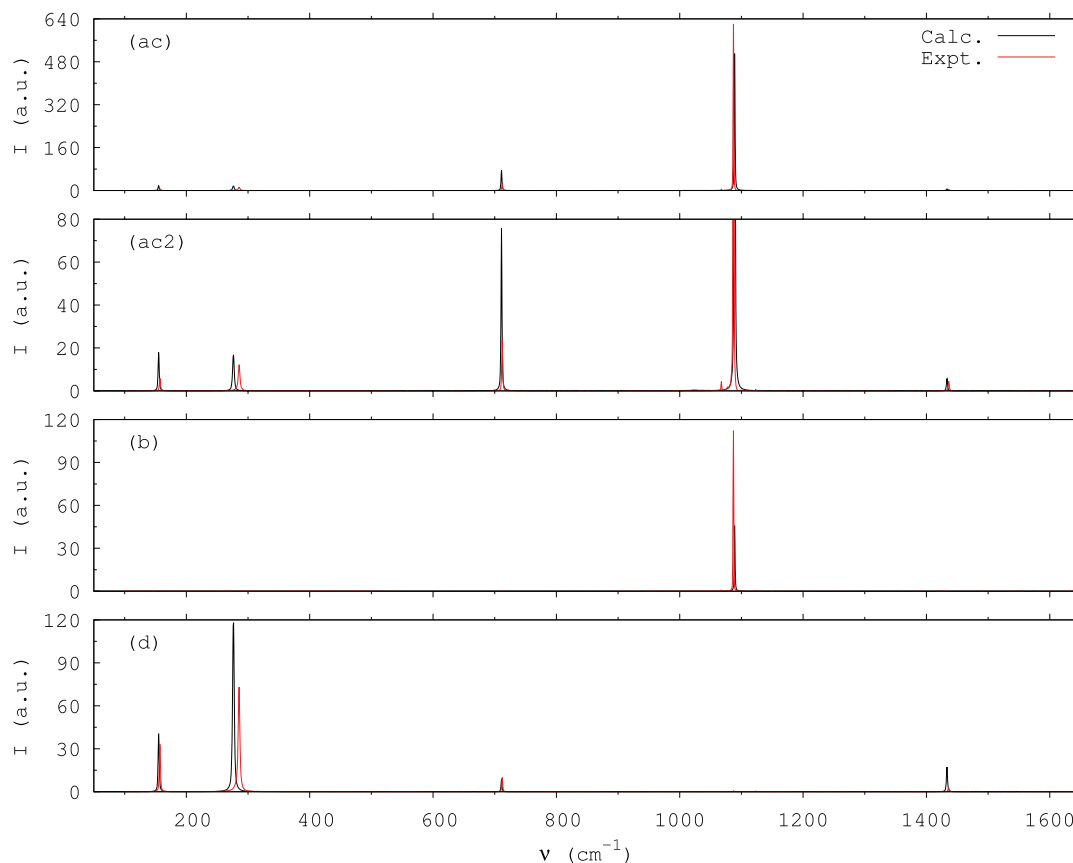


FIG. 3. Calculated and experimental Raman spectra of single crystal calcite at  $T = 80$  K. In the calculations the temperature effect was taken into account for integrated intensities only. Raman integrated intensities were re-normalized so that the XX value for  $A_{1g}$  mode at  $1089.1 \text{ cm}^{-1}$  is equal to 1000. (ac) XX polarization,  $a^2 + c^2$  contributions ( $A_{1g}$  plus  $E_g$  symmetries); (ac2) XX polarization, zoom on the  $c^2$  contributions scale; (b) ZZ polarization,  $b^2$  contributions ( $A_{1g}$ ); and (d) XZ polarization,  $d^2$  contributions ( $E_g$ ).

## VI. SINGLE CRYSTAL DATA

In the following discussion, we will focus on the data at 80 K, to take advantage of the improved spectral resolution in the experiments. Data for calcite and aragonite are given in Tables IV and V, respectively; the corresponding spectra are shown in Figures 3 and 4. Single crystal data at 300 K are reported in the supplementary material (Tables S1 and S2, and Figures S1 and S2).

### A. Calcite

Let us start discussing the 5 fundamental Raman modes of calcite (Table IV). Concerning the frequencies, as reasonably expected the mean absolute and maximum deviations are very close to the values found in the case of the powder data: 4.6 and  $-11.1 \text{ cm}^{-1}$ , respectively.

For single crystals, intensity data correspond to the elements of the Raman tensor of each mode (see also Table I). In this respect, single crystal data show a worse quality than the powder ones: the percentage mean absolute deviation is as large as 64%. The mean absolute deviation is 57; 7 out of 9 discrepancies are positive, resulting in a mean deviation of 36. The maximum deviation is 172 and corresponds to the  $d^2$  element of  $E_g$  mode computed at  $276.3 \text{ cm}^{-1}$ . Other two large deviations, 131 and  $-92$ , are found for the  $c^2$  element of  $E_g$

mode at  $711.0 \text{ cm}^{-1}$  and for the  $b^2$  element of the  $A_{1g}$  mode at  $1089.1 \text{ cm}^{-1}$ , respectively.

We now comment on two spectral features that are found in the experimental spectra but not in the computed data. They are two minor, but yet visible, peaks at  $1067.4$  and  $1749.8 \text{ cm}^{-1}$ . These peaks are observed only in the spectra corresponding to XX and ZZ polarizations, which implies they must have a  $A_{1g}$  symmetry (to observe the peak at  $1749.8 \text{ cm}^{-1}$ , see Figure S3 in the supplementary material). As regards the first mode, at  $1067.4 \text{ cm}^{-1}$ , simulated isotopic substitution suggests this peak is the  $^{18}\text{O}$  satellite mode of the very intense peak at  $1087.1 \text{ cm}^{-1}$  (isotopic shift  $19.7 \text{ cm}^{-1}$ ). Graphical animation ([www.crystal.unito.it/prtfreq/jmol.html](http://www.crystal.unito.it/prtfreq/jmol.html)) shows that the computed mode at  $1089.1 \text{ cm}^{-1}$  is the symmetric C–O stretching in the  $\text{CO}_3^{2-}$  units. Upon substitution of a  $^{16}\text{O}$  atom in a carbonate group with a  $^{18}\text{O}$  atom, the carbonate units with “regular”  $^{16}\text{O}$  atoms still stretch at this frequency; on the other hand, units bearing the  $^{18}\text{O}$  isotope decouple from the others and stretch at  $1068.3 \text{ cm}^{-1}$ . The corresponding computed shift is  $20.8 \text{ cm}^{-1}$ , in excellent agreement with the experimental one.

The second additional peak is at  $1749.8 \text{ cm}^{-1}$ , well beyond the highest frequency of the fundamental modes,  $1437.1 \text{ cm}^{-1}$ . As proposed by Gillet *et al.*,<sup>64</sup> this peak should be attributed to an overtone mode. By generating all the direct

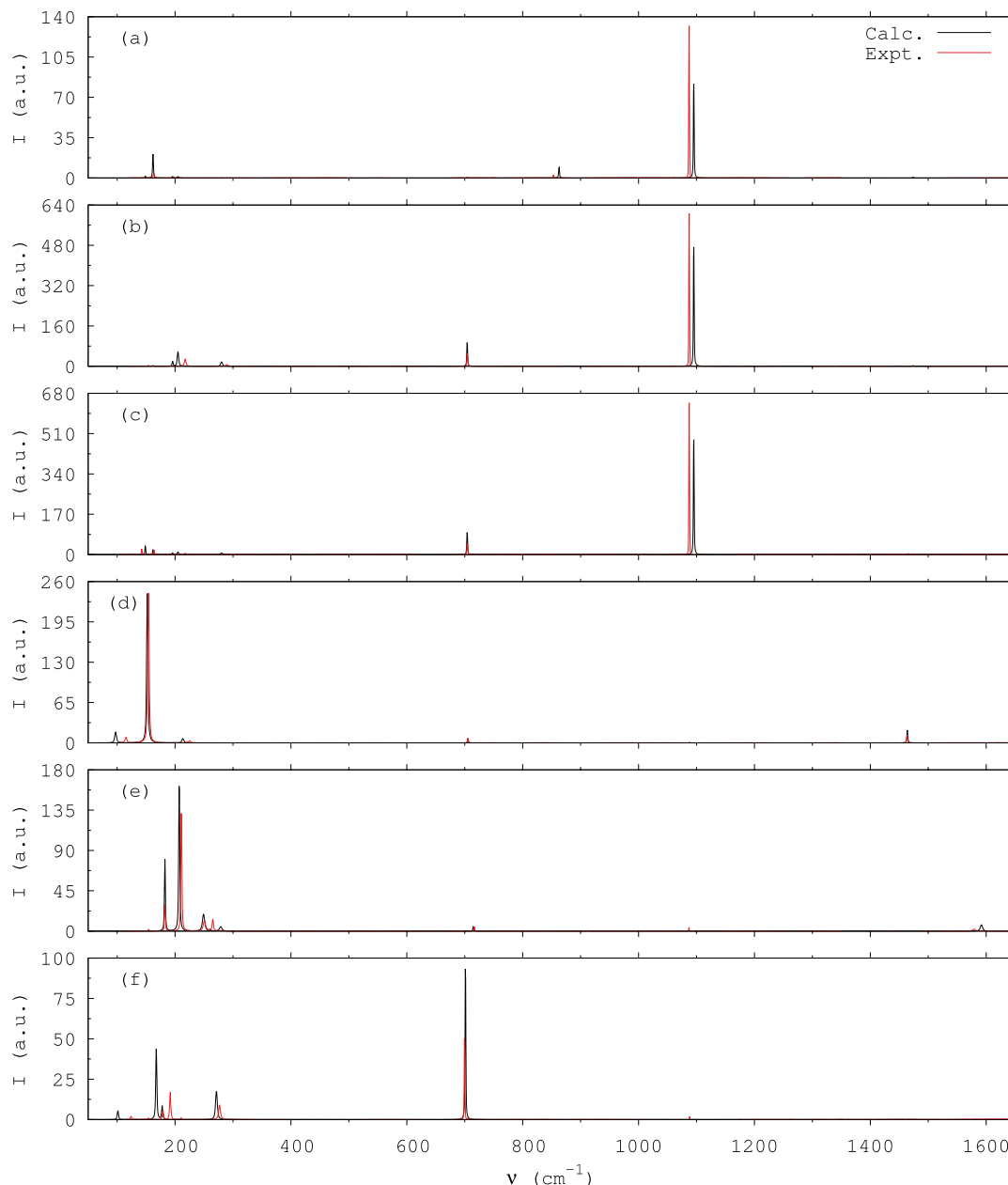


FIG. 4. Calculated and experimental Raman spectra of single crystal aragonite at  $T = 80$  K. In the calculations the temperature effect was taken into account for integrated intensities only. Raman integrated intensities were re-normalized so that the ZZ value for  $A_g$  mode at  $1095.3 \text{ cm}^{-1}$  is equal to 1000. (a) XX polarization,  $a^2$  contributions ( $A_g$  symmetry); (b) YY polarization,  $b^2$  contributions ( $A_g$ ); (c) ZZ polarization,  $c^2$  contributions ( $A_g$ ); (d) XY polarization,  $d^2$  contributions ( $B_{1g}$ ); (e) XZ polarization,  $e^2$  contributions ( $B_{2g}$ ); and (f) YZ polarization,  $f^2$  contributions ( $B_{3g}$ ).

products of couples of fundamental modes yielding a mode with  $A_{1g}$  symmetry, two compatible combinations are found:  $2 \times 874.5$  ( $A_{2u}$  symmetry, IR active) =  $1749.1 \text{ cm}^{-1}$  and  $2 \times 882.5$  ( $A_{2g}$ , silent) =  $1765.0 \text{ cm}^{-1}$ . The corresponding deviations with respect to the experimental value are  $-0.7 \text{ cm}^{-1}$  and  $15.2 \text{ cm}^{-1}$ . Even taking into account a possible anharmonic effect (not included in the calculations), the second value seems very far from the experiment, suggesting that the mode is more likely the overtone of the IR active mode computed at  $874.5 \text{ cm}^{-1}$ . Graphical animation ([www.crystal.unito.it/prtfreq/jmol.html](http://www.crystal.unito.it/prtfreq/jmol.html)) shows that both modes are bendings of the  $\text{CO}_3^{2-}$  units; in the former the two carbonates of the unit cell move in phase, in the latter they move in anti-phase.

## B. Aragonite

Of the 30 fundamental Raman modes of aragonite (Table V), 28 can be identified in the single crystal measurement. Both missing modes (23 and 30) have 0 computed intensity. Compared with the powder experiment, 4 additional modes are identified: 9, 12, 14, and 20. As for calcite, the mean absolute and the maximum deviations are very close to the values found in the case of the powder data:  $7.7$  and  $-24.9 \text{ cm}^{-1}$ .

A look at the intensities indicates that single crystal data for aragonite have a slightly larger percentage mean absolute deviation, 31%, than in the powder case (23%). The mean absolute deviation is 32. Thirty-six out of 45 discrepancies

are positive, giving a mean deviation of 26. The maximum deviation is 123, which corresponds to the  $e^2$  element of the  $B_{2g}$  mode computed at  $182.5\text{ cm}^{-1}$ . Other 5 intensities have discrepancies larger than 100:  $A_g$  modes computed at  $205.0\text{ cm}^{-1}$  ( $b^2$  element) and  $704.2\text{ cm}^{-1}$  ( $b^2$  and  $c^2$  elements), and  $B_{3g}$  modes at  $167.6$  and  $701.2\text{ cm}^{-1}$  ( $f^2$  element); also the  $e^2$  element of the  $B_{2g}$  mode at  $207.2\text{ cm}^{-1}$  shows a quite large deviation, 88.

Also in the case of aragonite we identify experimental peaks not belonging to the list of computed fundamental Raman modes: one at  $1061.2\text{ cm}^{-1}$  (spectra with XX, YY, ZZ polarization;  $A_g$  symmetry) and one at  $220.0\text{ cm}^{-1}$  (spectrum with XY polarization;  $B_{1g}$  symmetry). Similar to the case already discussed for calcite, the peak at  $1061.2\text{ cm}^{-1}$  is attributed to a  $^{18}\text{O}$  satellite mode of the intense peak at  $1087.2\text{ cm}^{-1}$  (isotopic shift  $26.0\text{ cm}^{-1}$ ). Graphical animation ([www.crystal.unito.it/prtfreq/jmol.html](http://www.crystal.unito.it/prtfreq/jmol.html)) indicates this mode as the symmetric C–O stretching in the  $\text{CO}_3^{--}$  units; the computed isotopic shift is  $23\text{ cm}^{-1}$ , again in good agreement with the experimental one. This satellite mode has already been discussed in more detail by these authors.<sup>10</sup> As regards the additional mode at  $220.0\text{ cm}^{-1}$ , a satisfactory assignment could not be found in the present study.

### C. Spectra

Figures 3 and 4 show the Raman single crystal spectra for calcite and aragonite at 80 K. Zoomed views of the spectra for aragonite are available in the supplementary material (Figure S4).

Despite the discrepancies described above for frequencies and, in particular, for intensities, the spectra for the various polarization directions are well characterized by the calculations, and clearly distinguished from each other. Indeed, for calcite the spectrum with  $a^2 + c^2$  contributions, labeled as (ac) in Figure 3, has 5 peaks, the most intense being at  $1089.1\text{ cm}^{-1}$ ; the spectrum for  $b^2$ , (b) in the figure, has only one peak at  $1089.1\text{ cm}^{-1}$ ; the spectrum for  $d^2$ , (d) in the figure, has 4 peaks, the most intense being at  $155.1$  and  $276.3\text{ cm}^{-1}$ .

Also in the case of aragonite, the computed spectra permit to clearly distinguish among different polarizations. There are three spectra involving modes with  $A_g$  symmetry; they feature the contributions from Raman tensors elements  $a^2$ ,  $b^2$ , and  $c^2$ , and are labeled as (a), (b), and (c), respectively, in Figure 4 (as well as in Figure S4 in the supplementary material). The three corresponding computed spectra all have the most intense peak at  $1095.3\text{ cm}^{-1}$ . Spectrum (a) has two relatively intense peaks at  $161.9$  and  $862.8\text{ cm}^{-1}$ ; spectrum (b) has three features at  $704.2$ ,  $205.0$ , and  $280.2\text{ cm}^{-1}$ , whereas spectrum (c) has three features at  $704.2$ ,  $148.7$ , and  $161.9\text{ cm}^{-1}$ . The most intense peak of the spectrum (d) for  $d^2$  is at  $152.1\text{ cm}^{-1}$ , with other two features at  $1463.9$  and  $97.4\text{ cm}^{-1}$ . Spectrum (e) for  $e^2$  has the most intense peak at  $207.2\text{ cm}^{-1}$  and another feature at  $182.5\text{ cm}^{-1}$ . Finally, spectrum (f) for  $f^2$  has the most intense peak at  $701.2\text{ cm}^{-1}$ ; as regards other relatively intense peaks, calculation indicates  $167.6$  and  $271.4\text{ cm}^{-1}$ , at slight variance with the experiment ( $193.7$  and  $278.5\text{ cm}^{-1}$ ).

## VII. CONCLUSIONS

We have performed accurate *ab initio* calculations and experimental measurements of the Raman spectra of both calcite and aragonite phases of calcium carbonate.

The agreement between calculated and experimental frequencies is very good: the mean absolute deviation for calcite at 80 and 300 K is around 4 and  $2\text{ cm}^{-1}$ , respectively; the same values for aragonite are 8 and  $6\text{ cm}^{-1}$ . Calculation generally underestimates the experiment, which is likely related to the particular choice of the model Hamiltonian.

Calculated Raman intensity values, spanning a range of four orders of magnitude, agree reasonably well with the experimental data. Upon rescaling of the intensities to the most intense peak of each dataset, computed values generally overestimate measured ones. This issue is probably related to approximations and uncertainties from both computational and experimental sides, and calls for the development of superior quality basis sets for solid state calculations.

Overall, the agreement between the computed and measured spectra is quite satisfactory. In particular, simulation permits to clearly distinguish between calcite and aragonite in the case of powder spectra, as well as among different polarization directions of each compound in the case of single crystal spectra.

The combined analysis permits to identify and characterize almost all the fundamental Raman peaks of the two compounds, with the exception of either modes with zero computed intensity or modes overlapping with more intense peaks. Two additional peaks have been identified in both calcite and aragonite. In the former case, they correspond to  $^{18}\text{O}$  satellite mode and to an overtone, respectively. In the latter case, one peak has been assigned to a  $^{18}\text{O}$  satellite mode, too; the second peak remains at the moment unassigned.

As a final remark, previous papers by these authors have shown how powerful can be the interplay between theory and experiment in analyzing and interpreting the infrared spectra of minerals. The detailed discussion presented in the current paper strengthens this argument, extending the synergistic action to the field of Raman spectroscopy.

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