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1	Thermo-elastic behaviour of grossular garnet at high pressures and
2	temperatures
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24	ABSTRACT
25	The thermo-elastic behaviour of synthetic single crystals of grossular garnet (Ca ₃ Al ₂ Si ₃ O ₁₂) has been
26	studied in situ as a function of pressure and temperature separately. The same data collection protocol has
27	been adopted to collect both the pressure-volume $(P-V)$ and temperature-volume $(T-V)$ datasets in order to
28	make the measurements consistent with one another. The consistency between the two datasets allows
29	simultaneous fitting to a single pressure-volume-temperature equation of state (EoS), which was performed
30	with a new fitting utility implemented in the latest version of the program EoSFit7c. The new utility
31	performs fully weighted <i>simultaneous</i> fits of the <i>P-V-T</i> and <i>P-K-T</i> data using a thermal pressure Equation of
32	State combined with any PV EoS. Simultaneous refinement of our $P-V-T$ data combined with that of K^{T} as a
33	function of T allowed us to produce a single $P-V-T-K^2$ equation of state with the following coefficients:
34 25	$V_0 = 1664.46(5)$ A ² , $K_{T0} = 166.5/(17)$ GPa and $K = 4.96(7) \alpha_{(300K, 1bar)} = 2.09(2)^{*}10^{-1}$ K ⁻¹ with a
35	refined Einstein temperature ($\theta_{\rm E}$) of 512K for a Holland-Powell-type thermal pressure model and a Tait 3 -
30 27	order EoS. Additionally, thermodynamic properties of grossular nave been calculated for the first time from
37 20	theory within the framework of the quesi hermonic approximation
20	theory within the namework of the quasi-narmonic approximation.
39 40	Kowwords: grossular high pressure high temperature diffraction bulk modulus P.V.T.K. fit. EosEit
41	Keywords . grossular, ingit-pressule, ingit-temperature, unitaction, burk modulus, 1 - v - 1 - K m, Ebst n
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1. INTRODUCTION

Stable over a wide range of pressure and temperature conditions, garnets are among the most 53 54 abundant phases in the Earth's upper mantle and transition zone. Yet coherent and homogeneous equation of state coefficients to reliably describe their behaviour at simultaneous high pressures and 55 temperatures still have to be determined with sufficient accuracy and precision. Recent papers have 56 shown the importance of a robust and valid description of the elastic behaviour of the most 57 58 abundant mineral phases for the accurate calculation of host-inclusion entrapment pressures that in 59 turn allow the growth conditions in the Earth to be inferred (e.g. Angel et al. 2014, 2015a, 2015b; 60 Ashley et al. 2015; Milani et al. 2015). However, the determination of a full set of elastic coefficients of a mineral is always challenging because of the limitations in performing 61 simultaneous high-pressure (P) and temperature (T) experiments on one side and because of the 62 lack of consistent measurements performed separately at high-P and high-T. Therefore, a reliable 63 description of the elastic behaviour of minerals at high-P,T with a single set of Equation of State 64 65 (EoS) coefficients can be extremely challenging.

Pressure-volume EoS are well-developed and are capable of reproducing the isothermal 66 volume or density changes of materials to within the experimental uncertainties. A natural approach 67 to describing the *P*-*V*-*T* behaviour of a material is therefore to describe how the volume (V_{0T}), bulk 68 moduli (K_{0T}) and its pressure derivatives (K_{0T}) change with increasing temperature (at room 69 pressure), and then use these parameters to calculate the isothermal compression at the temperature 70 of interest. A potential limitation of this approach is that while the variation of K_{0T} with T can be 71 measured, the variation of K_{0T} with T has been measured very rarely, although it is obvious that it 72 should increase slightly with increasing temperature. As Helffrich and Connolly (2009) pointed out, 73 the common assumption that K_{0T} does not change with temperature in combination with the 74 75 approximation that $\partial K_{0T}/\partial T$ is constant, often leads to the prediction of non-physical negative

¹ Notation: subscripts on K (e.g. $K_{P,T}$) indicate the reference pressure and temperature to which the bulk modulus is referred; superscripts S or T (e.g. K^{S} or K^{T}) stand for adiabatic or isothermal quantities.

thermal expansion coefficients at reasonably modest pressures for a large number of materials. An alternative approach is to employ the idea of thermal pressure (e.g. Anderson 1995). Then the total pressure (P) at a given V and T can be expressed as the sum of two terms:

$$P(V,T) = P(V,T_{ref}) + Pth(T)$$

The function $P(V, T_{ref})$ is the isothermal equation of state for material at the reference temperature, up to the volume at P and T_{ref} . As shown in the Figure 1, if $T > T_{ref}$, then $P(V, T_{ref}) < P(V, T)$. The thermal-pressure Pth(T) is the pressure that would be created by increasing the temperature from T_{ref} to T at constant volume, which is the isochor of the material passing through the final P,T point of interest. As shown in Figure 1, the concept of thermal pressure is therefore essentially a way of calculating P(V,T) along a different path in P-T space. We can consider that the isothermal compression part that gives us $P(V, T_{ref})$ is well-defined. The slope of the isochor is given by

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$$\left(\frac{\partial P}{\partial T}\right)_{V} = \frac{\alpha}{\beta} = \alpha K^{T}$$
 (where α is the volume thermal expansion and β is the compressibility of the

88 material) so the thermal pressure is $Pth = \int_{Tref}^{T} (\alpha K^T)_{\nu} dT$. In order to use a thermal-pressure EoS one

89 then has to know, or to make an assumption, about how the product αK^T varies along the isochor. 90 For many materials and *P*,*T* ranges of interest (e.g. Figure 1) the temperatures are in the region of 91 the Debye temperature, so the isochor is linear and thus αK_T is a constant. In this case 92 $Pth \sim (\alpha K^T)_V (T - Tref)$.

A number of other assumptions or models of the behaviour of the material can be made, including the Debye and similar models which make assumptions about the vibrational and hence thermal behaviour of materials. Holland and Powell (2011) noted that the form of αK^T against temperature should have the shape of a heat capacity function and developed an approximation for the integral of αK^T that involves an Einstein function, as follows:

$$Pth = \alpha_0 K_{00} \left(\frac{\theta_E}{\xi_0} \right) \left(\frac{1}{\exp(\theta_E/T) - 1} - \frac{1}{\exp(\theta_E/T_{ref}) - 1} \right)$$

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with ξ_0 being the same as in the thermal expansion equation of Kroll et al. (2012). This thermal-99 pressure model has the physically-correct properties that the product αK^T becomes constant at high 100 temperatures while it decreases to zero at low temperatures. This means that both the bulk modulus 101 102 K_{0T} and the thermal expansion become constant at low temperatures, and both have an approximately linear variation with temperature above θ_E . The exact expressions for thermal 103 expansion and bulk modulus as a function of temperature depend on the choice of isothermal 104 equation of state, but at T_{ref} and zero pressure $\alpha = \alpha_0$ (Holland and Powell 2011). In this study we 105 106 used the Holland and Powell (2011) model of thermal pressure in combination with the Tait isothermal EoS (so as to maintain compatibility with Thermocalc, Holland and Powell 2011) to 107 determine the EoS of grossular (Ca₃Al₂Si₃O₁₂) garnet. We show that the thermal pressure approach 108 109 allows the full P-V-T EoS to be determined by fitting simultaneously measurements of the volume 110 and bulk modulus over a range of temperature but only at room pressure. Additional separate 111 measurements of the volume compression curve at high-P can be used to provide either an 112 independent test of the EoS parameters, or be used to further constrain them in simultaneous global fits of all available data. This approach has the advantage of not requiring data from simultaneous 113 114 high-P, high-T measurements which remain challenging with respect to both the control and accurate measurement of P and T, and with respect to the precision in volume measurements. A 115 comparison with previous literature data (e.g. Skinner 1956; Isaak et al. 1992; Thieblot et al. 1998; 116 117 Kono et al. 2010; Gwanmesia et al. 2014; Du et al. 2015 etc..) is provided together with a further 118 crosscheck of our experimental data against the results obtained by density functional theory (DFT) 119 ab-initio simulation. Thermodynamic properties of grossular have been calculated for the first time 120 from crystal Helmholtz and Gibbs energies, including the contribution from phonons, using density 121 functional theory within the framework of the quasi-harmonic approximation.

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125	2. METHODS
126	2.1 Sample synthesis and characterization
127	Synthetic single crystals of pure end-member grossular (Gr ₁₀₀) garnets were synthesized at

high pressures with a multi-anvil apparatus at the Bayerisches Geoinstitut. Stoichiometric mixtures 128 129 of CaO (prepared by decarbonation of CaCO₃), Al(OH)₃ and SiO₂ were packed into platinum 130 capsules of 3.5 mm length and 2 mm diameter. Samples were synthesized at 6 GPa and 1300°C. The experimental sample was heated in 23 mins and the temperature was kept at 1300°C for 40 131 mins. Each multi-anvil experiment was performed with Cr-doped MgO octahedra of 18 mm edge 132 length combined with tungsten carbide cubes of 11 mm truncation-edge lengths. For all the 133 134 experiments a graphite heater was employed. Experiments were quenched by cutting the electrical power supply to the furnace. Single-crystals of up to ca. 100 µm in diameter were recovered from 135 136 the capsules. The compositions of the crystals have been determined by electron microprobe at the 137 University of Padova using a Cameca CAMEBAX-micro operating at 15 nA and 15 kV with standards of andradite for Si, spinel for Al and diopside for Ca. Results from the chemical analyses 138 are reported in Table 1. Two crystals (with size ca. 80 x 30 x 30 µm) were selected for the X-ray 139 140 diffraction study based on the absence of twinning and visible inclusions and on the quality of their 141 diffraction peak profiles (FWHM not greater than 0.15°).

142

143 2.2 Thermal expansion behaviour by single crystal X-ray diffraction experiments

144 The high-temperature single-crystal X-ray diffraction experiments were performed at the 145 University of Pavia with a HUBER four-circle point detector diffractometer operating at 50 kV and 146 30 mA (Mo $K\alpha$ radiation) and equipped with a 0.8 mm short collimator and automated with the 147 SINGLE software (Angel and Finger 2011). This diffractometer is equipped with a newly designed Page 5 of 23

micro-furnace controlled by a Eurotherm temperature regulator (Alvaro et al. 2015). For the high-T148 149 experiments, one single crystal of grossular was placed inside a thin-walled quartz vial (0.3 mm inner diameter and 26 mm long) and held in position by means of quartz wool. The vial was then 150 mounted on a metal goniometer head on the diffractometer. During the high-temperature 151 152 experiments the effects of the crystal offsets and the diffractometer aberrations were compensated by using the eight-position centring method (King and Finger 1979) before starting the high-T153 measurement. The unit-cell parameters (Table 2 and Figure 2) were determined at 23 different 154 155 temperatures from room-T up to 1020 K using only 4-position centring because of the spatial restrictions described in Alvaro et al. (2015). A further 7 measurements were then made on cooling 156 back to room temperature. Each individual measurement took about 12 hours, and the high-157 temperature experiment lasted for a total of 14 days. Unconstrained unit-cell parameters confirmed 158 the cubic symmetry within 1 e.s.d. Therefore, only constrained unit-cell parameters obtained by 159 160 vector least-squares fit (Ralph and Finger 1982) are reported in Table 2. Unit-cell parameters determined on the heating and cooling cycles are indistinguishable. 161

162 The low-temperature single-crystal X-ray diffraction experiment was performed on the very 163 same crystal used for the high-T experiment still mounted inside its quartz vial. The experiment was 164 carried out at the Institut für Mineralogie und Kristallographie at the University of Wien using a Stoe StadiVari diffractometer equipped with an Incoatec IuS microsource (Mo-target, with multi-165 166 layer mirror optics, operating at 50kV and 0.1mA) and a Pilatus 300K area detector. Low temperatures were achieved using an Oxford Cryojet system that allows measurements to be 167 performed between 100 and 500 K. Full intensity data collections up to $110^{\circ} 2\theta$ (with a coverage of 168 99% and redundancy higher than 4) were performed at 10 different temperatures from room-T down 169 170 to 100 K. Correction of the crystal offsets from the goniometer center were made by first 171 performing a short data collection and adjusting the crystal position to make the frame scale factors flat. Unconstrained unit-cell parameters confirmed the cubic symmetry within 1 e.s.d.; therefore, 172 only constrained unit-cell parameters are reported in Table 2 and Figure 2 together with those 173 Page 6 of 23

174 obtained at high-*T*.

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176 2.3 Compressibility behaviour by single crystal X-ray diffraction experiments

High-pressure single-crystal X-ray diffraction experiments for the grossular garnet were 177 178 carried out with a Huber four-circle diffractometer at the Bayerisches Geoinstitut. The diffractometer operates at 50 kV and 40 mA (MoKa radiation), is equipped with a point detector 179 and is automated by SINGLE software (Angel and Finger 2011). The sample was loaded in a ETH-180 181 type diamond-anvil cell (DAC, Miletich et al. 2000) using a steel gasket, pre-indented to 100 µm thickness and with a hole diameter of 250 µm. Methanol:ethanol mixture 4:1 was used as pressure 182 transmitting medium, which remains hydrostatic up to about ~ 9.5 GPa (Angel et al. 2007; Klotz et 183 184 al. 2009). A single crystal of quartz was loaded in the DAC together with the sample and used as a pressure standard (Angel et al. 1997). During the centring procedure, the effects of crystal offsets 185 186 and diffractometer aberrations were eliminated from the refined peak position by the eight-position centring method (King and Finger 1979). Unconstrained unit-cell parameters were obtained by 187 188 vector least-squares fit (Ralph and Finger 1982) on not less than 20 reflections up to $2\theta = 29^{\circ}$. The 189 symmetry-unconstrained unit-cell edges show deviations smaller than 1 e.s.d. from the constrained 190 ones thus confirming the cubic symmetry within 1 e.s.d.. Therefore, only the constrained unit-cell parameters have been reported in Table 3 and Figure 3. 191

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193 2.4 Fitting of the pressure-volume-temperature-modulus data using EoSFit7c

All fits described in this manuscript have been performed with the EosFit7c program (Angel et al. 2014). Previous versions of this program were able to fit isothermal EoS, thermal expansion models and the thermal-pressure EoS of Holland and Powell (2011) to P-V, V-T or P-V-T data, by the method of least-squares. We have now extended the capabilities of EosFit-7c to also fit all of these forms of EoS to bulk and linear moduli, either alone or in combination with volume or unitcell parameter data. Fits are performed by weighted non-linear least squares to minimise the Page 7 of 23

residuals P_{obs} - P_{calc} , using the effective variance method (Orear 1982) to convert the experimental 200 uncertainties in volume and bulk moduli, pressure and temperature of each datum into an effective 201 uncertainty in P (Angel 2000). Input bulk moduli can be either adiabatic (K^{s}) or isothermal values 202 (K^{T}) , but in either case they must be Reuss values, corresponding to the response of the material 203 204 under hydrostatic pressure. The Voigt and Reuss bulk moduli of a material quantify its volume response to the application of uniform strain and uniform stress (hydrostatic pressure) respectively. 205 When uniform stress is applied to a cubic material, it undergoes isotropic (i.e. uniform) strain. 206 207 Therefore, the Voigt and Reuss bulk moduli of a cubic material are identical. We remind the reader that this is not true for materials of lower symmetry and therefore the use of Voigt values, or Voigt-208 Reuss-Hill averages, in Equations of state of crystals with symmetry lower than cubic is 209 inconsistent with the theory underlying equations of state under hydrostatic conditions, and will 210 lead to incorrect results (e.g. Angel et al. 2009). 211

212 If the input moduli to EosFit7c are isothermal, no further conversion is made, and the observational equations for such data in the least-squares are simply the expression for the bulk 213 214 modulus in terms of pressure for each EoS. These expressions are direct for invertible EoS such as 215 the Murnaghan or Tait EoS (Freund and Ingalls 1989; Holland and Powell 2011; Angel et al. 2014). For non-invertible EoS the procedure used is to calculate the compression ratio V/V_{00} corresponding 216 to the observed K and then use the P-V expression of the EoS to calculate the corresponding 217 pressure. If adiabatic moduli are provided as input data, they are converted to isothermal values by 218 $K^{S} = (1 + \alpha_{V} \gamma T) K^{T}$ where the value of thermal expansion (α_{V}) is taken from the current EoS at the 219 220 pressure and temperature of interest. The Grüneisen parameter γ can be expressed in terms of measurable quantities, for example $\gamma = \frac{\alpha_V K^S}{C_p} = \frac{\alpha_V K^T}{C_V}$. However, since heat capacity data is not 221 available at elevated pressures, in EosFit7c we use the simple approximation (Anderson 1996) that 222

223 $\gamma(P,T) = \gamma_0 \left(\frac{V(P,T)}{V_0}\right)^q$, with the value of q close to 1 for 'normal' solids under modest P,T

conditions. Only Boehler and Ramakrishnan (1980) and Boehler (1982) appear to have attempted to 224 measure the value of q, showing that it scales with volume. This means that the Grüneisen 225 parameter scales with volume, increasing with increasing temperature and decreasing upon 226 compression, which when considered as a measure of the pressure-dependence of the frequencies of 227 vibrational modes, is reasonable. This will break down when the material becomes very 228 229 anharmonic, near to structural phase transitions, but should not be a problem for materials such as the grossular garnet in this study. The value of γ_0 (i.e. γ value at the reference conditions) and of q 230 can be set by the user in Eosfit7c. 231

For linear moduli that describe the variation of the unit-cell parameters with pressure as 232 $M_i^T = -a_i \frac{\partial P}{\partial a_i}$, the procedure is analogous to that used for linear dimensions in EosFit7 (Angel et 233 al., 2014). Internally the linear moduli are converted to volume-like bulk moduli as $K_i = \frac{M_i}{3}$ and 234 treated with volume EoS. For the conversion of adiabatic to linear moduli, we use the relationship 235 between isothermal and adiabatic compressibilities $\beta_i^T = \beta_i^S + \frac{\alpha_i \alpha_v T}{C_n}$ which can be derived from 236 between isothermal and adiabatic elastic compliances 237 the fundamental relationship $s_{ijkl}^{T} = s_{ijkl}^{S} + \frac{\alpha_{ij}\alpha_{kl}T}{C_{p}}$ with the C_{p} in units of Jm⁻³K⁻¹ (Nye 1985). The substitution $\frac{\gamma}{K^{S}} = \frac{\alpha_{V}}{C_{p}}$ gives 238 $\beta_i^T = \beta_i^S + \frac{\alpha_i \gamma T}{K^S} = \beta_i^S \left(1 + \frac{\alpha_i \gamma T \beta_V^S}{\beta_V^S} \right)$ which can be rewritten in terms of 239 moduli as

240
$$M_i^S = \left(1 + \alpha_i \gamma T \frac{M_i^S}{K^S}\right) M_i^T$$
. It is important to notice that for cubic materials $\alpha_i = \alpha_V/3$ and $\frac{M_i^S}{K^S} = 3$

so that only for cubic materials this can be written as $M^{s} = (1 + \alpha_{v}\gamma T)M^{T}$ which is the same relation as for the bulk moduli.

- 243
- 244 2.5 Ab-initio DFT (Density Functional Theory)

Ab-initio calculations have been performed in order to crosscheck our experimental results at 245 simultaneous high-P and T. Thermodynamic properties of grossular have been calculated from 246 crystal Helmholtz and Gibbs energies, including the contribution from phonons, using density 247 functional theory (DFT) within the framework of the quasi-harmonic approximation (OHA). 248 249 Periodic DFT calculations were performed with the plane-wave VASP package (Kresse and Furthmüller 1996) within GGA approximation, with projector-augmented waves (PAW, Blöchl 250 1994) and the PBEsol functional (Perdew et al. 2008). A plane-wave cutoff of 520 eV (1.3 x default 251 252 VASP value) was used as recommended in the VASP package. Energy convergence was tested for two k-meshes 2x2x2 and 4x4x4. Since the denser mesh did not lead to any significant reduction in 253 energy ($\Delta E < 0.001 \text{ eV}$) the 2x2x2 mesh was used both for unit cell optimization and subsequent 254 calculation of interatomic forces. The cell was optimized at 17 pressure points within the range [-8 255 to 8GPa] with step of 1 GPa. The threshold for electronic and ionic convergence was set to 10⁻⁸ and 256 10^{-6} eV, respectively. The atomic forces were converged to 10^{-5} eV/Å. Phonon frequencies at 257 constant cell volumes were calculated in program PHONOPY (Togo and Tanaka 2015) which takes 258 advantage of the direct method of Parlinsky et al. (1997). The force constants used to construct the 259 dynamical matrices (input for PHONOPY) were calculated in the VASP package using density 260 261 functional perturbation theory (DFPT, Baroni et al. 2001) and taking into account symmetries of the conventional body-centered cell of the garnet, containing 160 atoms. The results obtained with 262 263 DFT-GGA commonly overestimate cell volume values of ionic solids by up to 9.5% with respect to experimental results while PBEsol functional brings the discrepancy down to 2% (Perdew et al. 264 2008). In our case, the DFT calculations vielded volume values of 1664.6 $Å^3$ (static DFT), 1681.9 265 $Å^3$ (zero-point vibration included) and 1687.3 $Å^3$ (T=290 K) which agree with experimental values 266 of 1664.3 Å at room P and T, (0.0001 GPa and 300 K) within 0.0 %, 1.1% and 1.4% respectively. 267 This volume overestimation amounts to about -2 GPa, on the pressure scale, with respect to the 268 269 experimental datum.

271

3. RESULTS

272 *3.1 Temperature-volume data*

The room-PT unit-cell volumes (V_{00}) of grossular as determined in this study, agree with one 273 another to better than 1 e.s.d. thus confirming the reliability of the unit-cell lattice determination 274 275 with different instruments and measurement protocols. Therefore, we combined the high-Tmeasurements together with those at low-T in a single dataset without any normalization or re-276 scaling. The evolution of the unit-cell volume with temperature is reported in Figure 2 and Table 2 277 278 where the low-T and the high-T data are reported together. A continuous increase of the unit-cell 279 volume is observed as a function of temperature with no evidence of any irreversible change in the crystal occurred up to the maximum temperature reached in this study. Data collected both with 280 increasing and decreasing temperature overlap within experimental error, thus indicating good 281 experimental reproducibility. The temperature-volume data for the entire range of T studied were fit 282 283 using the software EoSFit7c (Angel et al. 2014) using a Kroll-type EoS (Kroll et al. 2012) allowing for refinement of the Einstein temperature (θ_E) but not refining K' value that was kept fixed to 5 284 (i.e. the value from compressibility data, see the following sections). The thermal expansion 285 coefficients obtained were $\alpha_{(298K, 1bar)} = 2.08(2)*10^{-5} \text{ K}^{-1}$, $V_{00} = 1664.45(7) \text{ Å}^3$ with $\theta_E = 515(20) \text{ K}$. 286 287 We compared our T-V for grossular with those from Isaak (1992) determined by dilatometry. Isaak et al. (1992) measured physical expansions, and converted such values to volume expansions 288 289 Yv(obs). Conversion of Isaak et al. (1992) data (taken from their Table 2) to unit-cell volumes has been performed according to $V = V_0 (1 + Y_V obs)$. In order to ensure that our measurements and those 290 by Isaak et al. (1992) were on the same volume scale we used our V_{00} at the same reference 291 temperature of 293K used by Isaak et al. (1992). A plot of the two datasets (see Figure 2) shows 292 293 that they agree within their mutual uncertainties.

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295 *3.2 Temperature-volume-bulk modulus data*

Isaak et al. (1992) also determined the adiabatic bulk modulus of grossular by RUS measurements and then used measured heat capacity data and their own thermal expansion data to also calculate the isothermal bulk modulus at each measurement temperature by means of $K^{s} = (1 + \alpha_{v}\gamma T)K^{T}$, using the value of the Grüneisen parameter $\gamma = \frac{\alpha_{v}K^{s}}{C_{p}}$. The reliability of the heat capacity (C_{p}) data by Isaak et al. (1992) has been confirmed by comparison with the more

recent data by Bosenick et al. (1996) (see Figure 4). The C_p values determined by Isaak et al. (1992) fall well within the scatter of the data determined by Bosenick et al. (1996) and are also in good agreement with those we have obtained from the DFT calculations.

Using K and V data as a function of T it is obviously possible to refine V_{00} , α_0 , θ_E and K_{00} . But 304 in addition it is possible and required to refine K' because this affects the results when using the 305 thermal pressure model as it appears in the isothermal EoS $P(V, T_{ref})$, and at room pressure 306 $P(V, T_{ref}) = -Pth(T)$. Simultaneous refinement of our T-V data and the T-K^S data from Isaak et al. 307 (1992) for grossular using $\gamma_0=1.22$ and q=0 (i.e. $\gamma=\gamma 0$) yielded identical results as to those obtained 308 from fitting the *T*-*V*- K^{T} data (see Table 4²). This shows that the assumption that γ is constant is a 309 reasonable approximation for grossular. This observation is also in agreement with Isaak et al. 310 311 (1992) who calculated that γ drops by 0.1(± 0.05) over the temperature range from 300K to 1200K and with our results from DFT calculations as shown in Figure 4 according to which γ drops to 1.24 312 at 650K and then it assumes constant values (e.g. changes of about 2 at the 3rd decimal place) above 313 650K. 314

315

316 *3.3 Pressure-volume data*

The unit-cell volume decreases smoothly with increasing pressure, as shown in Figure 3 and reported in Table 2, up to the maximum hydrostatic pressure reached in this study of ~7.5 GPa.

² The Equation of State coefficients reported in Table 4c representing the best fit of our data have been deposited as supplementary material in the form of .eos files (that can be read with EoSFit7c). They are also available directly from the authors.

Such pressure covers the stability pressure range for upper mantle garnets. Fit of the *P-V* data alone using a 3rd order Tait EoS yielded the following coefficients: $V_{00} = 1664.36(7)$ Å³, $K_{00} =$ 167.45(1.04) GPa and K_{00} = 4.93(31) also reported in Table 4a. The goodness of the fit results was confirmed by the low ΔP_{max} value (calculated as $|P_{obs} - P_{calc}|$) of 0.010 GPa. The equation of state coefficients obtained with this method agree with those obtained by *T-V-K* fitting to better than 1 e.s.d. thus confirming the reliability of the results from the two fitting approaches.

325

326 *3.4 P-V-T-K and PVT Equation of State*

The remarkably good agreement between the lattice parameters as determined with different 327 experimental settings and that of the EoS coefficients obtained with different procedures allowed us 328 to perform a simultaneous refinement of the entire set of EoS coefficients by combining our 329 pressure-volume and temperature-volume data together with those of temperature-bulk modulus 330 331 (Isaak et al. 1992) into a single dataset. Fully weighted simultaneous fit of the P-V-T-K data has been performed using a thermal pressure Equation of State with a 3rd order Tait EoS, allowing 332 simultaneous refinement of V_{00} , K_{00} , $K_{00'}$, α_0 and θ_E . As expected, the resulting coefficients V_{00} = 333 1664.46(5) Å³; K_{00} =166.57(17) GPa; K_{00} '=4.96(7); $\alpha_{(298K, 1bar)}$ = 2.09(2)*10⁻⁵ K⁻¹, with $\theta_{\rm E}$ = 512(19)K 334 $(\gamma_0 = \gamma = 1.22)$ agree within 2 σ to those from *P*-*V* and *T*-*V*-*K* separately (see Table 4). At this point it 335 is important to point out that the main difference between the *P*-*V*-*T*-*K* and *T*-*V*-*K* fits can be seen in 336 the description of the bulk moduli as a function of T (Figure 5a). It should be clear from Figure 5a 337 338 that while in the thermal regime between room-T and the maximum temperature reached with the experiments the bulk moduli calculated from the T-V-K and P-V-T-K Equations of State agree 339 within mutual uncertainties, in the low-T regime much larger discrepancies are observed. These 340 341 discrepancies are due to the different values of $\theta_{\rm E}$ in the two EoS.

Simultaneous refinement of the *P-V-T* data alone yielded the following coefficients $V_{00} =$ 1664.44(6) Å³; K_{00} =166.19(1.25) GPa; K_{00} '=5.14(40); $\alpha_{(298K, 1bar)} = 2.09(2)*10^{-5} \text{ K}^{-1}$, with θ_{E} = 512K (not refined). Standard deviations on the bulk modulus ($K_{0\text{T}}$) and its first derivative (K') obtained Page **13** of **23** from the *P*-*V*-*T* fitting are 6-8 times larger than those from the *P*-*V*-*T*-*K* ones and, when considered, the resulting coefficients from the *P*-*V*-*T* fit agree within 1 sigma with those obtained from the *P*-*V*-*T*-*K* fits. The *P*-*V*-*T* results are not showed in Figure 5a for sake of clarity, however, the resulting bulk moduli as a function of *T* agrees with the experimental data within 1 σ .

349

4. DISCUSSION

350 *4.1 Thermal expansion: comparison with literature data*

The measured room temperature-pressure volumes (V_{00}) reported in this study are in good 351 agreement with those previously reported in literature by several authors (see Table 4a and b). So 352 353 far, only few X-ray diffraction thermal expansion data are available for grossular garnet (e.g. Skinner 1956; Bosenick and Geiger 1997; Thieblot et al. 1998; Du et al. 2015) and most of them 354 have been obtained by means of powder X-ray diffraction. Data from Du et al. (2015) and those by 355 356 Skinner (1956) have been re-fitted to their entire temperature range 296 K - 859 K (excluding two outliers at 393 K and 477 K) and 284 K -980 K, respectively. On the other hand, data from Thieblot 357 et al. (1998) have been fitted from 300 K to 1453 K (i.e. well below the decomposition temperature) 358 excluding a few outliers at 607 K, 724 K, and 1241 K. The results of all the extrapolations at 300 K 359 are reported in Table 4b and shown Figure 2. As shown in Table 4b, there are considerable 360 differences between our data and those available in literature. These can be most likely ascribed to 361 the compositional differences between the various samples and/or to the uncertainties given by the 362 different experimental methods. For example, both Skinner (1956) and Thieblot et al. (1998) used 363 364 grossular-rich samples for their investigation. For sake of comparison, regardless the compositional differences, these literature data have been re-fitted to Fei-type EoS (Fei 1995). Results from the 365 fitting are reported in Table 4b, and Figure 2. 366

As shown in Figure 2, our data for the low-*T* regime are in good agreement with those reported by Bosenick and Geiger (1997) if their most scattered data (e.g. 195, 220 and 295 K) are not considered.

The thermal expansion coefficients at 300 K from both Skinner (1956) and Du et al. (2015) 370 are smaller than those calculated from our fitting. The discrepancies between our dataset and that by 371 Skinner (1956) could be ascribed to the differences in composition as their sample referred to as a 372 'grossularite' has not been properly characterized. The thermal coefficient value obtained fitting the 373 data by Du et al. (2015) is not reliable as demonstrated by the negative thermal expansion at high-T374 (regardless of the Equation of State model chosen). The same is true also for the data reported by 375 Thieblot et al. (1998), where the thermal expansion value decreases with increasing temperature. 376 377 Our datasets, ranging from 100 K to about 1000 K, allowed fitting of a Kroll-type EoS (Kroll et al. 2012), leading to a much more reliable description of the thermal behaviour of grossular. The 378 reliability of the Einstein temperatures used (θ_E =569(13) K) have been confirmed by comparison 379 with the saturation temperature (where $T_{sat} \approx \theta_E/2$) obtained from the fitting of the low-T data alone 380 (T_{sat}=270(6) K) using a Salje-type EoS (Salje et al. 1991). The agreement within 1 sigma is also 381 achieved when comparing with the θ_E from the *P-V-T-K* fit [θ_E =512(19) K, see Table 4c] 382

383

384 4.2 Compressibility: comparison with literature data

The elastic coefficients for grossular have been determined in several studies using different methods (e.g. ultrasonic wave velocity, static compression experiments in diamond-anvil cell, abinitio calculations) as summarised in Table 4a. For purpose of comparison with the literature data, the adiabatic bulk modulii (K^{S}) obtained by means of ultrasonic wave velocity techniques have been converted to K^{T} .

390 DFT calculations of a P(V) curve, fitted by a third order Birch-Murnaghan EoS, provide a 391 value of 156 GPa for the bulk modulus K^T at 300K, and a K' of 4.22. The *static* (no zero point and 392 thermal effects included) and the *athermal* (zero point corrected and a temperature of 0K) values 393 are respectively 163.9 and 160.1 GPa (K' fixed at 4.22: the value we obtained at 300K). These 394 results are in line with the *standard* performances of DFT-GGA Hamiltonians like PBEsol, as they 395 generally underestimate bulk moduli by about 10 GPa at room temperature. Another DFT study 396 **15** of **23**

from the literature (Erba et al. 2014) provided a somewhat higher bulk modulus (171 GPa), but that 396 was a result from a static calculation with a hybrid HF/DFT (B3LYP) Hamiltonian which includes 397 an Hartree-Fock correction to the DFT-GGA electronic exchange functional. Indeed, such 398 correction proved to be crucial to compute bulk moduli (and vibrational frequencies) very close to 399 400 the experimental data (Prencipe et al. 2009, 2011, 2012, 2014), but it is computationally very costly if a plane-wave basis set is used, as in the present case, and therefore it cannot be afforded. The 401 largest discrepancy in grossular bulk modulus value is with respect to the results obtained by Kawai 402 403 and Tsuchiya (2012) with a lattice dynamics approach together in combination with a LDA 404 functional. Furthermore, the first derivative of the bulk modulus (K') which we could now reliably refine clearly demonstrates that the correct value must be very close to 5 in agreement with some of 405 406 the previous studies (see Table 4a).

407

408 *4.3 Full elastic behaviour (P-V-T-K): comparison with literature data*

The new utility available in EoSFit7c allowed us to also apply the same method described in 409 410 the sections above (P-V-T-K fitting) to the available literature data obtained from independent 411 elastic measurements and PVT datasets (Kono et al. 2010; Gwanmesia et al. 2014). It is remarkable the extra constraint given by the simultaneous weighted fit of the K data (see results in Table 4c) 412 which however applied to the literature dataset does not allow to obtain the same quality as for our 413 414 elastic coefficients due to the lack of data in the low-temperature regime. This combined with the much lower accuracy and precision that in most cases arises by the limitation of the P-V-T415 simultaneous measurements (i.e. pressure calibrant, pressure transmitting media etc.) give rise to 416 discrepancies such as those showed in Figure 5b. In this case the discrepancies in bulk modulus 417 values at high-T reach up to 4% (more than 6 GPa at 1400 K). In order to crosscheck the 418 measurements with the DFT results, we performed fitting of the DFT data in the temperature range 419 between 400 and 800K using a thermal pressure equation of state combined with a 3rd order Tait-420 eos. For purpose of comparison with the other fits, bulk moduli values have been normalized to the 421 Page **16** of **23**

bulk moduli extrapolated at 0K from the P-V-T-K fit of the experimental data (Fig. 5a, red line).
The trend of the two fits agrees within 1σ thus confirming the reliability of our fitting results via PV-T-K fits even at the highest temperature where the DFT approach leads to the smallest deviations
from the experimental data.

426

4. CONCLUSIONS AND IMPLICATIONS

It is obvious that pressure-volume or temperature-volume measurements alone do not 427 provide strong constraints on the cross-terms in P-T space (e.g. $\partial K/\partial T$ or $\partial \alpha/\partial P$) which are 428 429 fundamental for the description of the elastic properties of a material at simultaneous high-P and high-T. Limitation in experimental measurements and/or in the description of the pressure-volume-430 temperature data (e.g. incorrect Equations of State, erroneous mis-use of Voigt, Reuss isothermal or 431 432 adiabatic bulk moduli) has often in the past lead to inaccurate unreliable or even completely wrong results (e.g. negative thermal expansion, erroneous $\partial K/\partial T$ etc...) specially when the EoS are 433 434 subsequently extrapolated to high-P,T.

Pressure-volume EoS are well-developed and, if correctly used, are capable of reproducing 435 436 the isothermal volume or density changes of materials to within the experimental uncertainties. 437 Previous versions of the EoSFit program were able to fit isothermal EoS, thermal expansion models and the thermal-pressure (Holland and Powell 2011) to P-V, V-T or P-V-T data, by the method of 438 least-squares. We have now extended the capabilities of EosFit7c to also fit all of these forms of 439 440 EoS to bulk and linear moduli, either alone or in combination with volume or unit-cell parameter data. With this approach we can now provide a reliable description of the variation of K' with T thus 441 avoiding mistaken extrapolations leading to non-physical behaviour as often occurs for several 442 materials (Helffrich and Connolly 2009). 443

The consistency of the results obtained with this new fitting procedure has been shown by comparison of results from using adiabatic and isothermal moduli. It has also been demonstrated the importance of performing fitting with this procedure in terms of both accuracy and precision of the thermoelastic coefficients when the *PT* cross-terms ($\partial K/\partial T$ or $\partial \alpha/\partial P$) are extrapolated in *PT* space Page **17** of **23**

448	(see Figure 5b). Additionally, thermodynamic functions were calculated for grossular using DFT
449	approach within QHA. The obtained results are in fairly good agreement with experiment, when
450	allowing for the known short-comings of the functionals employed in the DFT.
451 452 453	
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461 462	REFERENCES
463 464	Alvaro, M., Angel, R.J., Marciano, C., Milani, S., Scandolo, L., Mazzucchelli, M.L., Zaffiro, G.,
465	Rustioni, G., Briccola, M., Domeneghetti, M.C., and others (2015) A new micro-furnace for in
466	situ high-temperature single-crystal X-ray diffraction measurements. Journal of Applied
467	Crystallography, 48, 1192–1200.
468	Anderson, D.L. (1995) Equations of State of Solids for Geophysics and Ceramic Science. Oxford
469	University Press. Oxford, UK.
470	Anderson, O.L. (1996) Anharmonicity of forsterite and the thermal pressure of insulators.
471	Geophysical Research Letters, 23, 3031–3034.
472	Angel, R.J. (2000) Equations of state. In R.M. Hazen and R.T. Downs, Eds., Reviews in
473	Mineralogy: High temperature and high pressure crystal chemistry. Mineralogical Society of
474	America, Washington, D.C.
475	Angel, R.J., and Finger, L.W. (2011) SINGLE: a program to control single-crystal diffractometers.
476	Journal of Applied Crystallography, 44, 247–251.
477	Angel, R.J., Allan, D.R., Milletich, R., and Finger, L.W. (1997) The use of quartz as an internal

- 478 pressure standard in high-pressure crystallography. Journal of Applied Crystallography, 30,
 479 461–466.
- 480 Angel, R.J., Bujak, M., Zhao, J., Gatta, G.D., and Jacobsen, S.D. (2007) Effective hydrostatic limits
- 481 of pressure media for high-pressure crystallographic studies. Journal of Applied
 482 Crystallography, 40, 26–32.
- Angel, R.J., Jackson, J.M., Reichmann, H.J., and Speziale, S. (2009) Elasticity measurements on
 minerals: a review. European Journal of Mineralogy, 21, 525–550.
- Angel, R.J., Gonzalez-Platas, J., and Alvaro, M. (2014) EosFit7c and a Fortran module (library) for
 equation of state calculations. Zeitschrift fur Kristallographie, 229, 405–419.
- Angel, R.J., Nimis, P., Mazzucchelli, M.L., Alvaro, M., and Nestola, F. (2015a) How large are
 departures from lithostatic pressure? Constraints from host-inclusion elasticity. Journal of
 Metamorphic Geology, 33, 801–813.
- Angel, R.J., Alvaro, M., Nestola, F., and Mazzucchelli, M.L. (2015b) Diamond thermoelastic
 properties and implications for determining the pressure of formation of diamond-inclusion
 systems. Russian Geology and Geophysics, 56, 211-220.
- Ashley, K.T., Darling, R.S., Bodnar, R.J., and Law, R.D. (2015) Significance of "stretched" mineral
 inclusions for reconstructing P–T exhumation history. Contributions to Mineralogy and
 Petrology, 169, 1–9.
- Baroni, S., De Gironcoli, S., Dal Corso, A., and Giannozzi, P. (2001) Phonons and related crystal
 properties from density-functional perturbation theory. Reviews of Modern Physics, 73, 515–
 562.
- Blöchl, P.E. (1994) Projector augmented-wave method. Physical Review B, 50, 17953–17979.
- 500 Boehler, R. (1982) Adiabats of quartz, coesite, olivine, and magnesium oxide to 50 kbar and 1000
- 501 K, and the adiabatic gradient in the Earth's mantle. Journal of Geophysical Research: Solid 502 Earth, 87, 5501–5506.
- 503 Boehler, R., and Ramakrishnan, J. (1980) Experimental results on the pressure dependence of the Page **19** of **23**

- 504 Gruneisen parameter: a review. Journal of Geophysical Research, 85, 6996–7002.
- Bosenick, A., and Geiger, C.A. (1997) Powder X ray diffraction study of synthetic pyrope-grossular
 garnets between 20 and 295 K. Journal of Geophysical Research-All Series, 102, 22.
- Bosenick, A., Geiger, C.A., and Cemič, L. (1996) Heat capacity measurements of synthetic pyropegrossular garnets between 320 and 1000 K by differential scanning calorimetry. Geochimica et
 Cosmochimica Acta, 60, 3215–3227.
- 510 Du, W., Clark, S.M., and Walker, D. (2015) Thermo-compression of pyrope-grossular garnet solid
 511 solutions: Non-linear compositional dependence. American Mineralogist, 100, 215–222.
- 512 Erba, A., Mahmoud, A., Orlando, R., and Dovesi, R. (2014) Elastic properties of six silicate garnet
- end members from accurate ab initio simulations. Physics and Chemistry of Minerals, 41, 151–
 160.
- Fei, Y. (1995) Thermal Expansion. In Mineral Physics & Crystallography: A Handbook of Physical
 Constants pp. 29–44. American Geophysical Union.
- Freund, J., and Ingalls, R. (1989) Inverted isothermal equations of state and determination of B0,
 B'0 and B0. Journal of Physics and Chemistry of Solids, 50, 263–268.
- Gwanmesia, G.D., Wang, L., Heady, A., and Liebermann, R.C. (2014) Elasticity and sound
 velocities of polycrystalline grossular garnet (Ca3Al2Si3O12) at simultaneous high pressures
 and high temperatures. Physics of the Earth and Planetary Interiors, 228, 80–87.
- Helffrich, G., and Connolly, J.A.D. (2009) Physical contradictions and remedies using simple
 polythermal equations of state. American Mineralogist, 94, 1616–1619.
- Holland, T.J.B., and Powell, R. (2011) An improved and extended internally consistent
 thermodynamic dataset for phases of petrological interest, involving a new equation of state for
 solids. Journal of Metamorphic Geology, 29, 333–383.
- Isaak, D.G., Anderson, O.L., and Oda, H. (1992) High-temperature thermal expansion and elasticity
 of calcium-rich garnets. Physics and Chemistry of Minerals, 19, 106–120.
- 529 Kawai, K., and Tsuchiya, T. (2012) First principles investigations on the elasticity and phase Page 20 of 23

- 530 stability of grossular garnet. Journal of Geophysical Research: Solid Earth, 117, 1–8.
- King, H.E., and Finger, L.W. (1979) Diffracted Beam Crystal Centering and Its Application to
 High-Pressure Crystallography. Journal of Applied Crystallography, 12, 374–378.
- Klotz, S., Chervin, J.C., Munsch, P., and Marchand, G. Le (2009) Hydrostatic limits of 11 pressure
 transmitting media. Journal of Physics D: Applied Physics, 42, 75413.
- Kono, Y., Gréaux, S., Higo, Y., Ohfuji, H., and Irifune, T. (2010) Pressure and temperature
 dependences of elastic properties of grossular garnet up to 17 GPa and 1 650 K. Journal of
 Earth Science, 21, 782–791.
- Kresse, G., and Furthmüller, J. (1996) Efficient iterative schemes for ab initio total-energy
 calculations using a plane-wave basis set. Physical Review B, 54, 11169–11186.
- Kroll, H., Kirfel, A., Heinemann, R., and Barbier, B. (2012) Volume thermal expansion and related
 thermophysical parameters in the Mg, Fe olivine solid-solution series. European Journal of
 Mineralogy, 24, 935–956.
- 543 Milani, S., Nestola, F., Alvaro, M., Pasqual, D., Mazzucchelli, M.L., Domeneghetti, M.C., and 544 Geiger, C.A. (2015) Diamond-garnet geobarometry: The role of garnet compressibility and 545 expansivity. Lithos, 227, 140–147.
- 546 Miletich, R., Allan, D.R., and Kuhs, W.F. (2000) High-pressure single-crystal techniques. High547 Temperature and High-Pressure Crystal Chemistry, 41, 445–519.
- 548 Nye, J.F. (1985) Physical Properties of Cristals: Their Representation by Tensors and Matrices.
 549 Oxford University Press.
- 550 Orear, J. (1982) Least squares when both variables have uncertainties. American Journal of Physics,
 551 50, 912.
- Parlinski, K., Li, Z., and Kawazoe, Y. (1997) First-Principles Determination of the Soft Mode in
 Cubic ZrO2. Physical Review Letters, 78, 4063–4066.
- Perdew, J., Ruzsinszky, A., Csonka, G., Vydrov, O., Scuseria, G., Constantin, L., Zhou, X., and
- 555 Burke, K. (2008) Restoring the Density-Gradient Expansion for Exchange in Solids and Page 21 of 23

- 556 Surfaces. Physical Review Letters, 100, 136406.
- Prencipe, M., Noel, Y., Bruno, M., and Dovesi, R. (2009) The vibrational spectrum of lizardite-1T
 [Mg3Si 2O5(OH)4] at the Γ point: A contribution from an ab initio periodic B3LYP
 calculation. American Mineralogist, 94, 986–994.
- 560 Prencipe, M., Scanavino, I., Nestola, F., Merlini, M., Civalleri, B., Bruno, M., and Dovesi, R.
- (2011) High-pressure thermo-elastic properties of beryl (Al4Be6Si12O36) from ab initio
 calculations, and observations about the source of thermal expansion. Physics and Chemistry
 of Minerals, 38, 223–239.
- Prencipe, M., Mantovani, L., Tribaudino, M., Bersani, D., and Lottici, P.P. (2012) The Raman
 spectrum of diopside: a comparison between ab initio calculated and experimentally measured
 frequencies. European Journal of Mineralogy, 24, 457–464.
- Prencipe, M., Bruno, M., Nestola, F., De La Pierre, M., and Nimis, P. (2014) Toward an accurate ab
 initio estimation of compressibility and thermal expansion of diamond in the [0, 3000 K]
 temperature and [0, 30 GPa] pressures ranges, at the hybrid HF/DFT theoretical level.
 American Mineralogist, 99, 1147–1154.
- Ralph, R.L., and Finger, L.W. (1982) A Computer-Program for Refinement of Crystal Orientation
 Matrix and Lattice-Constants from Diffractometer Data with Lattice Symmetry Constraints.
 Journal of Applied Crystallography, 15, 537–539.
- Salje, E.K.H., Wruck, B., and Thomas, H. (1991) Order-parameter saturation and low-temperature
 extension of Landau theory. Zeitschrift für Physik B Condensed Matter, 82, 399–404.
- 576 Skinner, B.J. (1956) Physical properties of end-members of the garnet group. American
 577 Mineralogist, 41, 428–436.
- 578 Thieblot, L., Roux, J., and Richet, P. (1998) High-temperature thermal expansion and
 579 decomposition of garnets. European Journal of Mineralogy, 10, 7.
- Togo, A., and Tanaka, I. (2015) First principles phonon calculations in materials science. Scripta
 Materialia, 108, 1–5.