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Evidences of volcanic unrest on high-temperature fumaroles by satellite thermal monitoring: The case of Santa Ana volcano, El Salvador

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(Article begins on next page)

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Title: The effects of environmental parameters on diffuse degassing at Stromboli volcano: insights from joint monitoring of soil CO2 flux and radon activity

Article Type: Research Paper

Keywords: Stromboli volcano; Continuous geochemical monitoring; Soil CO2 flux; radon activity; Environmental parameters; Time series analyses.

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Abstract: Soil CO2 flux and 222Rn activity measurements may positively contribute to the geochemical monitoring of active volcanoes. The influence of several environmental parameters on the gas signals has been substantially demonstrated. Therefore, the implementation of tools capable of removing (or minimizing) the contribution of the atmospheric effects from the acquired timeseries is a challenge in volcano surveillance. Here, we present four years-long continuous monitoring (from April 2007 to September 2011) of radon activity and soil CO2 flux collected on the NE flank of Stromboli volcano. Both gases record higher emissions during fall-winter (up to 2700 Bq*m-3 for radon and 750 g m-2 day-1 for CO2) than during spring-summer seasons. Short-time variations on 222Rn activity are modulated by changes in soil humidity (rainfall), and changes in soil CO2 flux that may be ascribed to variations in wind speed and direction. The spectral analyses reveal diurnal and semidiurnal cycles on both gases, outlining that atmospheric variations are capable to modify the gas release rate from the soil. The long-term soil CO2 flux shows a slow decreasing trend, not visible in 222Rn activity, suggesting a possible difference in the source depth of the of the gases, CO2 being deeper and likely related to degassing at depth of the magma batch involved in the February-April 2007 effusive eruption. To minimize the effect of the environmental parameters on the 222Rn concentrations and soil CO2 fluxes, two different statistical treatments were applied: the Multiple Linear Regression (MLR) and the Principal Component Regression (PCR). These approaches allow to quantify the weight of each environmental factor on the two gas species and show a strong influence of some parameters on the gas transfer processes through soils. The residual values of radon and CO2 flux, i.e. the values obtained after correction for the environmental influence, were then compared with the eruptive episodes that occurred at Stromboli during the analysed time span (2007-2011) but no clear correlations emerge between soil gas release and volcanic activity. This is probably due to i) the distal location of the monitoring stations with respect to the active craters

and to ii) the fact that during the investigated period no major eruptive phenomena (paroxysmal explosion, flank eruption) occurred. Comparison of MLR and PCR methods in time-series analysis indicates that MLR can be more easily applied to real time data processing in monitoring of open conduit active volcanoes (like Stromboli) where the transition to an eruptive phase may occur in relatively short times.





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Prof. A. Aiuppa Editor at JVGR

Dear Editor,

We are submitting the revised copy of the manuscript "The effects of environmental parameters on diffuse degassing at Stromboli volcano: insights from joint monitoring of soil CO_2 flux and radon activity".

As suggested by the Referees, the paper has been reorganised: Chapter 4 is now including four subchapters investigating more deeply the relationships between gases and environmental parameters.

Two new figures have been added: Fig. 5 showing the influence of wind (direction and speed) on soil CO_2 flux, and Fig. 6 which summarizes the results of spectral analyses on raw data (1 year of hourly measurements) for soil CO_2 flux, ²²²Rn activity, air T and atmospheric P.

Figure 10 has been modified to include results of both the statistical treatments used.

We added in Table 3 the main parameters of the MLR analysis and the Table 4 has been modified to better explain the results of PCR statistical method. We also provided additional technical information of the two monitoring stations in the S1 Table for Supplementary Material.

Abstract, Discussion and Conclusions have been revised, as suggested by the Referees, to make immediately clear the scope of the paper and the obtained results.

We hope that the paper can now better satisfy the JVGR standards.

Sincerely, Marco Laiolo & co-workers

Torino February 03, 2016

Response to Reviewers (cited lines refer to the New Manuscript)

Referee # 1

- From a methodological point of view, I suggest the authors to add a table with the technical characteristics of the various sensors of the two stations (it can be added as supplementary material). Precision and accuracy of the measurements should be added.

As suggested, we added in Supplementary Material a Table (S1) with the technical characteristics, accuracy and sensitivity of the sensors of the two fully-automated stations.

- From an organizational point of view, I would suggest that chapter 4 was named "Results" with subchapters.

As suggested we changed the title of Chapter 4 and inserted four sub-chapters [line 205 – line 379].

- In page 12 authors mention that daily temperature variations have a weak influence on the gas flux, highlighting the correlation between wind and CO2 fluxes (also based on results obtained in previous works). However, looking at table 2, the correlation coefficient between CO2 flux and air temperature is higher than the coefficient between CO2 flux and wind speed. Authors should revisit these sentences and better discuss this aspect.

We rewrote this point trying to clarify this aspect (see Chapter 4.1 and 4.2). The relationship between CO_2 flux and wind has been more properly discussed (line 261 - 267 and Fig. 5). We performed a spectral analysis on both gases and on the main atmospheric variables (Air T and P) in order to outline the diurnal and semi-diurnal cycles of the analysed gas [line 278 – line 287].

- Pages 13 and 14 - Figure 6 is quite interesting and shows variations on the coefficients of correlation for different seasons. I would suggest the authors to add a shadow area in Figure 6 highlighting the first subset that authors refer "is somehow different compared to all the others". I suggest to better specify to which period authors refer to and to write the main differences (page 14, 3rd and 4th lines), as they are not easily observed just looking at the figure. Other aspect from Figure 6 is associated with the correlation established between CO2 flux and wind direction. Since the wind direction is a "circular" variable, how do the authors define the various wind quadrants? Do you define intervals or use the absolute values? In this last case 0 and 359 are almost the same, but will have a different weight for the correlation and for the regression and principal components models. Authors need to clarify how do they used this variable.

As suggested we outlined the subset SS 2007 with a grey field in Fig. 6 (now Fig. 8). We added a new figure (Fig. 5), in Chapter 4.1., to show the wind influence (both speed and direction) on soil CO_2 flux. Moreover, we removed the variations of correlation coefficient between soil CO_2 flux and wind direction and speed. The question of wind influence is now more clearly discussed in chapter 4.1 (line 261 - 267).

- Authors analyse data obtained in the soil surface (soil CO2 flux station) and at about 1 m depth (soil radon concentration station), and this can result in different effects of the environmental parameters,

just depending on the site of the measurement (for instance the soil water content). I suggest the authors to add some discussion about this aspect.

We have inserted some sentences discussing the difference of the two stations [line 187 - 198] and the related influence of environmental parameters [line 233 - 267]. We have explained that the applied methods represent the state of the art for measuring soil CO₂ flux and ²²²Rn concentration.

- On the chapter related with the statistical treatment, I suggest the authors to add some references that support the mathematical modelling. The reference "Hernandez et al., 2004" (page 15, line 2) refers to a work where this methodology was applied but do not explain the mathematical criteria. In addition, a previous work of Granieri et al. (2003 - EPSL) is, as far as I know, the first that applied regression analyses to the soil CO2 flux time series and should be mentioned. Those authors also discuss why selected regression models instead of the principal components analyses, aspect that should be here considered and discussed.

Following the suggestions of the referee, we inserted in Table 3 the variables used and the retrieved MLR equation by using the appropriate independent variables (i.e. environmental parameters) for both gases. Text has been adjusted accordingly [line 335 – 340]. Also Table 4 has been modified to better explain the PCR method.

- One question that was not clear to me was if the authors applied the statistical analyses to the raw data (hourly measurements) or to the daily values? This aspect needs to be clarified. If the statistical approach is done to the raw values, which in my opinion is the more adequate, did authors check also if there are daily variations on the gas time series? For the other side, if the statistical approach is applied to the daily averages (removing automatically the existence of daily cycles) I would suggest authors to add a sentence specifying that and to give some explanation how do they use the information of the wind speed for instance. Is it the average value selected, or the maximum speed of the day?

The statistical treatment was applied to the hourly raw values (wind included), as supposed by referee, and we have explained this aspect in the revised version of the paper [line 321 - 323].

- As already mentioned authors used two different analyses to filter external influences from the gas data. Authors try also to discuss some differences between the two approaches used. In what concerns MLR authors mention (page 15) that MLR takes into account only the factors that are "more correlated" with CO2 and radon. What was the criterion used? Granieri et al. (2003) applied the procedure of Garside to select the variables; Viveiros et al. (2008; 2015 - JVGR and GSL) and Silva et al. (2015 - EPJ) selected only the variables that caused an increment on the R2 higher than 1%. A sentence about the criteria used need to be added. In what concerns the statistical methodology, I strongly suggest the authors to add the final regression models for the CO2 fluxes and radon concentrations. Same comment to the PCR models that need to be shown. The created variables (PCR method) that will correlate with the gas data are requested. These models can be added as tables or as equations, and are mandatory to understand which variables were added to the model, and which was the influence of each one.

We accepted the suggestions of the referee and adjusted the paper accordingly (see revised chapter 4.4; line 320 - 379).

- One aspect that also called my attention and I would like to see some discussion about is the fact that authors show that correlation factors change depending on the seasons. Would it be wise to produce

different regression models considering different seasons of the year? One additional point is related with the fact that some variables may have different influences on the gas data depending on different intervals of values, i.e., variables may work as second order polynomial and not as "simple" linear variables. Authors did not consider this possibility and evaluated only the linear correlations, what is acceptable, however, I would suggest to be more careful when in page 15 they refer: "...major fluctuations that obviously cannot be due to environmental variations" (line 17). Considering that some environmental parameters may not have only a straight linear influence on the gas flux, some of the explanation may be attributed to that aspect, thus I would recommend the authors to use the word "suggest" instead of "obviously".

We changed the sentence as suggested and added some sentences in chapters 4.3 and 5 [see line 383 – 385]. Additionally, regarding variations exhibited by radon residuals, we included a short discussion on the sensitivity of the radon sensor [line 389 – 395].

- Considering that MLR explains more gas variations than PCR (table 3), I do not understand why authors decided to use the standard residuals from the PCR to compare with the volcanic activity (Fig. 8). I would expect exactly the opposite. Authors should better explain this decision that has impact in the final discussion/observations. Authors mention that "PCR ensures a more accurate estimation of warning thresholds..." (pags 18, line 9), but in my opinion this sentence needs to be better explained.

Following the Referee suggestion, in the revised text we included the analyses of MLR residuals as well (see modified Fig. 10) and tried to better explain the reasons for preferring MLR for a quick monitoring application and PCR for a more accurate post-processing of data [see Chapter 5 and line 474 -481].

- In the abstract authors mention that no clear correlation emerges between soil gases and volcanic activity, which I agree. In fact, from the analyses of Fig. 8, some periods with anomalous CO2 or radon values are contemporaneous with the volcanic activity, but other periods do not show any relation. A pattern in the bahaviour of the CO2 fluxes - radon activity and volcanic activity is not established. Authors should better emphasize this, since in some paragraphs of sub-chapter 6.2, as well as in the conclusions, authors seem to highlight some correlations. Potential explanations for the differences observed between CO2 flux and radon data would also be welcome, as well as some attempt to explain why alternation between positive and negative anomalies is observed (essentially in the radon time series). Maybe authors have some idea why this occur and could be an interesting step forward to understand the subsuperficial phenomena.

The abstract, discussion and conclusions have been adequately reviewed following the suggestions in order to improve the comparison between our results and the Stromboli volcanic activity during the analysed period. We have also included an inference on a possible difference depth for Rn and CO_2 sources [line 396 – 401; line 433 – 437 and line 453 – 458].

I suggest the authors to read the recent work of Viveiros et al. (2014 - JGR), where similar seasonal trends were highlighted in the soil CO2 flux time series of data recorded at Furnas Volcano (S. Miguel Island). In addition, those authors, together with Rinaldi et al. (2012 - JGR), identified diurnal variations on the CO2 flux time series and modelled these variations with a geothermal simulator. In my opinion the results obtained by those authors can be useful to explain some of the observations here presented; in addition, those authors applied quite a similar filtering process to the data that is also confirmed in this

work (similar to Fig. 8). I suggest the authors to read the modelling done considering the thermal gradients, what is explained in the present work as a major cause for the fluctuations observed.

We acknowledge the referee for this suggestion that it allowed to describe more properly the capability of atmospheric variables to modulate gas release from soil [see subchapter 4.2].

Reviewer #2: General comments

In this manuscript, the authors show a long time data series for two different gas species, consisting of continuous monitoring of Rn concentration and soil CO2 flux at Stromboli volcano between April 2007 and September 2011. The knowledge of the degassing processes in active volcanoes are of great interest for the volcanology community. In my opinion, the topics covered might be considered for a new contribution to the journal, however, some concerns regarding the manuscript should be considered before publication.

First, I suggest a revision of both the abstract and conclusion in order to make the novelty of the work and the findings immediately clear. Moreover, much of the original data considered in the paper are still lacking in clarity and scope. Specifically, section 4 is often redundant and the narrative explaining the correlation between the meteorological parameters and the Rn concentration and CO2 fluxes is confusing, despite the important role the authors seem to give this in the paper. Finally, I strongly recommend reviewing the conclusion. In my opinion, rather than focusing on the general literature published on Stromboli, the authors could be much more focused in demonstrating and underlining the relevance and importance of the results of their work.

We carefully revised in detail abstract, discussion and conclusions in order to focus better on the results obtained in this work and, specifically, highlighting the contribution that our results can give to the improvement of geochemical monitoring of active volcanoes.

I, therefore, recommend a major revision before publication, including a thorough proofreading in terms of use of English and general text cohesion.

Following the recommendations and previous suggestions, the text has been deeply revised and partly reorganized, avoiding repetitions and giving attention to its cohesion. We hope that it can now be accepted.

The following are specific comments for the improvement of the manuscript and requests for better explanation:

Pag. 3 "Due to its short half-life (t1/2=3.82 days)"

It should be specified that this refers to 222Rn

We changed the sentence according to this comment [line 60].

Pag 3 "In volcanic areas, the combined surveys of soil 222 Rn emissions and CO₂ flux give us the opportunity to track diffuse degassing and related processes (Giammanco et al., 2007)."

I suggest exploring this specific aspect in your work much more as it seems to be relevant to the topics within this paper and could also be reasonably considered one of the main arguments justifying the work.

The combined measurements of soil CO₂ flux and ²²²Rn activity and specifically the ²²²Rn/²²⁰Rn ratio are mainly employed in order to reveal sites characterized by deep magmatic degassing (Giammanco et al., 2007). Following the suggestion we presented a more deep analysis of the CO2 flux and Rn activity reaching the conclusion that the two gases have likely a different source [line 396 – 401; line 433 - 437 and line 453 - 458].

Pag 6 "Since then, the scientific community acquired a tremendous amount of geophysical, geochemical and geodetic data were collected during the most recent eruptions 2007 and 2014 (Barberi et al., 2009; Rizzo et al., 2014)."

Grammar: Please insert "which"

The sentence was modified according to this comment [line 134 – 138].

Pag 8 "Therefore, two additional soil CO2 flux and multiparametric fully automated stations were installed along the ENE flank of the volcano (respectively at Rina Grande and Nel Cannestrà) where gaseous emissions are higher (Carapezza et al., 2009) (Fig. 2b)."

I suggest including in the Fig 2 b an inset showing the location of the stations on the island in order to identify the positions of the stations. In addition, it would be interesting to indicate the scale in order to recognize the "order of density" of the sampling shown in the maps both in fig. 2a and 2b.

In Figure 2a and 2b the location of the automated 222 Rn (2a) and soil CO₂ flux (2b) stations is shown. We added scale in both sub-Figures, as suggested.

Pag 8 "The gas automated stations consist of two units, one for measuring the isotope of the radon progeny ..."

I suggest that the authors make a link with the previous paragraphs. It seems that there are four monitoring stations in operation, but in the following text you consider only one. It would be useful to understand why and what has motivated the choice. In addition, it is not clear which monitoring station is used by the authors; is it the station Nel Canestrà (as it seems indicated in tab 2), or at Rina Grande?

We clarified this point [line 163 – 165 and line 173 – 175] and clearly outlined (Figs. 1 and 2 and Tab. 1) the dataset analysed in this work.

Pag 9 "A preliminary analysis of the trend of environmental parameters, supported by the correlation factors reported in Table 2 ..."

The authors should make clear what kind of "correlation factors" are considered between the parameters investigated (r, R2, CV ...), also in table 2. In addition, in table 2, second line, the names in the column related to the CO2 monitoring station are not visible.

We changed correlation factors with correlation coefficients [line 226] and remarked in the text that in Table 2 we show the coefficient R [line 303].

Pag 10 "The wind speed signal is ... (Fig. 4b)."

It should be Fig. 4e.

This sentence has been removed.

Pag 10 "Radon concentrations (see Fig. 3a) show relatively low values (Cigolini et al., 2009)"

For the sake of clarity, the authors should supply some more information in this sentence. Is a different regime of radon activity for Stromboli stated in Cigolini et al., 2009? If it is not, then it is preferable not to add self-reference in the paper.

We clarified this sentence and in order to better understand the mean of "relatively low values" we compare the discussed values with other measures performed in different areas [line 215 and line 219 – 221].

Pag 10 "Radon emissions are strongly negatively correlated to soil temperature, as already pointed out by Cigolini et al. (2009) and Laiolo et al. (2012). "

This is redundant.

The sentence has been removed.

Pag 11 "soil and air temperatures, ... Therefore, a marked difference between ..."

This is redundant.

The sentence has been removed.

Pag 11 "... an increase in soil moisture is capable of raising the 222Rn emanation coefficient of one order of magnitude ..."

Do the authors intend perhaps to say rate?

We properly changed this sentence in order to clarify this point [line 237 – 240].

Pag 12 "...does not appear on radon long-term signals that, instead, show an excellent stability around the average of 2000 Bq/m3 (see straight line in Fig. 3a..."

Excellent? This adjective is very strong and I question its appropriacy.

The sentence has been modified.

Pag 12 "...It is important to outline that, according to this view, the concurrent measurements of CO2 and 222Rn can be considered as an important tool for detecting the variation in time of the deep vs. shallow contribution to the soil gas, as it was found in other volcanoes ..."

This is an aspect that should be explored more, and the authors should explain in what way the data shown can contribute to make that link more evident. Perhaps it is also better to move this discussion to a different section of the paper, providing space for a wider argumentation.

The revised text includes a more carefully analysis of the temporal variations of CO_2 flux and ²²²Rn activity leading to infer the deep vs shallow contribution of the two gases.

Pag 12 "...However, during a short term period, it is clear that..."

I suggest not introducing the phrase with "However" (here but also in other parts of the text and also "nevertheless"), because the contrast it introduces is not immediately obvious; i.e. the contrast refers to the topic sentence of the preceding paragraph and not its concluding sentence.

The comment has been considered in the English revision of the text.

Pag 13 "...This behaviour is evident by calculating the mean value of the whole data on each specific day of the year. The annual trend from Jan-1 to Dec-31 (Fig. 5) remarks the inverse relation between temperature and soil gas release, as well as the correlation between the two gas species..."

It would be preferable to make this sentence clearer rather than generate confusion amongst the relationship between daily variation of temperature and wind speed (previous sentence); and with the annual variation of temperature. In addition, I suggest adding the temperature in the figure 5.

We clarified the main steps of this analysis, which was applied on the raw dataset to investigate the annual trend affecting both gases [line 288 – 292]. Moreover, in the related Figure we added the trend of soil and air temperature (Fig. 7).

Pag 13 "Is also evident that the most significant day-by-day variations occur du..."

It is also ...

We corrected this sentence [line 294].

Pag 13 "In Fig. 6 we report ... both CO2 flux and radon activity display a more distinctive negative ..."

These arguments have already been considered in part, therefore the discussion in this section of the text becomes redundant and risks muddling the reader. The information included in fig. 6 should be rearticulated in a revised and synthesized text of the entire section 4.

This section of the paper has been reviewed taking into account the comments of both referees (see subchapter 4.3).

Pag 14 "In addition, seasonal variations in geochemical data and environmental correlation coefficients are somewhat different in the first subset compared to all the others."

This is not clear. Please explain better and introduce a description about what is meant by subset.

A specific comment has been introduced in subchapter 4.3 [line 315 – 319] and the first subset has been marked on Fig. 8.

Pag 14 "In this time span , there"

Please delete space

We deleted space.

Pag 15 "Particularly, soil and air temperature were considered by the regression for both gas species, together with wind conditions for CO2 ..."

Please be specific with wind "conditions": speed; directions; or both?

The question has been clarified [line 339] and we added also a new Fig. 5.

Pag 15 "Results show that the atmospheric variables taken into account for this analysis are able to predict 45% and 51% (R=0.67 and R=0.71) ..."

Is it R or R2? In table 3 it is indicated as "Determination Coefficient"

We clarified this point in the text [line 341 - 342] and in the related Tables (Table 3 - 4), in order to avoid misunderstanding.

Pag 15 "Moreover, soil CO2 flux and 222Rn values show low dispersions, respectively 4.0% and 4.8% of the computed residuals exceed the average ±2 standard deviation (Table 3)"

This should be inverted: ... respectively 4.8% and 4.0%

We corrected the sentence [line 344].

Pag 15 "By looking at the residuals, it can be seen that..."

The bell shape in the last two years does not seem to have disappeared exactly. Indeed, the high CO2 flux fluctuation in the previous years masks the bell shape in the later years (it could be made better visible using a different scale). In addition, the CO2 residual shown in the graph appears to be significantly negative for most of the time. Can you give an interpretation/explanation for this; is there an error in the right scale of the graph?

We can assure the referee that the "bell-shape" of CO_2 flux does not depend on the scale. There was actually an error on the right scale of the graph now corrected in Fig. 9 and we are grateful to the referee for allowing this correction.

Pag 16 "So, we created one dataset for radon with soil temperature and atmospheric pressure of radon station and added the other variables from the CO2 station...."

I think the authors should also indicate in the text why it is reasonable to use meteorological parameters measured from a monitoring station sited at distance; are they not far?; is there significant soil homogeneity...?

As shown in Figure 1 and 2 the two stations are actually very close, at a distance of only 30 m. Actually, only soil moisture is extrapolated for ²²²Rn station.

Pag 17 "However, it can be noted that the radon treatment provides many significant negative residual values $\geq 2\sigma$, whereas only one negative residual is recorded for soil CO2 flux (see Fig. 8)..."

The graph shows an Rn time series that seems to vary without a specific connotation but, in some way, appears as a sort of noise. Please could you be more specific about the volcanological information that we can deduct from this signal?

In the Discussion and Conclusions we addressed this topic and suggested an explanation for the radon behaviour [line 387 – 395].

Pag 18 "Comparison of the standard residuals (by PCR) for CO2 soil flux and 222Rn activity versus time (Fig. 8) shows that ... considered, only radon shows frequent positive anomalies whereas anomalous CO2 flux values are rarely recorded....."

It is hardly possible to recognize if there is any similar behaviour in the graph, both for the length of the data series and the difficulty in overlapping the two signals (fig 8a and b). In order to state that there is, and how significant the correlation may be, it would be better to divide the graph into smaller windows of time. It would help also to overlap the signals and, if possible, add some numerical correlation values (like matching positive and negative cases).

The Discussion (and figure) has been modified to address this question.

Pag 18 "Finally, the comparison between ... compare our emerging results with the 2007-2011 Stromboli volcanic activity"

It is not clear why:

- MLR seems to be a more adequate method for getting the quick results needed in near-real time volcano monitoring.

- whereas PCR ensures a more accurate estimation of warning thresholds (which thresholds ?)

Please explain

The chapter has been revised to better explain the characteristics of the two methods and their possible use [Chapter 5 and line 474 -481]. Results of both treatments have been considered for the comment on the residual values.

Pag 18 "...are slightly influenced by volcanic activity that, in the time-span considered, did not produce any relevant explosions or major lava effusion.

This section should be improved with more details in the interpretation. The authors stress that in the first two years there is a good correlation between the data recorded and the volcanic activity. But they should explain or interpret the scarce correlation from 2009 to 2011, where actually 7 major explosions and 3 lava overflows definitely confirm that Stromboli volcano was fairly active.

We modified the Discussion, as well as Conclusions and Abstract in order to remove any misunderstanding regarding the relationship between volcanic activity and gas signals obtained by using the PCR and MLR analyses. We enphasize that in the considered period volcanic activity was related only to changes in the upper plumbing system without involvement of deep gas-rich magma [line 416 – 421 and line 442 – 443].

Pag 19 "... are slightly influenced by volcanic activity that, in the time-span considered, did not produce any relevant explosions or major lava effusion."

This contradicts what is stated in the previous section and shown in fig. 8. In addition, in the recent literature and also in the references quoted in the text (Calvari et al. 2014; Rizzo et al. 2015) it is pointed out that Stromboli, in the time considered, has shown "dynamic" volcanic activities.

In the deep revision on the Discussion and Conclusions sections we carefully considered the referee comments.

Pag 19 "...As recently stressed by De Gregorio et al. (2014) ... Thus, combining CO2 flux and 222 Rn concentration measurements gave us the opportunity to better investigate the changes in volcanic activity associated with magma rise."

The link is not clear between these two sentences. Moreover, it does not seem that there has been enough investigation into what way the combination of CO2 and Rn provided information in this work regarding the volcanic activity and the "associated magma rise".

We modified this sentence [line 451 – 453].

Pag 20 "We finally emphasize that the monitored area has been characterized by marked anomalous radon emissions prior and during the paroxysmal explosion of March 15, 2007 (Cigolini et al., 2013). Therefore, it is not excluded that multiparametric geochemical monitoring may play ..."

It is surely acceptable the statement that geochemical monitoring may play a very important role in order to forecast volcanic activities, but data and argumentations supporting this aspect in this paper are still weak and necessitate major deepening.

We changed this sentence remarking the contribution given by continuous measurements in the analysed sites and also considering previous observations [line 481 – 486].

Figure captions

Fig 1 Please correct CO2 in CO₂

We changed the figure according to this comment.

Fig 2 Please correct CO2 in CO₂ flux. Please add some information in relation to the structural features shown in the maps.

We changed the figure according to this comment.

Fig 3 It seems that the circles are the raw data; please check.

As specified in the caption, we confirm that circles are the raw data.

Fig 5 Please correct CO2 in CO₂

We changed the figure (now Fig. 7) according to this comment.

Fig 7 All the graphs are in different grey gradations, but CO_2 in the legend b is in red. Please check.

We changed the figure (now Fig. 10) according to this comment.

Highlights

CO₂ flux and Rn activity were continuously measured for 4 years on Stromboli NE flank CO₂ flux and Rn are affected by environmental parameters (soil H and T, wind, air T) Two statistical methods were used to identify and remove atmospheric effects on gas In the reference period no deep gas-rich magma was involved in the eruptive activity No clear correlation was found between volcanic activity and CO₂ and radon emissions

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18	monitoring of active volcanoes. The influence of several environmental parameters on the gas	
19	signals has been substantially demonstrated. Therefore, the implementation of tools capable of	
15	signals has been substantiany demonstrated. Therefore, the implementation of tools capable of	
20	removing (or minimizing) the contribution of the atmospheric effects from the acquired	
21	timeseries is a challenge in volcano surveillance. Here, we present a four years-long period of	
22	continuous monitoring (from April 2007 to September 2011) of radon activity and soil carbon	
23	dioxideCO ₂ flux acquiredcollected on the NE flank of Stromboli volcano, from April 2007 to	Formatted: Subscript

24	September 2011. Previous soil degassing surveys on this sector allowed us to decode the
25	volcano tectonic structures where diffuse degassing is actively taking place. Radon and CO_2
26	stations were installed in a distal area from the active vents, at ~520 m a.s.l. (named Liscione or
27	Nel Cannestrà). Collected time series define a rather good correlation between the two Both
28	gases and both record higher emissions during fall-winter periods (up to 2700 Bq/m ³ *m ⁻³ for
29	Rnradon and 750 g m ⁻² day ⁻¹ for CO ₂). The ²²² Rn concentration shows a remarkable steady state)
30	than during spring-summer seasons. Short-time variations on ²²² Rn activity are modulated by
31	changes in soil humidity (rainfall), and changes in soil CO ₂ flux that may be ascribed to
32	variations in wind speed and direction. The spectral analyses reveal diurnal and semi-diurnal
33	cycles on both gases, outlining that atmospheric variations are capable to modify the gas release
34	rate from the soil. The long-term trend whereas soil and air temperatures define a typical
35	sinusoidal curve and soil humidity variations strictly control short time changes. Soil CO2 flux
36	exhibits a similar behaviour with drastic changes ascribed to wind conditions. However, the
37	long term datasets show a soil CO ₂ flux shows a slow decreasing trend, not visible in ²²² Rn
38	activity, suggesting a reduced CO ₂ contribution from a deep source (i.e. magmatic) within the
39	most recent years.
40	Correlation coefficients outline that radon variations are essentially modulated by soil
41	temperature and humidity, whereas CO2 fluxes mainly depend on soil air temperature and wind
42	speed and direction. In order to possible difference in the source depth of the of the gases, CO_2
43	being deeper and likely related to degassing at depth of the magma batch involved in the
44	February-April 2007 effusive eruption. To minimize the effect of the environmental parameters
45	on the ²²² Rn concentrations and soil CO ₂ fluxes, we applied two different statistical treatments
46	were applied: the Multiple Linear Regression (MLR) and the Principal Component Regression

47 (PCR). These approaches allow to quantify the weight of each atmospheric environmental factor on the two gas species. Both methods and show thea strong influence of some of the 48 environmental parameters on the gas transfer processes through the soil. 49

Residualsoils. The residual values of radon and CO₂ flux, i.e. the values obtained after correction 50 51 for the environmental influence, were then compared with the volcanic eruptive episodes that 52 occurred at Stromboli during the analyzed analyzed time span (2007-2011) but no clear correlation emergescorrelations emerge between soil gas release and volcanic activity. This is 53 54 probably due to *i*) the distal location of the monitoring stations inwith respect to the active craters and to *ii*) to the fact that during the investigated period is characterized by the "ordinary" 55 Strombolian activity (occurring between two eruptions, i.e., 2007 and 2014) during which no 56 major eruptive phenomena (paroxysmal explosions has been identified. 57

Results obtained by explosion, flank eruption) occurred. Comparison of MLR and PCR have been 58 evaluated methods in terms of "cost benefit" for improving time-series analyses. The 59 60 methodology hereby tested analysis indicates that MLR can be more easily applied in near-to real 61 time measurements to improve geochemical data processing in monitoring in volcanic areas of open conduit active volcanoes (like Stromboli) where the transition to mitigate volcanic riskan 62 eruptive phase may occur in relatively short times. 63

1 - Introduction

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66 Geochemical measurements and realReal-time monitoring of gas species are release (output and 67 composition) at active volcanoes is useful to identify forecast changes in volcanic activity and magma degassing. It is well known that diffuse and concentrated soil degassing may release high 68 69 amount of gases at active volcanoes. Active volcanoes are characterized by persistent huge gas

70 emissions from craters, fumaroles and also diffusively from soils (Allard et al., 1991; Burton et al., 2013; Inguaggiato et al., 2013; Burton et al., 2013), thus) and systematic gas monitoring may 71 help to detect precursory signals of incoming eruptions (e.g., Aiuppa et al., 2009; Padrón et al., 72 2013). In recent years, this approach was applied at several volcanoes to record geochemical 73 74 changes during volcanic activity, and to investigate their role before, during and after major 75 eruptive episodes (including flank instabilities; e.g., Carapezza et al., 2004 and 2009; Alparone et al. 2005; Cigolini et al., 2005). Another open and debated issue is the role of degassing prior the 76 77 onset of earthquakes (Toutain and Baubron, 1999; Salazar et al., 20022001) and during earthquake-volcano interactions including seismic-volcanic unrest (Cigolini et al., 2007; Padilla 78 79 et al., 2014).

Carbon dioxide, after water, is the most abundant volatile dissolved in magmas and, because of
its relatively low solubility in magmatic liquids, it is essentially released at higher depths and
before other gas species (Pan et al., 1991, Papale et al., 2006). Notably, measurements onof soil
CO₂ fluxes or CO₂ concentrations in volcanic plumes, are critical for detecting degassing
processes related to changes in the plumbing system of the volcano.

Radon is a noble gas, a daughter decay product of ²²⁶Ra and belongs to the ²³⁸U decay chain. Due 85 to its short half-life $(t_{1/2}=3.82 \text{ days})$ -it), ²²²Rn can be used as a tracer of both diffuse and 86 localized degassing since it can substantially be measured everywhere. However, radon 87 concentration in Radon concentrations may be moderate during diffuse degassing magmas may 88 be low, but during fracture opening it they may reach extremely high values (well over 89 $\frac{10^{6}}{Bq}$ higher than 10^{6} Bq/m³ within the, as measured in Stromboli crater area; Cigolini et al., 90 2013). Its ascent towards the surface is strictly ruled by the mobility of other gas phases, such as 91 CO₂ and H₂O defined as "carrier gases" (Gauthier and Condomines, 1999). In volcanic areas, the 92

93	combined surveys of soil 222 Rn emissions and CO ₂ flux give us the opportunity to track diffuse
94	degassing and related processes (Giammanco et al., 2007). However, it is well known that
95	environmental parameters are critical in modulating gas release from soils (including radon and
96	CO2, The joint measurements of soil CO2 flux and ²²² Rn activity have been used to search
97	possible volcanic and seismic precursors (Makario Londoño, 2009), as well as to track fluid
98	migration and outgassing along active faults, fractures or fumaroles (Baubron et al., 2002; Faber
99	et al., 2003; Zimmer and Erzinger, 2003). Moreover, combined surveys of ²²² Rn, ²²⁰ Rn and CO ₂
100	give us a clue to discriminate distinct gas sources (i.e. rock fracturing, hydrothermal, magmatic)
101	(Giammanco et al., 2007; Siniscalchi et al., 2010) and, to track the evolution of a volcanic unrest
102	phase (Padilla et al., 2013). Pinault and Baubron, 1996; Carapezza and Granieri, 2004; Pérez et
103	al., 2007; Cigolini et al., 2009).
104	AutomaticContinuous and real-timeautomatic measurements substantially increase the potential
105	role of these gases in forecasting eruptionspossibility to identify precursory signals, since the
106	data are easily collected, transferred, and processed and filtered thus minimizing the effects of
107	environmental parameters on soil degassing in near real-time (Brusca et al., 2004; Viveiros et al.,
108	2008; Cigolini et al., 2009; Carapezza et al., 2009). Environmental parameters are critical in
109	modulating gas release from soils, including radon and CO ₂ (Pinault and Baubron, 1996;
110	Carapezza and Granieri, 2004; Pérez et al., 2007; Cigolini et al., 20102009) and their effects
111	must be considered during continuous geochemical monitoring.
112	In this respect, a promising challenge is to establish a fully-automated data processing able to
113	minimize the effects of environmental factors on the acquired data. In this way, data obtained by
114	the geochemical monitoring networks can be easily transferred to the authority responsible of
115	volcano surveillance. The statistical treatment or the spectral analysis of the data are the mostly

116	used methods to recognize and remove the contribution of the atmospheric factors (e.g.
117	Carapezza et al., 2009; Laiolo et al., 2012).; Rinaldi et al., 2012; Silva et al., 2015; Viveiros et
118	al., 2008; 2014). Particularly, the spectral analysis may be positively applied to recognize diurnal
119	to seasonal cycles and to investigate the processes ruling the release of gases from soils (Rinaldi
120	et al., 2012: Martin-Luis et al., 2015).

Radon concentrations can be diluted by major fluxes of CO2 and water vapor (e.g., Giammanco 121 et al., 2007; Siniscalchi et al., 2010). Recently, Girault et al. (2014) and Girault and Perrier 122 123 (2014) have shown, at the Syabru-Bensi hydrothermal system (Central Nepal), that radon is essentially incorporated into the released CO₂generated from a shallow radium sourcessource (a 124 125 rock thickness of 100 m is sufficient to account for the observed radon discharge). An alternative model with a deeper source of CO_2 (carrying) and incorporated into upraising CO_2 . In active 126 volcanoes radon can be carried to the surface from greater depth) is also possible great depths 127 along major faults-(Girault and Perrier, 2014), Cigolini et al. (2013) have shown that extremely 128 129 high radon emissions at Stromboli volcano can be related to the ascent of hot to supercritical CO_2 -bearing hot fluids (up to 270 410 °C) along the fractures (200-300 m deep) that 130 surroundsurrounding the crater rim- of Stromboli volcano (at about 700-720 m a.s.l.) and well 131 correlate with the estimated depth of the source region of VLP events (e.g., (Chouet et al., 1997; 132 Marchetti and Ripepe, 2005). Previous investigations showedhave shown that CO2 fluxes and 133 ²²²Rn concentrations at Stromboli are within the range of those measured in other open-conduit 134 active volcanoes (InguaggiatoCigolini et al., 2013; CigoliniInguaggiato et al., 2013). 135 136 In this paper, we present a four years-long period (from April 2007 to October 2011) of continuous monitoring of ²²²Rn activity and soil CO₂ flux recorded<u>collected</u> by two automatic

stations located on the north-eastern upper flank of Stromboli Island (Fig. 1). These The 138

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139	measurement sites have been chosen in the light of previous surveys. In fact,: anomalous radon
140	values were recorded in this site during periods of sustained volcanic activity and before, during
141	and after the paroxysmal explosion of March 15, 2007 (Cigolini et al., 2013). Similarly,
142	systematic measurements of soil CO_2 flux revealed anomalous degassing zonesareas on the
143	volcano slopes and this site has been identified this site as a potential target for continuous
144	monitoring (Carapezza et al., 2009). Collected data on both gas species were analyzed to define
145	statistical parameters (background, threshold and anomalies; Carapezza et al., 2009; Cigolini et
146	al., 2009; 2013). Here, we applied two different statistical methods (Multiple Linear Regression
147	and Principal Component Regression) for a critical evaluation on the effects of some
148	environmental parameters on the acquired gas species. Both methods allow to minimize the
149	contribution of the atmospheric factors on the CO2 and radon signal and attempt to correlate
150	gases anomalous variations with the volcanic regimes operating at Stromboli during the recorded
151	time span. Finally, we briefly discuss the future application of the statistical treatments on near
152	real-time measurements adopted in monitoring volcanic activity.

153

154 **2 – Stromboli volcano**

Stromboli is the north-easterneasternmost island of the Aeolian archipelago and-it reaches an elevation of 924 m a.s.l. (Fig. 1). It is a composite stratovolcano consisting of lava flows alternated with abundant tephra deposits. The emerged part of the volcanic edifice was built withinin the last 100 ky (Francalanci et al., 1989; Hornig-Kjiarsgaard et al., 1993). The morphology of the island results from periods of extrusive growth alternated to lateral collapses, in turn related to the dyke intrusions, magma upwelling and regional tectonics (Tibaldi, 2003 and 2004; Corazzato et al., 2008). The volcano is well known for its typical persistent explosive related to the typical persistent explosive related to the typical persistent explosive related to the typical persistent explosive related to typical persistent explosive related to typical persistent explosive related to typical persistent explosive

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162	activity called Strombolian, that started approximately 2 ky ago (Rosi et al., 2000; Arrighi et al.,
163	2004). Strombolian activity is characterised by continuous degassing with the emission, on
164	average every 15-20 minutes, of juvenile material (glowing scoriae, lapilli and ash) ejected from
165	the active vents located within the crater terrace at \sim 700 <u>m.</u> a.s.l. This mild explosive activity is
166	episodically interrupted by lava flows, major and paroxysmal explosions (Barberi et al., 1993
167	and 2009) that can be accompanied by flank failure and collapses, which may also generate
168	tsunamis, like in 1930 and recently in December 2002. These events may affect the South-
169	Central Tyrrhenian sea (Tinti et al., 2006). Paroxysmal events, such as the ones occurred on
170	April 5, 2003 and March 15, 2007, are the most violent volcanic explosions of Stromboli and are
171	characterized by the ejection of the so-called "golden pumices" (nearly aphyric, phenocrysts < 10
172	vol%, highly vesicular > 50 vol%, low viscosity K-basaltic pumiceous materials; Métrich et al.,
173	2005 and 2010). These ejecta are generally mixed with degassed scoriaesscorias (the latter also
174	ejected during the typically mild Strombolian activity) and with ballistic solid blocks. The CO ₂
175	and H ₂ O contents measured in primitive melt inclusions, found within forsteritic olivines of the
176	golden pumices, indicate that these materials represent the undegassed magma residing in the
177	deeper part of the Stromboli plumbing system (Bertagnini et al., 2003; Francalanci et al., 2004;
178	Métrich et al., 2005; Cigolini et al., 2008; 2015<u>2014</u>).
179	Soon after the 2002-2003 effusive event, a great improvement of the monitoring system was
180	undertaken under the coordination of the Italian Civil Protection Department. Since then, This

undertaken under the coordination of the Italian Civil Protection Department. Since then, This
advance on ground-based monitoring allowed the scientific community acquired a a
tremendousgreat amount of geophysical, geochemical and geodetic data were collected during
the most recent eruptions 2007 and 2014 (effusive episodes, as well as during the span of time
characterized by low to high explosive activity (cf. Barberi et al., 2009; Ripepe et al., 2009;

185 <u>Calvari et al., 2014;</u> Rizzo et al., 2014). Recent investigations showhave also shown that inwithin
186 the latter years there has beenwas an increase in thermal activity the radiative heat power
187 associated to several minor lava overflows within the summit area (Coppola et al., 2012; Calvari
188 et al., 2014).

189 In recent years, geochemical Geochemical monitoring at Stromboli has involved the following 190 activities: soil radon concentrations concentration (Cigolini et al., 2009; and 2013), soil CO₂ flux 191 (Carapezza et al., 2004 and 2009; Federico et al., 2008; Rizzo et al., 2009 and 2014), fully-192 automated SO₂ plume measures by COSPEC (Burton et al., 2009), continuous 193 measurements of CO_2/SO_2 ratios within the volcanic plume (Aiuppa et al., 2009 and 2011). 194 These methods were tested to eventually forecast paroxysms and major explosions. However; in 195 addition, the continuous monitoring of low-temperature fumaroles has also been found aswas 196 useful-tool to detect short-time changes in volcanic activity (Madonia and Fiordilino, 2013). The 197 extension and structure of the complex hydrothermal system of the volcano has-also been 198 investigated by multidisciplinary studies (involving electrical resistivity, soil CO₂ 199 concentrationsconcentration, temperature and self-potential measurements, ef.; Finizola et al., 200 2006 and 2009; Revil et al., 2011). Nevertheless systematic data collectionGeochemical studies 201 on the geothermal aquifer (e.g. chemical physical parameters, dissolved gas, isotopes) has been 202 very usefulat the periphery of the volcano is an additional tool to detect precursory signals of an impending eruption (Carapezza and Federico 2000et al., 2004; Capasso et al., 2005). 203

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205 3 – Methods and Techniques

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206	Preliminary radon and CO2carbon dioxide surveys were conducted to find the most appropriate
207	sites for continuous monitoring. A network of 21 radon stations has been operative at Stromboli
208	since 2002 (e.g. Cigolini et al., 2005; 2009; 2013). Systematic measurements were undertaken by
209	using LR115 track-etches alpha-detectors exposed from two to six weeks (Bonetti et al., 1991),
210	in order to obtain continuous time series on ²²² Rn emissions. Additionally, periodic short-term
211	measurements has been performed by means of EPERM® electretes (Kotrappa et al., 1993) that
212	allowed us to better correlate radon emissions with the variations of volcanic activity (Cigolini et
213	al., 2005 and 2007). These periodic measurements demonstrated that diffuse degassing occurs at
214	Stromboli mainly along the main structural discontinuities-(Fig. 2). After the February-April
215	2007 effusive-explosive event, a real-time station for radon measurements was first installed at
216	520 m a.s.l. at Liscione, on the northeastern side of the cone (see Fig. 1 and Fig. 2a). Similarly, a
217	soil CO ₂ flux survey first outlined the main sectors of gas emanation (Carapezza and Federico,
218	2000) and two automatic soil CO_2 flux stations (and environmental parameters) were installed at
219	Stromboli: one at the summit (Pizzo sopra La Fossa) in 1999, and the second one near the sea-
220	shore in 2001 (Pizzillo). In the following years, CO ₂ soil concentration surveys, within the crater
221	terrace and surrounding areas, were performed to identify the sectors of major degassing and
222	higher hydrothermal activity (Finizola et al., 2002 and 2003). Furthermore, Carapezza et al.
223	(2009) performed a wide detailed survey of soil CO ₂ flux on the island and sectors of anomalous
224	degassing were detected. Therefore, two-additional soil CO2 flux and multiparametric fully-
225	automated stations were installed along the ENE flank of the volcano (respectively at Rina
226	Grande and Nel Cannestrà) where gaseous anomalous gas emissions are higher were found
227	(Carapezza et al., 2009) (Fig. 2b). The area is confined between two major structural

228	discontinuities (the N40E fault, and the N60E fault) that cross-cut the northeastern sector of
229	Stromboli.
230	The gasDataset analysed in this paper refer to the ²²² Rn and CO ₂ measurements acquired by fully
231	automated stations located in the Nel Cannestrà sector (see Fig. 1 and 2). The area is confined
232	between two major structural discontinuities (the N40°E fault, and the N60°E fault) that cross-
233	cut the north-eastern sector of Stromboli. These automated stations consist of two units, one for
234	measuring the isotope of the radon progeny (together with soil temperature and atmospheric
235	pressure) and the other for measuring soil CO_2 flux (by accumulation chamber) together with
236	environmental parameters (atmospheric temperature, humidity and pressure; soil humidity and
237	temperature; wind direction and horizontal speed). The radon unit provides near real-time
238	measurements of ²²² Rn concentrations (by using a DOSEMan, Sarad Gmbh, Germany)
239	connected to an electronic board able to acquire and transfer the collected data to a radio-modem
240	that sends, by means of a directional antenna, the signal to the COA volcano observatory
241	(Cigolini et al., 2009). Specifically, data are acquired. The station acquires data every 30 minutes
242	and the radon concentration and soil temperature are measured at 1 m depth-(Cigolini. The
243	DOSEMan radonmeter measures α-particles within a 4.5-10 MeV energy window, including
244	both ²¹⁸ Po and ²¹⁴ Po peaks (Gründel and Postendörfer, 2003). An exhaustive description of the
245	radon dosimeter and of the real-time ²²² Rn station can be found in Cigolini et al. (2009) and
246	<u>Laiolo</u> et al., <u>2009. (2012</u>). Soil CO ₂ flux and environmental parameters are measured hourly
247	with a fully equipped automated station produced by West Systems (see Carapezza et al., 2009
248	for method description). Soil temperature and humidity are measured at 50 cm depth; air CO_2
249	concentration is measured 30 cm above the soil/air interface (Carapezza et al., 2009). Data are
250	stored on a non-volatile memory and can be retrieved by means of a telemetry system at the

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252	both stations are reported in the supplementary materials (Table S1).	
253	The timespan investigated in this work (April 2007 - September 2011) matches the period in	
254	which both instrumentations were mostly operativesoperative. In fact, the ²²² Rn station was	
255	installed in early April 2007 and is still operative whereas the automated CO ₂ flux station,	
256	installed in mid March 2007, was dismissed in late September 2011.	
257		
258	4 – Continuous monitoring <u>Results</u>	
259	<u>4.1 – Time series ρf radon activity and soil CO₂ flux from the soil and ²²²Rn activity.</u>	
260	-The time series for ²²² Rn activity and soil CO ₂ -flux (April 2007 September 2011) together with	
261	those of the environmental parameters, are reported in Figures 3 and 4, respectively. The main	
262	descriptive statistics are reported in Table 1.	
263	The overall behaviour of the CO ₂ and ²²² Rn signals is somehow similar in the first two years:	
264	they both show bell-shaped profiles, strongly ruled by seasonseasonal trend, that reach their	
265	lower and stable values during summer and the higher values and, with a wider variation range,	
266	in winter. A similar behaviour is shownobserved also in-terms of multi-modal distributions (see	
267	histograms in Fig. 3). Both trends display several marked spikes within each time series (Fig. 3a	
268	and 3b), with) and numerous peaks of boththe two gases beingare essentially concordant, thus	
269	indicating. In the last two years, soil CO ₂ flux shows a common origin. decreasing trend,	
270	whereas radon activity maintains nearly the same annual average, or simply increase (see Table	

<u>1).</u>

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251 volcano observatory (COA, see Fig. 1). The main technical characteristics of the sensors used in

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272	A preliminary analysis of the trend of environmental parameters, supported by the correlation
273	factors reported in Table 2, Compared with other active volcanic areas, radon shows that both the
274	²²² Rn and CO ₂ signals are inversely correlated to soil and air temperature, and positively
275	correlated to soil humidity (especially radon). The decoupling of soil temperature from the radon
276	signals is well known: convective cells that rule gaseous transfer toward the surface do not reach
277	the surface due to the inversion of the thermal gradient within surface soils due to summer
278	heating (Mogro Campero and Feischer, 1977; Cigolini et al., 2001; 2009; Laiolo et al., 2012). It
279	is interesting to note that this phenomenon is affecting CO_2 fluxes as well.
280	The wind speed signal is rather variable and fluctuating. However, it seems to be more stable
281	during late spring summer, when winds are substantially below 7 m/sec (Fig. 4b).
282	Radon concentrations (see Fig. 3a) show-relatively low values (concentration (cf. Cigolini et al.,
283	2009),2013 and references therein). Values are essentially below 2000 Bq/m ³ for a large part of
284	the year, and may exhibit a short-term variability. In fact,; the average radon-activity in the four
285	<u>years</u> is around 2000 Bq/m ³ with a standard deviation of 1200 Bq/m ³ (Table 1). During winter
286	(November-February), radon typically exhibits higher average values (Fig. 3a) with peaks up to
287	7900 Bq/m ³ . We remind that the 222 Rn activity in the summit area, close to the active vents,
288	shows significantly higher average values reaching 12,500 Bq/m ³ (\pm 4,200; see-Laiolo et al.,
289	2012; Cigolini et al., 2013).
290	Radon seasonal minima refer to late spring-summer periods (March-October) with average
291	values of ~1200 Bq/m ³ and hourly minima close to 200 Bq/m ³ . This behaviour is <u>The ²²²Rn long-</u>
292	term stability on low values seems closely related to the absence of marked weather changes

293 during this season.

294	Radon emissions The time-series of environmental parameters are strongly negatively shown in
295	Fig. 4; a preliminary analysis of their effects, supported by the correlation coefficients (Table 2),
296	shows that both the ²²² Rn and CO ₂ signals are inversely correlated to soil and air temperature, as
297	already pointed out by Cigolini et al. (2009) and Laiolo et al. (2012). In turn, soil humidity
298	(higher after rainfalls) is somehowand positively correlated with(especially radon emissions
299	(Table 2).) to soil humidity, in turn depending on rainfall. It is interesting to point out-that
300	recorded data show that sudden variations in radon concentrations normally occur within few
301	hours of continuous raining and/or temperature drops. Notably, aA similar phenomenon has been
302	already-observed at Furnas volcano (Azores archipelago) during continuous monitoring of CO2
303	flux measurements (Viveiros et al., 2008) as well as and radon measurements activity (Silva et al.,
304	2015).

The relation between temperature and ²²²Rn activity is ascribed to the local thermal gradient 305 (between soil and air temperatures) that affects the efficiency of the in-soil convective cells and, 306 307 consequently, the migration of gasesgas toward the surface (Mogro-Campero and Fleisher, 1977; 308 Cigolini et al., 2001). Therefore, a marked difference between soil and air temperatures, typical of the fall-winter season, causes an increase in the measured radon activities. Conversely, the The 309 310 entire dataset shows a clear positive correlation (R= 0.74; Table 2) between soil moisture and radon emissions. Indeed; indeed, an increase in soil moisture is capable of raisingmay increase 311 the ²²²Rn emanation coefficient of(i.e. exhalation rate) by one order of magnitude (Nazaroff, 312 1992; Sakoda et al., 2010; Girault and Perrier, 2012). 313

However, to evaluate the relation between ²²²Rn activity and soil gas concentration and humidity in the specific site of radon station, we have also to consider the confinement of the box containing the radon detector. The device is placed in an impermeable polycarbonate case

317	(permeable to ²²² Rn but not to water) at a depth of about <u>80 cm1 m</u> . Thus, if the soil matrix
318	surrounding the case is affected by water saturation, the preferential pathway for radon migration
319	will follow the interface soil-bottom of the case, leading to an increase in α decay counts. An
320	alternative explanation may be Another possibility is that, during a rainfall episode, only the
321	higher portions of soil (down to about 10-30 cm) undergo water saturation that, in turn,
322	temporarily inhibits the free motion of the radon particles toward the surface. Consequently,
323	radon will preferentially be confined at lower levels (i.e., where water content is low or absent)
324	so that decay counts will be drastically higher in the portion of soil where the case (containing
325	the detector) is inserted. Notably, as As the relation between soil humidity and radon activity
326	essentially depends from soil permeability, the observed behaviour can be inhomogeneous over a
327	given sector of the volcano (Perrier et al., 2009).
328	Soil CO ₂ flux measurements were-acquired by the automatic station startingstarted on mid
329	March 2007. Average values The four years average value for CO ₂ fluxes are flux is $\sim 600 (\pm 643)$
330	g m ⁻² day ⁻¹ with 1 with a rather high standard deviation (as for radon) (Table 1). The maximum
331	values were reached in January 2008 with fluxes up to aboutslightly exceeding 7000 g m ⁻² day ⁻¹ .
332	Also in this case As already observed for ²²² Rn activity, minima in soil CO ₂ fluxes occur during
333	summer-early fall when values approaching to zero were recorded (or well below 100 g m ⁻² day ⁻¹
334	were recorded. A similar trend outlines that the variation of the local thermal gradient is capable
335	to affect also the CO ₂ flux from the soil (Viveiros et al., 2014).
336	Surprisingly, there is a noticeable correlation between soil CO ₂ flux and wind, both speed
337	(positive) and direction (negative) (Table 2). In Fig. 5 it is clear how winds blowing toward SE at
338	speed > 8 m/s are able to produce an efficient gas escape from the soil, causing an increase of the
339	CO ₂ flux. Such a behaviour is mainly related to a Venturi effect due to a local condition of the
	15

340 <u>Nel Cannestrà station site, that cannot be extrapolated to other sectors of the volcano,</u>
341 <u>considering that it was not observed in the Rina Grande station (see Fig. 2b for location)</u>

342 (Carapezza et al., 2009).

343 It is interesting to note that the four years-long dataset exhibits a declining long-term trend (Fig. 344 3b and annual average in Table 1) which can be likely viewed as a decreasing supply of CO_2 -rich 345 magma from the deeper to the upper plumbing system. This hypothesis is supported by the long-346 term trend observed in the CO_2 emissions from the plume, retrieved by combining CO_2/SO_2 347 ratio ratio and SO₂ flux measurements (see Aiuppa et al., 2011). The decreasing trend of the soil CO_2 flux, marked by the annual average shifting from 920 (in 2007-2008) to $\frac{310330}{310330}$ g m⁻² day⁻¹, 348 does_(in 2010-2011), is not appear on evident in the radon long-term signals that, instead, show 349 an excellent stability around the average of 2000a slight increase from 1777 to 2264 Bq/m³ in 350 annual averages (see straight linelines in Fig. 3a and 3b, respectively). It is important to outline 351 that, according to this view, the concurrent measurements of CO₂ and ²²²Rn can be considered as 352 353 an important tool for detecting the variation in time of the deep vs. shallow contribution to the soil gas, as it was found in other volcanoes (i.e. Etna; Giammanco et al., 2007 Table 1). 354 However, during a short term period, it is clear that relative increases in CO2 flux coexist with 355 356 and/or are immediately followed by similar trends in ²²²Rn activity. Generally, CO₂ flux shows a more stable signal during short-term periods, indicating that daily temperature variations have a 357 feeble influence on the gas flux. Conversely, as pointed out by Carapezza at al. (2009), there is a 358 359 definite correlation between winds (including both speed and direction) and soil CO₂ fluxes 360 (Table 2). Winds toward the SE blowing at speed higher than 8 m/s cause a more efficient escape 361 of the gas from soil. This is related to a Venturi effect due to a local condition that cannot be extrapolated to other sectors of the volcano (Carapezza et al., 2009). 362

363 This behaviour is evident by calculating the

364 <u>4.2 – Short-term periodicity and long-term trends</u>

In order to identify diurnal and semidiurnal cycles affecting the gas signals, we performed a 365 spectral analysis (Power Spectral Density) over a one year subset of data (sample time = 1 hour), 366 using the method suggested by Viveiros et al. (2014). The analysis identified the 12h and the 24h 367 frequency peaks in both CO_2 flux and ²²²Rn activity (Fig. 6), confirming previous findings 368 (Perrier et al., 2009 and 2012; Rinaldi et al., 2012). In our case, the ²²²Rn signal seems to be 369 370 modulated by temperature and barometric changes, although we do not exclude that this 371 periodicity could be related with solar tides (e.g., Steinitz et al. 2011). On the other hand, the soil 372 CO₂ flux reveals a main 12h period. It is worth noting that such a behaviour slightly differs from previous results suggesting a major influence of temperature rather than pressure (Rinaldi et al., 373 374 2012).

By analysing the long-time series of soil CO_2 flux and ²²²Rn activity, we performed a calculation 375 376 of the mean value of the whole data onfor each specific day of the year. The and the same 377 computation was carried out also on soil and air temperature data. The emerging annual trend from Jan-1 to Dec-31 ((see Fig. 5) remarks7a, b) highlights the inverse relation between 378 379 temperature and soil gas release, as well as thean apparent correlation between the two gas 380 species. Overall, we observe a 100% increment of the mean values comparing the spring-381 summer with the fall-winter period. Is also evident that the most significant day-by-day 382 variations occur during the fall and spring season, when the likelihood of drastic atmospheric 383 changes (i.e. heavy rainfall or windstorm) is higher than during summer-or winter.

The long-period behaviour is ruled by soil and air temperature (i.e. thermal gradient), whereas short-time/high frequency oscillations are modulated by soil humidity (e.g., rainy events) and

wind conditions (speed and direction), respectively. The value of the correlation coefficients (R) shown in Table 2, allowsallow to better evaluateassess the effects effect of the environmental parameters that actively modulate the trends of 222 Rn activity and soil CO₂ flux-trends.

389 In Fig. 6 we report

390 4.3 – Variation of the correlation coefficients

391 The seasonal time variation variations of average correlation coefficients of somethe main environmental parameters and compare them-with ²²²Rn activity and soil CO₂ flux. Here are 392 393 reported in Fig. 8, where seasons are gathered and simply subdivided in spring-summer and fall-winter subsets. It can be seen that ²²²Rn shows positive correlation with soil water content 394 and negative correlation with air and soil temperatures. Similarly, soil CO₂ flux is mostly 395 correlated positively to wind speed and negatively to air and soil temperatures. However, as 396 397 shown in Fig. 6, the correlation factors are the correlation coefficients are not so stable throughout the investigated period time span, but appear slightly modulated by seasonal effects. 398 399 For example, both <u>soil CO_2 flux and radon activity display a more distinctive negative</u> 400 correlation with air and soil temperatures during the spring-summer subsets. In addition, seasonal variations in geochemical data and environmental correlation coefficients are This 401 402 behaviour is likely due to the lack of drastic variations in weather conditions during the "dry" 403 season at Stromboli Island. Hence, the correlation between temperature and gas flux and concentration is not perturbed by other atmospheric factors (e.g. soil humidity). Correlation 404 405 coefficients seem somewhat different in the first subset (spring-summer 2007) compared to all 406 the others. This; this subset (grey field in Fig. 8) was sampled obtained just after the March 15, 407 2007 explosive paroxysm when the effects of environmental conditions on degassing dynamics, 408 even in relatively distal areas, wereseem somehow weaker. In fact, the April-June 2007 period

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409	represents a time span characterized by the interruption of the lava effusion ceased and the	
410	Strombolian activity and by the ceasing of lava effusion. was not resumed, or more precisely, the	
411	source of explosions was too deep to allow glowing scoriae to reach the crater surface. In this	
412	time span-, there was still a remarkable degassing rate at the erater vents likely related to a	
413	continuous supply of undegassedcraters from a relatively deep-seated magma to the upper	
414	feeding system of the volcano (Burton et al., 2009; Carapezza et al., 2009; level (Aiuppa et al.,	
415	2011; Cigolini2009; Barberi et al., 20132009).	
416		
417	54.4 Statistical Trootmonttreastment	Formatt After: 6
417	5 <u>4.4</u> – Statistical Treatment<u>treatment</u>	Formatt
418	Two different statistical methods have been applied toon the entire datasets raw dataset (sample	Formatt
419	<u>time = 1 hour</u>) of soil CO ₂ flux and ²²² Rn concentration in order to identify and remove gas	
420	variationsthe effects due to environmental parameters.	
421	54.4.1 - Multiple Linear Regression Statistics (MLR)	Formatt After: 3
422	The data sets datasets of radon concentration, soil CO_2 flux and environmental parameters have	
423	been analyzed analysed by the Multiple Linear Regression (MLR) which is a simple and largely	
424	applied method used to identify the contributions of several independent variables and model the	
425	fluctuations observed in the investigated signal. The analysis has been performed to predict the	
426	values of a dependent variable (Y) given a set of predictor variables ($X_1, X_2,, X_n$). The	
427	relationship between the dependent variable (Ycalc) and the predicted variables is expressed as	
428	$\underline{Y_{calc}} = Y_0 + b_1 X_1 + b_2 X_2 + \dots + b_n X_n $ (1)	
429	Y_0 is the intercept, X_n are the acquired variables and b_n the calculated regression coefficients	

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431	simplify and restrainreduce the number of predictor variables, MLR takes into account only the	
432	factors that are more correlated (positively or negatively) with CO ₂ <u>flux</u> and ²²² Rn <u>activity</u> (Table	
433	2).3). By considering previous research, we selected only the environmental factors (i.e.	
434	independent variables) causing an increment of the R ² greater or equal to 1% (Viveiros et al.,	
435	2008; Silva et al., 2015). Particularly, soil and air temperature were considered indicated by the	
436	regression for both gas species, together with wind $\frac{\text{conditionsspeed and direction}}{\text{conditionsspeed and direction}}$ for CO ₂ and	
437	soil humidity for ²²² Rn. This statistical approach can be quickly performed, thus it can be easily	
438	applied on near real time measurements. In Table 3 we report the main parameters for both gases	
439	by the MLR analysis.	
440	Results show that the atmospheric variables taken into account for this analysis are able to	
441	predict 45% and 51% (R=_0.67 and R=_0.71) of the variations observed in soil CO ₂ flux and	
442	222 Rn concentration, respectively. Moreover, soil CO ₂ flux and 222 Rn values show low	
443	dispersions, <u>as</u> respectively <u>the 4.08</u> % and 4.80% of the computed residuals exceed the average	
444	± 2 standard deviation (Table 3). <u>range.</u> Predicted values by MLR both for radon concentration	
445	and soil CO ₂ flux, together with the observed values and the calculated residuals, are plotted in	
446	Fig. $\frac{7a9a}{2}$ onto the recorded time series. By looking at the residuals, it can be seen that <i>i</i>) the bell	
447	shape of the CO_2 flux is somehow smoothed due to the removal of the seasonal trend and but it	
448	disappears totally in the last two years; whereas the bell shape still persists for radon, and <i>ii</i>) the	
449	residuals of both gases show peaks and major fluctuations that obviously cannot be duerelated to	
450	environmental variations. Moreover, in both cases the signals are still characterized characterised	
451	by noisy components (similar conditions were also previouslyas reported by Viveiros et al.,	
452	2008; Carapezza et al., 2009; Laiolo et al., 2012 and Viveiros et al., 2014).	

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453 Notably, this<u>This</u> statistical approach has been used at different volcanoes in the attempt to
454 detect short-term variations in volcanic activity (e.g. major explosions at Stromboli; Laiolo et al.,
455 2012), as well as the effects of seismic sequences (due to stress/strain structural changes) on the
456 shallow part of a volcanic edifice (e.g., Masaya as analyzedanalysed by Padilla et al., 2014).

457

54.4.2 - Principal Component Regression (PCR)

The second statistical treatment applied $\frac{10^{222}}{10^{222}}$ Rnto 222 Rn activity, soil CO₂ flux and environmental 458 parameters is the Principal Component Regression (PCR). This method differs from the previous 459 one in how predictors are treated: first, a factor analysis is performed on the environmental 460 dataset (X); then a forward step-wise linear regression of measured soil CO_2 flux and radon 461 activity (Y) is performed on the estimated factors. The goal of this approach is firstly to obtain a 462 463 reduction in the X data set in a way that maintains the maximum amount of information (i.e. 464 largest possible variance), and secondly to perform regression of Y on orthogonal (uncorrelated) 465 components. The aim is to ensure that highly correlated principal components are not overlooked 466 (Vandeginste et al., 1998).

We had to perform performed factor analysis on two separate datasets because the radon station 467 measures soil temperature and air pressure, whereas soil CO₂ flux station also measures air and 468 soil humidity, wind speed and direction. SoTherefore, we created one dataset for radon with soil 469 470 temperature and atmospheric pressure of measured at the radon station and added the other variables from measured at the CO₂ station. The factor analysis of the two datasets showshows 471 472 that three eigenvalues are higher than 1.0 and these three factors can explain the 73% of the total variance- (Table 4). In the second step, we performed forward step-wise regressions of ²²²Rn 473 concentration and soil CO_2 flux on the first three factors. We obtained two theoretical models 474

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475	which explain the 25% and the 47% (R=_0.50 and R=_0.68) respectively of the soil CO ₂ flux and	
476	²²² Rn concentration measurements variance .	
477	However, as (Table 4b). As in the MLR model, residual values show low dispersion, being less	
478	than 5% the portion of the data that exceeds the mean $\pm 2\sigma$ range (4.11% and 4.89% for CO ₂ and	
479	²²² Rn, respectively) of the data that exceed the mean standard deviations ($\pm 2\sigma$; cf. Table 3).). The	
480	timeseriestime series of observed, predicted and residual values of radon concentration and soil	
481	CO_2 flux are reported in Fig. 7b9b. An overall comparison of the latter with Fig. 7a9a shows that	
482	the two statistical treatments provide basically the same results.	
483		
484	65 – DISCUSSION	 Formatted: Space Before: 12 pt, After: 12 pt
485	6.1 Considerations on For ²²² Rn, the applied statistical treatment	Formatted: Font: Not Italic
486	On the whole, the results obtained by applying the two-methods are apparently very similar (Fig.	
486 487	On the whole, the results obtained by applying the two-methods are apparently very similar (Fig. 7) but they provide some significant differences. In fact, the indicate nearly the same percentage	
487	7) but they provide some significant differences. In fact, the indicate nearly the same percentage	
487 488	7) but they provide some significant differences. In fact, the indicate nearly the same percentage of the data that can be attributed to environmental variations is quite similar for ²²² Rn, , whereas	
487 488 489	7) but they provide some significant differences. In fact, the indicate nearly the same percentage of the data that can be attributed to environmental variations is quite similar for 222 Rn, , whereas the predicted soil CO ₂ flux values differvary from 45% in MLR to 25% in MLR and PCR	
487 488 489 490	7) but they provide some significant differences. In fact, the indicate nearly the same percentage of the data that can be attributed to environmental variations is quite similar for 222 Rn, whereas the predicted soil CO ₂ flux values differvary from 45% in MLR to 25% in MLR and PCR methods, respectively (see Table 2). Variations in The residual computed values are related to	
487 488 489 490 491	7) but they provide some significant differences. In fact, the indicate nearly the same percentage of the data that can be attributed to environmental variations is quite similar for 222 Rn, , whereas the predicted soil CO ₂ flux values differvary from 45% in MLR to 25% in MLR and PCR methods, respectively (see Table 2). Variations in. The residual computed values are related to processes that bypass meteorological factors and are likely related to the volcanic system and	
487 488 489 490 491 492	7) but they provide some significant differences. In fact, the indicate nearly the same percentage of the data that can be attributed to environmental variations is quite similar for 222 Rn, , whereas the predicted soil CO ₂ flux values differvary from 45% in MLR to 25% in MLR and PCR methods, respectively (see Table 2). Variations in. The residual computed values are related to processes that bypass meteorological factors and are likely related to the volcanic system and occur either within the shallow hydrothermal aquifer or in the deep magmatic plumbing system.	
487 488 489 490 491 492 493	7) but they provide some significant differences. In fact, the indicate nearly the same percentage of the data that can be attributed to environmental variations is quite similar for ²²² Rn, , whereas the predicted soil CO ₂ flux values differvary from 45% in MLR to 25% in MLR and PCR methods, respectively (see Table 2). Variations in The residual computed values are related to processes that bypass meteorological factors and are likely related to the volcanic system and occur either within the shallow hydrothermal aquifer or in the deep magmatic plumbing system. However, it It can be noted that the radon treatment provides treatments provide many significant	
487 488 499 490 491 492 493 494	7) but they provide some significant differences. In fact, the indicate nearly the same percentage of the data that can be attributed to environmental variations is quite similar for ²²² Rn, , whereas the predicted soil CO ₂ flux values differvary from 45% in MLR to 25% in MLR and PCR methods, respectively (see Table 2). Variations in. The residual computed values are related to processes that bypass meteorological factors and are likely related to the volcanic system and occur either within the shallow hydrothermal aquifer or in the deep magmatic plumbing system. However, it <u>It</u> can be noted that the radon treatment provides treatments provide many significant negative residual values $\geq -2\sigma$, whereas only one negative residual is recorded for soil CO ₂	

497	the effects of the environmental factors. We explain the high fluctuations showed by the
498	residuals of radon signal with the relative low sensitivity of the radon dosimeter (Table S1) when
499	settled with a high sampling rate (1 hour) in areas characterised by general low emissions (<
500	2000 Bq/m ³). In fact, such a noisy signal was not observed in the datasets acquired where the
501	radon emissions are higher (Laiolo et al., 2012). Moreover, a trend characterised by positive and
502	negative fluctuations in the calculated ²²² Rn residual values, has been already observed in a non-
503	volcanic area (Hayashi et al., 2015).
504	The residual time series (Fig. $\frac{7a}{7b}$) retrieved for $\frac{CO_2}{CO_2}$ -soil $\frac{CO_2}{CO_2}$ flux shows in both
505	treatments a decreasing trend for the first two years (visible also in the raw data) followed by a
506	nearly steady-state signal close or below the zero value. This behaviour supports the hypothesis
507	that the supply of $\underline{CO_2}$ -rich magma from the deep plumbing system (that started before the 2007
508	eruption, cf., Aiuppa et al., 2011) likely increased), besides increasing CO ₂ emissions not only
509	within emission in the gas plume itself, but induced also created a higher CO_2 flux in the more
510	distal zones, which lasted for nearly two years.
511	Comparison of the standard residuals (both by PCR and MLR) for CO ₂ soil flux and ²²² Rn
512	activity versus time (Fig. <u>810</u>) shows that in the initial two years of monitoring (up to May 2009)
513	the two gases display a similar behaviour with nearly synchronous alternation of periods with
514	anomalous emissions and periods with few or no anomalies (such as the summer of 2007 and
515	2008). In the last two years considered, only radon shows frequent positive anomalies whereas

516 anomalous CO_2 flux values are <u>only</u> rarely recorded.

Finally, the <u>The</u> comparison between the two statistical treatments suggests that the MLR seems
to be the <u>is</u> more adequate method for getting the quick results needed in near-real time volcano
monitoring, whereas PCR ensures a more accurate <u>estimation estimate</u> of <u>warning thresholdsdata</u>,

520 which might be eventually useful for a more accurate post-process analyses and for a more reliable monitoring. In this view, because of the more reliability of the PCR treatment, in 521 the next section we briefly compare our emerging results with the 2007 2011 Stromboli volcanic 522 activity. 523 524 6.2 Diffuse degassing vs. volcanic activity 525 526 In order to testassess the reliability of radon activity and CO₂soil CO₂ flux, measured in the distal 527 site of Liscione/Nel Cannestrà, as possible geochemical-precursors of major changes in the volcanic activity, the residuals time-series of both species obtained by PCR methodand MLR and 528 exceeding 2σ , have been compared with the main volcanic and seismic events occurred at 529 530 Stromboli volcano duringin the considered timespansame time span (Fig. 8). During the 10). The 4.5 years of gas monitoring (April 2007 - September 2011) represent a phase of ordinary 531 volcanic activity of Stromboli. In this period, twelve major explosions, four minor lava 532 overflows and a local earthquake (with M_L=_2.2) were recorded at Stromboli by the INGV 533 monitoring system (ef.-Calvari et al., 2014). HNo explosive paroxysm or effusive eruption 534 occurred. There is no clear correlation between our data and the recorded volcanic events, but 535 536 some useful considerations can be done. 537 From Fig. 10, it can be seen from Fig. 8 that in the first two years (following the 2007 explosive paroxysm), periods of and effusive eruptive phase), frequent and high CO₂ flux and ²²²Rn 538 539 anomalies are essentially concordantwere recorded in coincidence with periods of some 540 anomalous volcanic activityepisodes. Actually, the high number of positive residuals, from October 2007 to June 2008 and from November 2008 to May 2009, coincide with five major 541 explosions and one lava overflow. InDuring the summer of 2007, 2008 and 2009, though no 542

543	radon data are summers, neither major explosion/lava overflow nor residual CO2 flux peaks were
544	recorded (very few for radon, apart from 2009 summer when data were not available, no
545	anomalous gas emissions and no major explosions were recorded. However in the-). In summer
546	of 2010, two major explosions occurred in a period of no anomalous gas release. Finally, during
547	September-During December 2010 August - February 2011, a major explosion and three two
548	lava overflows occurred during in concurrence with isolated peaks of CO ₂ flux and during a
549	phase of anomalous ²²² Rn emission. Also the most-recent lava overflow in September 2011
550	coincided with an anomalous emission, with only few isolated peaks of soil CO2-fluxradon
551	activity value.
552	7Conversely, we have to consider that the eruptive events that occurred in the above time span
553	seem to be connected to minor changes associated to the dynamics of the upper part of the
554	conduit (e.g., Barberi et al., 1993 and 2009) without any involvement of the deep seated gas-rich
555	magma pockets (typically occurring during major effusive-explosive cycles of Stromboli
556	volcano, such as those of 2002-2003 and 2007).
557	
558	<u>6</u> - CONCLUSIONS
559	The presented data show that therefer to more than four years of soil gas measurements (²²² Rn
560	concentration and soil CO ₂ flux) at relatively distal sites from the active vents (Liscione and Nel
561	Cannestrà sites, located on the NE flank of Stromboli) are slightly influenced by volcanic
562	activity that, in, Fig. 1) during a the time-span considered, did not produce any relevant
563	explosions or (April 2007-September 2011) without major lava effusion. effusions and
564	paroxysmal explosions.

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565	The long-time averages for CO_2 flux and radon concentration exhibit relatively low values (585
566	g m ⁻² day ⁻¹ and 2050 Bq/m ³ , respectively) when compared to those measured at the summit crater
567	area (Carapezza et al., 2009; Cigolini et al., 2009 and 2013). This means that the advective
568	processes, able to enhance the gas release from soil, are considerably reduced moving away from
569	the crater area. However, the The long term declining trend observed for the soil CO_2 flux (Fig. 9)
570	and Table 1) suggests that the large supply of CO ₂ -rich magma associated with the 2007 eruption
571	(and invoked to explain the exceptional CO ₂ emissions from the plume <u>; Aiuppa et al., 2009 and</u>
572	2011) affected also the soil gas release in relatively distal areas. So, as already stressed by De
573	Gregorio et al. (2014) for Etna volcano, the soil CO ₂ flux measurements represent a key tool to
574	infer the magma supply dynamics and to evaluate the local degassing regime. Furthermore, the
575	combination of soil CO ₂ flux and ²²² Rn concentration measurements can better constrain the gas
576	source in relation to changes in volcanic activity (Faber et al., 2003; Perez et al., 2007; Padilla et
577	al., 2013)., e.g., Aiuppa et al., 2011) affected the soil gas release in distal areas as wellIn the last
578	two years (2010 - 2011) of our monitoring, anomalous radon concentrations have been
579	frequently recorded in periods with rare or absent soil CO ₂ flux anomalies; this likely indicates a
580	different source for the two gases, deeper for CO ₂ and somehow shallower for radon.
581	As recently stressed by De Gregorio et al. (2014), measuring the soil CO2 flux at active
582	volcanoes is a valued tool to infer the magma supply dynamics. Thus, combining CO2 flux and
583	²²² Rn concentration measurements gave us the opportunity to better investigate the changes in
584	volcanic activity associated with magma rise.
585	The four years monitoring of both gas species at Stromboli provided the opportunity of better
586	decoding how gaseous transfer toward the surface is ruled by environmental changes. Our data
587	show that both gases are affected by seasonal temperature variations that givegiving to the time

series a trend with a bell shaped profile. In particular, higher emissions occur during fall-winter, because fluid convection is promoted by the higher soil-air temperature gradient. Conversely, during summer, this gradient is reversed and near-surface convection is inhibited. Moreover, soil CO_2 flux is <u>locally</u> influenced <u>also</u> by wind <u>speed</u>, <u>and(>8 m/s in the SE sector)</u>, <u>while</u> radon activity by soil humidity.

In summary, the combined effect of decreasing surface temperaturestemperature, eventually coupled with increases in soil moisture, seems the main factor that controls the variations of radon emissions. The effect of soil humidity on radon activity probably reflects the adopted measurement techniques. In fact the radon measurements at 1 m depth are likely affected by soil humidity (particularly during the raining falls) which affects the radon diffusion and exhalation rates (i.e., Papachristodoulou et al., 2007).

We have shown that the influence of environmental parameters on gaseous timeseries time series 599 600 can be removed minimised by means of linear statistics and give us the opportunity to better track 601 the evaluate possible variations related to changes in volcanic activity. The statistical 602 analyses methods presented in this paper can be adopted for different purposes. Particularly; the 603 Principal Component Regression (i.e. Factor Analysis) appears to be the better suitedmore 604 suitable for an accurate analysis of large datasets following major changes in volcanic activity 605 (post-event data processing): in fact, the application of this method carefully evaluates the 606 contribution of each independent factor by means of precise cross correlations. Conversely, 607 Multiple Linear Regression analysis can be more quickly and easily applied duringto a nearly 608 real-time soil gas monitoring. This useful in volcano surveillance since it gives us the 609 possibilityopportunity to efficiently track anomalies inanomalous gas concentrations, or fluxes, that are not necessarily related to environmental factors. We thus emphasize that the reported 610

611	datasets represent a rather unique case, at the global scale as well, of geochemical and	
612	environmental data acquired in a very active volcanic area for such a long time. The monitored	
613	area represents an anomalous degassing zone (Carapezza et al., 2009; Cigolini et al., 2013) and	
614	our results show that a multiparametric geochemical monitoring may play a key role in decoding	
615	precursory signals related to major changes of Stromboli volcanic activity.	
616	We finally emphasize that the monitored area has been characterized by marked anomalous	
617	radon emissions prior and during the paroxysmal explosion of March 15, 2007 (Cigolini-et al.,	
618	2013)Therefore, it is not excluded that multiparametric geochemical monitoring may play, in	
619	the future, a key role in decoding precursory signals related to major explosive events that affect	
620	the NE sector of Stromboli.	
621		
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625 626 627	"Presidenza del Consiglio dei Ministri – Dipartimento della Protezione Civile (DPC)" through the DEVnet Project (a cooperative program between the Departments of Earth Sciences of the University of Torino and the University of Florence)) and through the "Potenziamento Monitoraggio Stromboli"	
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636	measurement of the volcanic gas plume CO2/SO2 ratio. J. Volcanol. Geotherm. Res. 182
637	(3-4), 221–230.
638	Aiuppa, A., Burton, M., Allard, P., Caltabiano, T., Giudice, G., Gurrieri, S., Liuzzo, M., Salerno,
639	G., 2011. First observational evidence for the CO2-driven origin of Stromboli's major
640	explosions. Solid Earth 2 (2), 135-142.
641	Allard, P., Carbonelle, J., Dajlevic, D., Le Bronec, J., Morel, P., Robe, M.C., Maurenas, J.M.,
642	Faivre-Pierret, R., Martin, D., Sabroux, J.C., Zettwoog, P., 1991. Eruptive and diffuse
643	emissions of CO ₂ from Mount Etna. Nature 35, 387–391.
644	Alparone, S., Behncke, B., Giammanco, S., Neri, M., Privitera, E., 2005. Paroxysmal summit
645	activity at Mt.Etna (Italy) monitored through continuous soil radon measurements.
646	Geophys. Res. Lett. 32 (16). doi: 10.1029/2005GL023352.
647	Arrighi, S., Rosi, M., Tanguy, J., Courtillot, V., 2004. Recent eruptive history of Stromboli
648	(Aeolian Islands, Italy) determined from high-accuracy archeomagnetic dating. Geophys.
649	Res. Lett. 31. doi: 10.1029/2004GL020627.
650	Baldi, P., Fabris, M., Marsella, M., Monticelli, R. 2005. Monitoring the morphological evolution
651	of the Sciara del Fuoco during the 2002-2003 Stromboli eruption using multi-temporal
652	photogrammetry. ISPRS J. Photogramm. Remote Sens. 59 (4), 199–211.
653	Barberi, F., Rosi, M., Sodi, A., 1993. Volcanic hazard assessment at Stromboli based on review

Aiuppa, A., Federico, C., Giudice, G., Giuffrida, G., Guida, R., Gurrieri, S., Liuzzo, M., Moretti,

R., Papale, P., 2009. The 2007 eruption of Stromboli volcano: insights from real-time

of historical data. Acta Vulcanol. 3, 173-187.

655	Barberi, F., Civetta, L., Rosi, M., Scandone, R., 2009. Chronology of the 2007 eruption of	
656	Stromboli and the activity of the Scientific Synthesis Group. J. Volcanol. Geotherm. Res.	Formatted: English (United Kingdom)
657	182 (3-4), 123-130.	
658	Baubron, J.C., Rigo, A., Toutain, J.P., 2002: Soil gas profiles as a tool to characterize active	
659	tectonic areas: the Jaut Pass example (Pyrenees, France). Earth. Planet. Sci. Lett., 196, 69-	
660	<u>81.</u>	
661	Bertagnini, A., Métrich, N., Landi, P., Rosi, M., 2003. Stromboli volcano (Aeolian Archipelago,	Formatted: English (United States)
662	Italy): An open window on the deep-feeding system of a steady state basaltic volcano. J.	Formatted: Indent: Left: 0 cm
663	Geophys. Res. B Solid Earth 108 (7), ECV 4-1-4-15.	
664	Bonetti, R., Capra, L., Chiesa, C., Guglielmetti, A., Migliorini, C., 1991. Energy response of	
665	LR115 cellulose nitrate to α -particle beams. Nucl. Radiat. Measur. 18, 321-338.	
666	Brusca, L., Inguaggiato, S., Longo, M., Madonia, P., Maugeri, R., 2004. The 2002-2003	
667	eruption of Stromboli (Italy): Evaluation of the volcanic activity by means of continuous	
668	monitoring of soil temperature, CO ₂ flux, and meteorological parameters. Geochem.	
669	Geophys. Geosyst. 5 (12), Q12001. doi:10.1029/2004GC000732.	
670	Burton, M.R., Caltabiano, T., Murè, F., Salerno, G., Randazzo, D., 2009. SO ₂ flux from	
671	Stromboli during the 2007 eruption: Results from the FLAME network and traverse	
672	measurements. J. Volcanol. Geotherm. Res. 182 (3-4), 214-220.	Formatted: English (United Kingdom)
673	Burton, M., Sawyer, G., Granieri, D., 2013. Deep carbon emissions from volcanoes. Rev.	
674	Mineral. Geochem. 75, 323-354.	
675	Calvari S., Bonaccorso, A., Madonia, P., Neri, M., Liuzzo, M., Salerno, G.G., Behnke, B.,	
676	Caltabiano, T., Cristaldi, A., Giuffrida, G., La Spina, A., Marotta, E., Ricci, T.,	
677	Spampinato, L., 2014. Major eruptive style changes induced by structural modifications of	

678	a shallow conduit system: the 2007-2012 Stromboli case. Bull. Volcanol. 76, 841. doi:	
679	10.1007/s00445-014-0841-7.	
680	Capasso, G., Carapezza, M.L., Federico, C., Inguaggiato, S., Rizzo, A., 2005. Geochemical	Formatted: English (United Kingdom)
681	monitoring of the 2002-2003 eruption at Stromboli volcano (Italy): precursory changes in	
682	the carbon and helium isotopic composition of fumarole gases and thermaland thermal	
683	waters. Bull. Volcanol. 68, 118–134. doi: 10.1007/s00445-005-0427-5.	Formatted: English (United Kingdom)
684	Carapezza, M.L., Federico, C. 2000. The contribution of fluid geochemistry to the volcano	
685	monitoring of Stromboli. J. Volcanol. Geotherm. Res. 95 (1-4), 227-245. doi:	Formatted: English (United States)
686	10.1016/S0377-0273(99)00128-6.	
687	Carapezza, M.L., and D. Granieri (2004). CO2 soil flux at Vulcano (Italy): comparison of active	
688	and passive methods and application to the identification of actively degassing structure,	
689	<u>Appl. Geochem. 19, 73-88.</u>	
690	Carapezza, M.L., Inguaggiato, S., Brusca, L., Longo, M. 2004. Geochemical precursors of the	Formatted: Indent: Left: 0 cm
691	activity of an open-conduit volcano: The Stromboli 2002-2003 eruptive events. Geophys.	
692	Res. Lett. 31 (7), L07620. doi: 10.1029/2004GL019614.	
693	Carapezza, M.L., Ricci, T., Ranaldi, M., Tarchini, L., 2009. Active degassing structures of	
694	Stromboli and variations in diffuse CO ₂ output related to the volcanic activity. J. Volcanol.	
695	Geotherm. Res. 182 (3–4), 231–245.	
696	Carapezza, M.L., Cigolini C., Coppola D., Laiolo M., Ranaldi, M., Ricci T., Tarchini, L., 2010.	
697	The role played by the environmental factors on diffuse soil degassing at Stromboli	
698	volcano. IAVCEI - Cities on Volcanoes 6th, CoV6/1.3/P/47, Tenerife, Canary Islands,	
699	Spain.	
700		

701	Chouet, B., Saccorotti, G., Martini, M., Dawson, P., De Luca, G., Milana, G., Scarpa, R., 1997.	
702	Source and path effects in the wavefields of tremor and explosions at Stromboli Volcano,	
703	<u>Italy. J. Geophys. Res. 102, 15,129 – 15,150.</u>	
704	Cigolini, C., Salierno, G., Gervino, G., Bergese, P., Marino, C., Russo, M., Prati, P., Ariola, V.,	Formatted: Indent: Left: 0 cm
705	Bonetti, R., Begnini, S., 2001. High-resolution Radon Monitoring and Hydrodynamics at	
706	Mount Vesuvius. Geophys. Res. Lett. 28 (21), 4035-4039.	
707	Cigolini, C., Gervino, G., Bonetti, R., Conte, F., Laiolo, M., Coppola, D., Manzoni, A., 2005.	
708	Tracking precursors and degassing by radon monitoring during major eruptions at	
709	Stromboli Volcano (Aeolian Islands, Italy). Geophys. Res. Lett. 32, L12308. doi:	
710	10.1029/2005GL022606.	
711	Cigolini, C., Laiolo, M., Coppola, D., 2007. Earthquake-volcano interactions detected from	
712	radon degassing at Stromboli (Italy). Earth Planet. Sci. Lett. 257, 511-525.	
713	Cigolini, C., Laiolo, M., Bertolino, S., 2008. Probing Stromboli volcano from the mantle to	
714	paroxysmal eruptions. In: Zellmer, G., Hammer, J., (Eds.), Dynamics of Crustal Magma	
715	Transfer, Storage, and Differentiation – integrating geochemical and geophysical	
716	constraints. Geological Society, London, Special Publication, 304, pp. 33-70.	
717	Cigolini C. Poggi, P., Ripepe, M., Laiolo M., Ciamberlini C., Delle Donne, D., Ulivieri, G.,	
718	Coppola D., Lacanna, G., Marchetti, E., Piscopo, D., Genco, R., 2009. Radon surveys and	
719	real-time monitoring at Stromboli volcano: Influence of soil temperature, atmospheric	
720	pressure and tidal forces on ²²² Rn degassing. J. Volcanol. Geotherm. Res. 184 (3-4), 381-	Formatted: Superscript
721	388.	

722	Cigolini C., Laiolo, M., Ulivieri, G., Coppola, D., Ripepe, M., 2013. Radon mapping, automatic
723	measurements and extremely high ²²² Rn emissions during the 2002–2007 eruptive Formatted: Superscript
724	scenarios at Stromboli volcano. J. Volcanol. Geotherm. Res. 264, 49- 65.
725	Cigolini, C., Laiolo, M., Coppola, D., 2014. Revisiting the last major eruptions at Stromboli
726	volcano: inferences on the role of volatiles during magma storage and decompression. In:
727	Zellmer, G.F., Edmonds, M., Straub, S.M. (Eds.), The Role of Volatiles in the Genesis,
728	Evolution and Eruption of Arc Magmas. Geological Society, London, Special Publication,
729	304, pp. 33-70.
730	Coppola, D., Piscopo, D., Laiolo, M., Cigolini, C., Delle Donne, D., Ripepe, M., 2012. Radiative
731	heat power at Stromboli volcano during 2000-2011: twelve years of MODIS observations.
732	J. Volcanol. Geotherm. Res. 215-216, 48-60, doi: 10.1016/j.jvolgeores.2011.12.001.
733	Corazzato, C., Francalanci, L., Menna, M., Petrone, C.M., Renzulli, A., Tibaldi, A., Vezzoli, L.,
734	2008. What controls sheet intrusion in volcanoes? Structure and petrology of the Stromboli
735	sheet complex, Italy. J. Volcanol. Geotherm. Res. 173 (1-2), 26-54.
736	De Gregorio, S., Camarda, M., Gurrieri, S., Favara, R., 2014. Change in magma supply
737	dynamics identified in observations of soil CO ₂ emissions in the summit area of Mt. Etna.
738	Bull. Volcanol.76 (8), 1-8. doi: 10.1007/s00445-014-0846-2.
739	Faber, E., Morán, C., Poggenburg, J., Garzón, G., Teschner, M., 2003. Continuous gas
740	monitoring at Galeras Volcano, Colombia: First evidence, J. Volcanol. Geotherm. Res., Formatted: English (United States)
741	<u>125 (1-2), 13-23.</u>
742	Federico, C., Brusca, L., Carapezza, M.L., Cigolini, C., Inguaggiato, S., Rizzo, A., Rouwet, D.,
743	2008. Geochemical prediction of the 2002–2003 Stromboli eruption from variations in CO ₂ Formatted: English (United Kingdom)
744	and Rn ²²² Rn emissions and in Helium and Carbon isotopes. In: Calvari, S., Inguaggiato, S.,

745	Ripepe, M. &_Rosi, M. (Eds.), The Stromboli volcano: an integrated study of the 2002-
746	2003 eruption. AGU, Geophysical Monograph Series, Washington D.C. 182, pp. 117-128.
747	Finizola, A., Sortino, F., Lenat, J.F., Valenza, M., 2002. Fluid circulation at Stromboli volcano
748	(Aeolian Islands, Italy) from self-potential and CO ₂ surveys, J. Volcanol, Geotherm, Res.

749 116, 1-18.

Finizola, A., Sortino, F., Lénat, J.F., Aubert, M., Ripepe, M., Valenza, M., 2003. The summit
hydrothermal system of Stromboli. New insights from self-potential, temperature, CO₂ and
fumarolic fluid measurements, with structural and monitoring implications. Bull. Volcanol.
65, 486–504.

- Finizola, A., Revil, A., Rizzo, E., Piscitelli, S., Ricci, T., Morin, J., Angeletti, B., Mocochain, L.,
 Sortino, F., 2006. Hydrogeological insights at Stromboli volcano (Italy) from geoelectrical,
 temperature, and CO₂ soil degassing investigations. Geophys. Res. Lett. 33 (17), L17304.
- Finizola, A., Aubert, M., Revil, A., Schütze, C., Sortino, F., 2009. Importance of structural history in the summit area of Stromboli during the 2002–2003 eruptive crisis inferred from temperature, soil CO₂, self-potential, and electrical resistivity tomography. J. Volcanol. Geotherm. Res. 183 (3–4), 213–227.
- Francalanci, L, Manetti, P, Peccerillo, A., 1989. Volcanological and magmatological evolution
 of Stromboli volcano (Aeolian Islands): the roles of fractional crystallisation, magma
 mixing, crustal contamination and source heterogeneity. Bull. Volcanol. 51, 355-378
- Francalanci, L., Tommasini, S., Conticelli, S., 2004. The volcanic activity of Stromboli in the
 1906-1998 AD period: Mineralogical, geochemical and isotope data relevant to the
 understanding of the plumbing system. J. Volcanol. Geotherm. Res. 131 (1-2), 179-211.

Formatted: English (United Kingdom)

767	Gauthier, P.J., Condomines, C., 1999. ²¹⁰ Pb- ²²⁶ Ra radioactive disequilibria in recent lavas and	
768	radon degassing: inferences on the magma chamber dynamics at Stromboli and Merapi	
769	volcanoes. Earth Planet. Sci. Lett. 172, 111-126.	
770	Giammanco, S., Sims, K.W., Neri, M., 2007. Measurements of 220 Rn and 222 Rn and CO ₂	
771	emissions in soil and fumarole gases on Mt. Etna Volcano (Italy): implications for gas	
772	transport and shallow ground fracture. Geochem. Geophys. Geosyst. 8, Q10001. doi:	Formatted: English (United Kingdom)
773	10.1029/2007GC001644.	
774	Girault, F., Perrier, F., 2012. Estimating the importance of factors influencing the radon-222 flux	
775	from building walls. Sci. Tot. Environm. 433, 247-263.	
776	Girault, F., Perrier, F., 2014. The Syabru-Bensi hydrothermal system in central Nepal: 2.	
777	Modeling and significance of the radon signature, J. Geophys. Res. 119, 4056-4089.	
778	Girault, F., Perrier, F., Crockett, R., Bhattarai, M., Koirala, B.P., France-Lanord, C., Agrinier, P.,	
779	Ader, M., Fluteau, F., Gréau, C., Moreira, M., 2014. The Syabru-Bensi hydrothermal	
780	system in central Nepal: 1. Characterization of carbon dioxide and radon fluxes. J.	Formatted: Italian (Italy)
781	Geophys. Res. 119, 4017-4055.	
782	Granieri, D., Chiodini, G., Marzocchi, W., Avino, R., 2003. Continuous monitoring of CO2 soil	
783	diffuse degassing at Phlegraean Fields (Italy): influence of environmental and volcanic	
784	parameters, Earth Planet. Sci. Lett. 212, 167-179.	
785	Gründel, M., Postendörfer. J., 2003. Characterization of an electronic Radon gas personal	
786	Dosimeter. Rad. Prot. Dosim. 107 (4), 287-292.	
787	Hayashi, K., Yasuoka, Y., Nagahama, H., Muto, J., Ishikawa, T., Omori, Y., Suzuki, T., Homma,	
788	Y., Mukai, T., 2015. Normal seasonal variations for atmospheric radon concentration: a	
789	sinusoidal modelJ Environ Radioact. 53, 139:149. doi: 10.1016/j.jenvrad.2014.10.007.	

790	Hernandez, P., Perez, N., Salazar, J., Reimer, M., Notsu, K., Wakita, H., 2004. Radon and	Formatted: Indent: Left: 0 cm
791	helium in soil gases at Cañadas caldera, Tenerife, Canary Islands, Spain. J. Volcanol.	
792	Geotherm. Res. 131, 59-76.	
793	Hornig-Kjarsgaard., I., Keller., J., Koberski, U., Stadbauer, E., Francalanci, L., Lenhart, R.,	
794	1993. Geology, stratigraphy and volcanological evolution of the island of Stromboli,	
795	Aeolian arc, Italy. Acta Vulcanol. 3, 21–68.	
796	Inguaggiato, S., Jácome Paz, M.P., Mazot, A., DelgadoGranados, H., Inguaggiato, C., Vita, F.	
797	2013. CO ₂ output discharged from StromboliIslandStromboli Island (Italy). Chem. Geol.	
798	339, 52-60.	
799	Kotrappa, P., Dempsey, J.C., Stieff, L.R., 1993. Recent advances in electret ion chamber	
800	technology. Radiat. Protect. Dosim. 47, 461-464.	
801	Laiolo, M., Cigolini, C., Coppola, D., Piscopo, D., 2012. Developments in real-time radon	
802	monitoring at Stromboli volcano. J. Environm. Radioact. 105, 21-29.	
8 03 .	pina, A., Burton, M.R., Harig, R., Mure, F., Rusch, P., Jordan, M., Caltabiano, T., 2013. New	
804	insights into volcanic processes at Stromboli from Cerberus, a remote controlled open path FTIR	
805	scanner system, J. Volcanol. Geotherm. Res. 249, 66-76.	Formatted: English (United States)
		Formatted: Italian (Italy)
806	Madonia, P., Fiordilino, E., 2013. Time variability of low-temperature fumaroles at Stromboli	Formatted: Indent: Left: 0 cm
807	island (Italy) and its application to volcano monitoring. Bull. Volcanol. 75, 776. doi:	Formatted: English (United Kingdom)
808	10.1007/s00445-013-0776-4.	
809	Makario Londoño, J., 2009. Radon and CO ₂ emissions in different geological environments as a	
810	tool for monitoring volcanic and seismic activity in central part of Colombia. Boletin de	
811	<u>Geologia, 31(2), 83-95.</u>	

812	Marchetti, E., Ripepe, M., 2005. Stability of the seismic source during effusive and explosive						
813	activity at Stromboli Volcano. Geophys. Res. Lett. 32 (3), 1-5.						
814	doi:10.1029/2004GL021406.						
815	Martin-Luis, M.C., Steinitz, G., Soler, V., Quesada, M.L., Casillas, R. 2015. ²²² Rn and CO ₂ at						
816	Las Cañadas Caldera (Tenerife, Canary Islands). Eur. Phys. J. Special Topics 224 (4), 641-						
817	<u>657.</u>						
818	Métrich, N., Bertagnini, A., Landi, P., Rosi, M., 2005. Triggering mechanism at the origin of Formatted: English (United States)						
819	paroxysm at Stromboli (Aeolian Archipelago, Italy): The 5 April 2003 eruption. Geophys.						
820	Res. Lett. 32, L10305. doi: 10.10129/2004GL022257.						
821	Métrich, N.A., Bertagnini, A. & Di Muro, A. 2010. Conditions of Magma Storage, Degassing						
822	and Ascent at Stromboli: New Insights into the Volcano Plumbing System with Inferences						
823	on the Eruptive Dynamics. J. Petrol. 51, 603-626.						
824	Mogro-Campero, A., Fleischer, R.L., 1977. Subterrestrial fluid convection: a hypothesis for long						
825	distance migration of radon within the earth. Earth Planet. Sci. Lett. 34, 321-325.						
826	Nazaroff, W.W., 1992. Radon transport from soil to air. Rev. Geophys. 30, 137-160. doi:						
827	10.1029/92RG00055.						
828	Padilla, G.D., <u>Hernández, P.A., Padrõn, E., Barrancos, J., Pérez, N.M., Melián, G., Nolasco,</u>						
829	D., Dionis, S., Rodríguez, F., Calvo, D., Hernández, I., 2013. Soil gas radon emissions						
830	and volcanic activity at El Hierro (Canary Islands): The 2011-2012 submarine eruption.						
831	Geochem. Geophys. Geosyst., 14 (2), 432-447.						
832	Padilla, G.D., Hernandez, P.A., Pérez, N.M., Pereda, E., Padron, E., Melian, G., Barrancos, J.,						
833	Rodriguez, F., Dionis, S., Calvo, D., Herrera, M., Strauch, W., Munoz, A., 2014.						
834	Anomalous diffuse CO ₂ emissions at the Masaya volcano (Nicaragua) related to seismic-						

835	volcano unrest. Pure Appl. Geophys. 171 (8), 1791-1804. doi: 10.1007/s00024-013-0756-	
836	9.	
837	Padrón, E., Padilla, G., Hernández, P.A., Pérez, N.M., Calvo, D., Nolasco, D., Barrancos, J.,	
838	Melián, G.V., Dionis, S., Rodríguez, F., 2013. Soil gas geochemistry in relation to eruptive	
839	fissures on Timanfaya volcano, LanzaroteIslandLanzarote Island (Canary Islands, Spain).	
840	J. Volcanol. Geotherm. Res. 250, 91–99.	Formatted: English (United Kingdom)
841	Pan, V., Holloway, J.R., Hervig, R.L., 1991. The pressure and temperature dependence of carbon	
842	dioxide solubility in tholeiitic basalt melts. Geochim. Cosmochim. Acta 55, 1587–1595.	
843	Papachristodoulou, C., Ioannides, K., Spathis, S. (2007). The effect of moisture content on	
844	radon diffusion through soil: Assessment in laboratory and field experiments. Health Phys	
845	<u>92 (3), 257-264.</u>	
846	Papale, P., Moretti, R., Barbato, D., 2006. The compositional dependence of the multicomponent	Formatted: Indent: Left: 0 cm
847	volatile saturation surface in silicate melts. Chem. Geol. 229, 78–95.	
848	Pérez, N.M., Hernández, P.A., Padrón, E., Melián, G., Marrero, R., Padilla, G., Barrancos, J.,	
849	Nolasco, D., 2007. Precursory subsurface ²²² Rn and ²²⁰ Rn degassing signatures of the 2004	
850	seismic crisis at Tenerife, Canary Islands. Pure Appl. Geophys. 164, 2431-:2448, doi:	
851	10.1007/s00024-007-0280.	
852	Perrier, F., Girault, F., 2012. Harmonic response of soil radon-222 flux and concentration	
853	induced by barometric oscillations. GeophysJ. Int., doi: 10.1093/gji/ggt280.	
854	Perrier, F., Richon, P., Sabroux, J.C., 2009. Temporal variations of radon concentration in the	Formatted: French (France)
855	saturated soil of Alpine grassland: The role of groundwater flow. Sci. Tot. Environm. 407,	Formatted: Indent: Left: 0 cm
856	2361-2371.	

857	Pinault, J.L., Baubron, J.C., 1996. Signal processing of soil gas radon, atmospheric pressure and	
858	soil temperature data: a new approach for radon concentration modelling. J. Geophys. Res.	
859	B: Solid Earth 101 (2), 3157-3171.	
860	Revil, A., Finizola, A., Ricci, T., Delcher, E., Peltier, A., Barde-Cabusson, S., Avard, G., Bailly,	
861	T.,Bennati, L., Byrdina, S., Colonge, J., Di Gangi, F., Douillet, G., Lupi, M., Letort, J.,	
862	Tsang Hin Sun, E., 2011. Hydrogeology of Stromboli volcano, Aeolian Islands (Italy) from	
863	the interpretation of resistivity tomograms, self-potential, soil temperature and soil CO2	
864	concentration measurements. Geophys. J. Int. 186 (3), 1078-1094.	Formatted: English (United Kingdom)
865	Rinaldi, A.P., Vandemeulebrouck, J., Todesco, M., Viveiros, F. 2012. Effects of atmospheric	
866	conditions on surface diffuse degassing. J. Geophys. Res. Solid Earth 117, B11201. doi:	
867	<u>10.1029/2012JB009490.</u>	
868	Ripepe, M., Delle Donne, D., Lacanna, G., Marchetti, E., and Ulivieri, G., 2009. The onset of	Formatted: Italian (Italy)
		Formatted: Italian (Italy)
869	the 2007. Infrasonic monitoring at Stromboli Volcano during the 2003 effusive eruption:	Formatted: Italian (Italy)
870	insights on the explosive and degassing process of recorded by an open conduit	Formatted: Indent: Left: 0 cm
		Formatted: English (United Kingdom)
871	systemintegrated geophysical network. J. Volcanol. Geotherm. Res. 182(3-4): 131-136- J.	Formatted: English (United Kingdom)
872		
	Geophys. Res112, B09207. doi: 10.1029/2006jB004613,2007.	Formatted: English (United Kingdom)
873	Geophys. Res112, B09207. doi: 10.1029/2006jB004613,2007. Rizzo, A., Grassa, F., Inguaggiato, S., Liotta, M., Longo, M., Madonia, P., Brusca, L., Capasso,	
873 874		
	Rizzo, A., Grassa, F., Inguaggiato, S., Liotta, M., Longo, M., Madonia, P., Brusca, L., Capasso,	Formatted: English (United Kingdom)
874	Rizzo, A., Grassa, F., Inguaggiato, S., Liotta, M., Longo, M., Madonia, P., Brusca, L., Capasso,G., Moricia, S., Rouwet, D., Vita, F., 2009. Geochemical evaluation of observed changes	
874 875	Rizzo, A., Grassa, F., Inguaggiato, S., Liotta, M., Longo, M., Madonia, P., Brusca, L., Capasso, G., Moricia, S., Rouwet, D., Vita, F., 2009. Geochemical evaluation of observed changes in volcanic activity during the 2007 eruption at Stromboli (Italy). J. Volcanol. Geotherm.	

879	(Italy): Inferences from soil CO_2 flux and ${}^{3}He/{}^{4}He$ ratio in thermal waters. Geophys. Res.	
880	Lett. 42, doi: 10.1002/2014GL062955.	
881	Rosi, M., Bertagnini, A., Landi, P., 2000. Onset of persisting activity at Stromboli Volcano	
882	(Italy). Bull. Volcanol. 62, 294-300.	
883	Sakoda, A., Ishimori, Y., Hanamoto, K., Kataoka, T., Kawabe, A., Yamaoka, K., 2010.	
884	Experimental and modeling studies of grain size and moisture content effects on radon	
885	emanation. Radiat. Measurem. 45, 204-210.	
886	Salazar, J.M.L., Hernández, P.A., Pérez, N.M., Melián, G., Álvarez, J., Segura, F., Notsu, K.,	
887	2001. Diffuse emission of carbon dioxide from Cerro Negro volcano, Nicaragua, Central	
888	America. Geophys. Res. Lett. 28 (22), 4275-4278.	
889	Silva, C., Ferreira, T., Viveiros, F. &., Allard P., 2015. Soil radon (²²² Rn) monitoring at Furnas	Formatted: Superscript
890	Volcano (São Miguel, Azores): Applications and challenges. Eur. Phys. J. Spec.	
891	Top.Special Topics 224 (4), 659-686. doi: 10.1140/epjst/e2015-02398-6.	
892	Siniscalchi, A., Tripaldi, S., Neri, M., Giammanco, S., Piscitelli, S., Balasco, M., Behncke, B.,	
893	Magri, C., Naudet, V., Rizzo, E., 2010. Insights into fluid circulation across the Pernicana	
894	Fault (Mt. Etna, Italy) and implications for flank instability. J. Volcanol. Geotherm. Res.	Formatted: English (United Kingdom)
895	193, 137-142	
896	Steinitz, G., Piatibratova, O., Kotlarsky, P., 2011. Possible effect of solar tides on radon signals.	
897	J. Environm. Radioact. 102 (8), 749 – 765. doi: 10.1016/j.jenvrad.2011.04.002.	
898	Tibaldi, A., 2003. Influence of cone morphology on dykes, Stromboli, Italy. J. Volcanol.	Formatted: English (United Kingdom)
899	Geotherm. Res. 126, 79–95.	Formatted: Indent: Left: 0 cm
900	Tibaldi, A., 2004. Major changes in volcano behaviour after a sector collapse: Insights from	
901	Stromboli, Italy. Terra Nova 16 (1), 2-8.	

902	Tibaldi, A., Corazzato, C., Marani, M., Gamberi, F., 2009. Subaerial-submarine evidence of	
903	structures feeding magma to Stromboli Volcano, Italy, and relations with edifice flank	
904	failure and creep. Tectonophys. 469 (1-4), 112-136.	
905	Tinti, S., Maramai, A., Armigliato, A., Graziani, L., Manucci, A., Pagnoni, G., Zaniboni, F.,	
906	2006. Observations of physical effects from tsunamis of December 30, 2002 at Stromboli	
907	volcano, southern Italy. Bull. Volcanol. 68, 450-461.	
908	Toutain, J. P., Baubron, J. C., 1999. Gas geochemistry and seismotectonics: a review.	
909	Tectonophys. 304, 1–27.	
910	Vandeginste, B.G.M., Massart, D.L., Buydens, L.M.C., De Jong, S., Lewi, P.J., Smeyers-	
911	Verbeke, J., 1988. Handbook of Chemometrics and Qualimetrics: Part B. Elsevier,	
912	Amsterdam.	
913	Viveiros, F., Ferreira, T., Cabral Vieira, J., Silva, C., Gaspar, J.L., 2008. Environmental	
914	influences on soil CO ₂ degassing at Furnas and Fogo volcanoes (São Miguel Island,	
915	Azores archipelago). J. Volcanol. Geotherm. Res. 177, 883–893.	Formatted: English (United Kingdom)
916		
917		
918		
919	Viveiros, F., Vandemeulebrouck, J., Rinaldi, A.P., Ferreira, T., Silva, C., Cruz, J.V., 2014.	
920	Periodic behavior of soil CO2 emissions in diffuse degassing areas of the Azores	
921	archipelago: Application to seismovolcanic monitoring. J. Geophys. Res. 119, 7578-7597.	
922	<u>doi:10.1002/2014JB011118.</u>	

923	Zimmer, M., Erzinger, J., 2003. Continuous H ₂ O, CO ₂ , ²²² Rn and temperature measurements on
924	Merapi Volcano, Indonesia. 8J. Volcanol. Geotherm. Res. 125 (1-2), 25-38. doi:
925	<u>10.1016/S0377-0273(03)00087.</u>
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929	Fig. 1. Digital Elevation Model of Stromboli island (Island (from Baldi et al., 2005) with the overprint of	(Formatted: Font: Bold
930	the major faults and collapsed sectors (simplified from Finizola et al., 2002 and Tibaldi et al., 2009). The		
931	locationLocations of the Volcano Observatory (COA) and of ²²² Rn activitythe radon and soil CO ₂ flux	(Formatted: Subscript
932	stations is also indicated are reported.		
933			
934	Fig. 2 (a) Map of radon activity measured in March 10-18, 2007 on the NE flank of Stromboli. The		Formatted: Font: Bold
935	pointsFull circles indicate measurement sites; the contour lines of radon emissions have been obtained by		Formatted: Space Before: 6 pt, After: 6 pt
936	kriging (Cigolini et al., 2013). Triangle <u>The triangle</u> indicates the <u>location of the</u> ²²² Rn automatic station.		
937	Dotted white rectangle marks the area of the $\frac{\text{CO}_2}{\text{CO}_2}$ flux map of March 2007 reported in b) and		
938	where stars). Stars are the sites of the automatic CO ₂ stations (Carapezza et al., 2009).	(Formatted: Subscript
938 939	where stars). Stars are the sites of the automatic CO_2 stations (Carapezza et al., 2009).	(Formatted: Subscript
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939940941942	Fig. 3. Time-series of daily averages of radon activities ²²² Rn activity (a) and soil CO ₂ fluxesflux (b) recorded hourly from April 2007 to September 2011: (grey dots). Black curves and the straight dotted lines represents the daily and the annual average, respectively. Histograms show the multi-modal		Formatted: Font: Bold Formatted: Space Before: 6 pt, After: 6 pt
 939 940 941 942 943 944 	Fig. 3. Time-series of daily averages of radon activities ²²² Rn activity (a) and soil CO ₂ fluxesflux (b) recorded hourly from April 2007 to September 2011- (grey dots). Black curves and the straight dotted lines represents the daily and the annual average, respectively. Histograms show the multi-modal distributions.		Formatted: Font: Bold Formatted: Space Before: 6 pt, After: 6 pt
 939 940 941 942 943 944 945 	Fig. 3. Time-series of daily averages of radon activities ²²² Rn activity (a) and soil CO ₂ fluxesflux (b)* recorded hourly from April 2007 to September 2011 ; (grey dots). Black curves and the straight dotted lines represents the daily and the annual average, respectively. Histograms show the multi-modal distributions. Fig. 4. Time-series of the daily mobile average of the main environmental parameters measured hourly*		Formatted: Font: Bold Formatted: Space Before: 6 pt, After: 6 pt Formatted: Subscript
 939 940 941 942 943 944 	Fig. 3. Time-series of daily averages of radon activities ²²² Rn activity (a) and soil CO ₂ fluxesflux (b) recorded hourly from April 2007 to September 2011- (grey dots). Black curves and the straight dotted lines represents the daily and the annual average, respectively. Histograms show the multi-modal distributions.		Formatted: Font: Bold Formatted: Space Before: 6 pt, After: 6 pt Formatted: Subscript Formatted: Font: Bold Formatted: Space Before: 6 pt,

948	Fig. 5. Soil CO ₂ flux vs. wind direction. Red and blue circles refer to data acquired with wind speed		Formatted: Font: Bold
949	above or below 8 m/s respectively. Note that most of the high soil CO ₂ fluxes are recorded for wind speed		
950	>8m/s in the SE sector.		
951	Fig. 6. Spectral amplitude for soil CO ₂ flux (a), ²²² Rn concentration (b), atmospheric pressure (c) and air		
952	temperature (d). The analyses were made over one year of hourly data (see text for details).		
953	<u>Fig. 7.</u> Bulk annual trend of Radonradon concentration and soil CO ₂ flux (a) retrieved from the mean-		Formatted: Space Before: 6 pt, After: 6 pt
954	values measured each day in the 4 years monitoring. Results are compared with the annual trend of soil		Formatted: Subscript
955	and air temperatures (b).		
956			
957	Fig. 68 . Seasonal variations of the <u>average</u> correlation coefficients (<u>R</u>) between main atmospheric factors		Formatted: Font: Bold
958	and ²²² Rn concentration (above) and soil CO ₂ flux- (below). Air T: air temperature; Soil T: soil	$\langle \rangle$	Formatted: Space Before: 6 pt, After: 6 pt
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959	temperature; Air P: barometric pressure; Soil H: soil humidity; Wind Sp: wind speed; Wind Dir: wind		
960	direction. Air H: air humidity. SS and FW refer to Spring-Summer and Fall-Winter period, respectively.		
961			
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962	Fig. 79. Results from Multiple Linear Regression (a) and Principal Component Regression (b) for radon	\sim	Formatted: Space Before: 6 pt,
963	activity (above) and soil CO ₂ flux (below) during the 4 $\frac{1}{2}$ years of monitoring. Data are reported as daily		After: 6 pt Formatted: Font: Bold
964	averageaverages. The observed, calculated and residual values are shownindicated.		
965			
966	Fig. 8. Plot <u>10. Time series</u> of the hourly calculated residuals of ²²² Rn (a)concentration and soil CO ₂ flux		Formatted: Font: Bold
967	(b) obtained by MLR (a) and PCR (b) methods with indication of the main volcanic and seismic events		
968	recorded <u>occurred</u> at Stromboli from April 2007 to September 2011-(see legend). The residuals were		
969	obtained by the results of PCR outcoming from about 95% of the dataset. Vertical axes express the		
970	standard deviation from the mean; only values $\geq 2 \frac{\text{sigmage}}{2}$ are plotted. Grey <u>filed field</u> marks the time span		
971	where ²²² Rn <u>when no radon</u> data are not available. 44		

972 <u>were recorded.</u>

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		average	SD	max	min	data no.
Rn station (LSC site)						
²²² Rn	$Bq m^{-3}$	2051.00	1196.00	7879.00	0	35,178
Soil T (-1m)	$^{\circ}C$	17.44	5.70	30.09	5.90	35,178
Air P	hPa	951.38	6.40	974.37	925.16	35,178
CO ₂ station (NC site)						
Soil CO ₂ flux	$g m^{-2} da y^{-1}$	585.45	643.17	7021.72	10.92	39,412
Air RH	%	77.62	16.02	100.00	12.36	39,412
Air T	$^{\circ}C$	16.45	6.55	38.78	0	39,412
Air P	hPa	954.63	5.66	973.45	921.69	39,412
Soil T (-0,5m)	$^{\circ}C$	19.87	5.04	28.86	11.89	39,412
Soil RH (-0,5m)	%	12.30	2.11	21.38	8.63	39,412
Wind direction	°N			359	0	39,412
Wind speed	$m s^{-1}$	3.67	2.96	21.91	0.02	39,412

	222	
\mathbf{T} -LL-1 D \mathbf{D} -	11 + 12 + 12 = 100	$A = A^{2} = $
Lanie I. Descriptive statistics of the main	parameters acquired by the R n and C U	² stations from April 2007 to September 2011

Annual averages	Apr. 2007 - Mar. 2008	Apr. 2008 - Mar. 2009	Apr. 2009 - Mar. 2010	Apr. 2010 - Mar. 2011
²²² Rn activity ($Bq m^{-3}$)	1777.4	1878.1	2158.4	2264.0
Soil CO_2 flux $(g m^{-2} day^{-1})$	919.9	658.0	406.9	329.5

		²²² Rn station (LSC site)			CO ₂ station (NC site)							
		²²² Rn	Soil T	Air P	Air RH	Air T	Air P	CO ₂ flux	Soil T	Soil RH	Wind direction	Wind speed
²²² Rn	$Bq m^{-3}$	1.000	-0.736	-0.282	0.387	-0.749		0.346		0.737	-0.104	0.225
Soil T	$^{\circ}C$	-0.736	1.000	0.320	-0.360	0.874		-0.532		-0.745	0.171	-0.195
Air P	hPa	-0.282	0.320	1.000	-0.305	0.330		-0.204		-0.447	0.222	-0.251
CO ₂ flux	$g m^{-2} da y^{-1}$	0.346			0.187	-0.541	-0.090	1.000	-0.566	0.370	-0.298	0.431
Air RH	%	0.387			1.000	-0.489	-0.228	0.187	-0.304	0.517	-0.089	0.021
Air T	$^{\circ}C$	-0.749			-0.489	1.000	0.155	-0.541	0.867	-0.772	0.083	-0.231
Air P	hPa	-0.114			-0.228	0.155	1.000	-0.090	0.125	-0.295	0.205	-0.227
Soil T	$^{\circ}C$	-0.694			-0.304	0.867	0.125	-0.566	1.000	-0.713	0.198	-0.193
Soil RH	%	0.737			0.517	-0.772	-0.295	0.370	-0.713	1.000	-0.120	0.211
Wind direction	$^{\circ}N$	-0.104			-0.089	0.083	0.205	-0.298	0.198	-0.120	1.000	0.148
Wind speed	$m s^{-1}$	0.225			0.021	-0.231	-0.227	0.431	-0.193	0.211	0.148	1.000

Table 2. Correlation coefficients (R) between the daily average values of the parameters measured at the monitoring stations from April 2007 to September 2011

	Coefficient B	Standard Error of B	Coefficient β	t-test	p-level			
Dependent variable	= soil CO ₂ flux; Data	n^{o} . = 39,412						
Intercept	2284.44			67.27	0.00			
Air T	-27.51	0.07	-0.43	-36.31	0.00			
Soil T	-35.27	0.05	-0.44	-37.91	0.00			
Soil RH	-48.43	0.02	-0.45	-26.20	0.00			
Wind direction	-1.47	0.98	-0.40	-52.63	0.00			
Wind speed	81.41	0.03	-0.31	96.98	0.00			
Dependent variable = 222 Rn activity; Data n ^o . = 35.178								
Intercept	-6190.63		-1.79	-7.93	0.00			
Air T	7.66	0.03	0.44	-21.47	0.00			
Soil T	-38.02	0.03	0.42	-45.24	0.00			
Atmospheric P	-48.60	0.03	0.42	10.01	0.00			
Soil RH	198.10	0.01	0.51	53.95	0.00			

Table 3. Summary of the Multiple Linear Regression analysis (MLR) on the acquired dataset

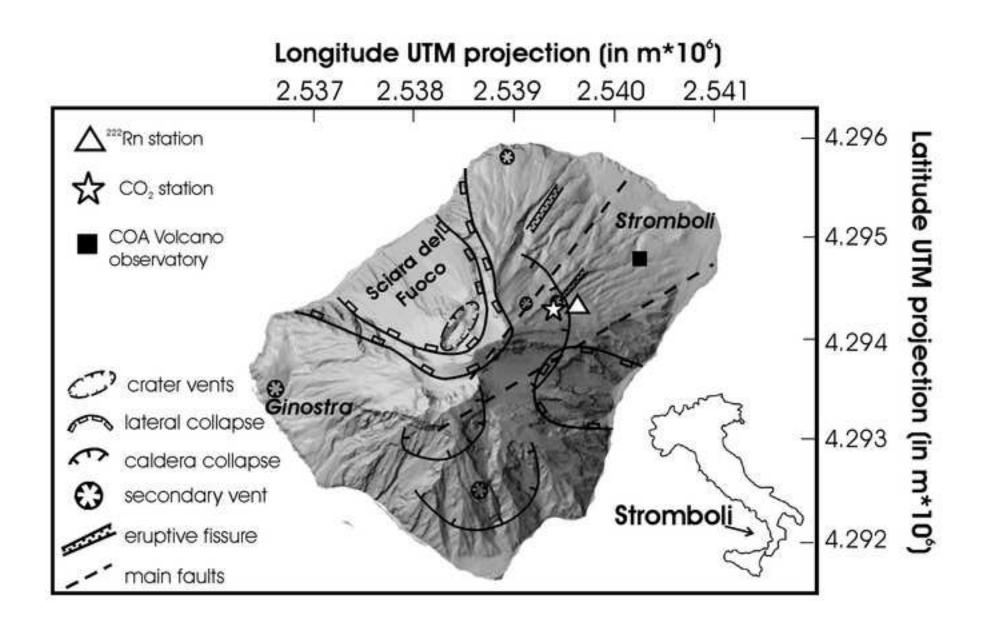
Factor1	Factor2	Factor3	Factor4	Factor5	Factor6	Factor7
0.59	-0.16	-0.05	0.55	0.43	-0.07	-0.07
-0.92	-0.03	-0.25	0.05	0.04	0.17	-0.26
-0.34	-0.06	0.82	-0.19	0.39	0.10	0.00
-0.85	0.08	-0.22	0.34	0.13	0.20	0.22
0.89	-0.03	0.03	0.00	-0.12	0.44	-0.01
-0.06	0.77	0.40	0.39	-0.28	-0.02	-0.04
0.27	0.72	-0.36	-0.34	0.41	0.02	0.00
2.900	1.147	1.076	0.854	0.621	0.276	0.126
0.414	0.164	0.154	0.122	0.089	0.039	0.018
	0.59 -0.92 -0.34 -0.85 0.89 -0.06 0.27 2.900	0.59 -0.16 -0.92 -0.03 -0.34 -0.06 -0.85 0.08 0.89 -0.03 -0.06 0.77 0.27 0.72 2.900 1.147	0.59 -0.16 -0.05 -0.92 -0.03 -0.25 -0.34 -0.06 0.82 -0.85 0.08 -0.22 0.89 -0.03 0.03 -0.06 0.77 0.40 0.27 0.72 -0.36 2.900 1.147 1.076	0.59 -0.16 -0.05 0.55 -0.92 -0.03 -0.25 0.05 -0.34 -0.06 0.82 -0.19 -0.85 0.08 -0.22 0.34 0.89 -0.03 0.03 0.00 -0.06 0.77 0.40 0.39 0.27 0.72 -0.36 -0.34 2.900 1.147 1.076 0.854	0.59 -0.16 -0.05 0.55 0.43 -0.92 -0.03 -0.25 0.05 0.04 -0.34 -0.06 0.82 -0.19 0.39 -0.85 0.08 -0.22 0.34 0.13 0.89 -0.03 0.03 0.00 -0.12 -0.06 0.77 0.40 0.39 -0.28 0.27 0.72 -0.36 -0.34 0.41 2.900 1.147 1.076 0.854 0.621	0.59 -0.16 -0.05 0.55 0.43 -0.07 -0.92 -0.03 -0.25 0.05 0.04 0.17 -0.34 -0.06 0.82 -0.19 0.39 0.10 -0.85 0.08 -0.22 0.34 0.13 0.20 0.89 -0.03 0.03 0.00 -0.12 0.44 -0.06 0.77 0.40 0.39 -0.02 0.44 -0.06 0.77 0.40 0.39 -0.28 -0.02 0.27 0.72 -0.36 -0.34 0.41 0.02 2.900 1.147 1.076 0.854 0.621 0.276

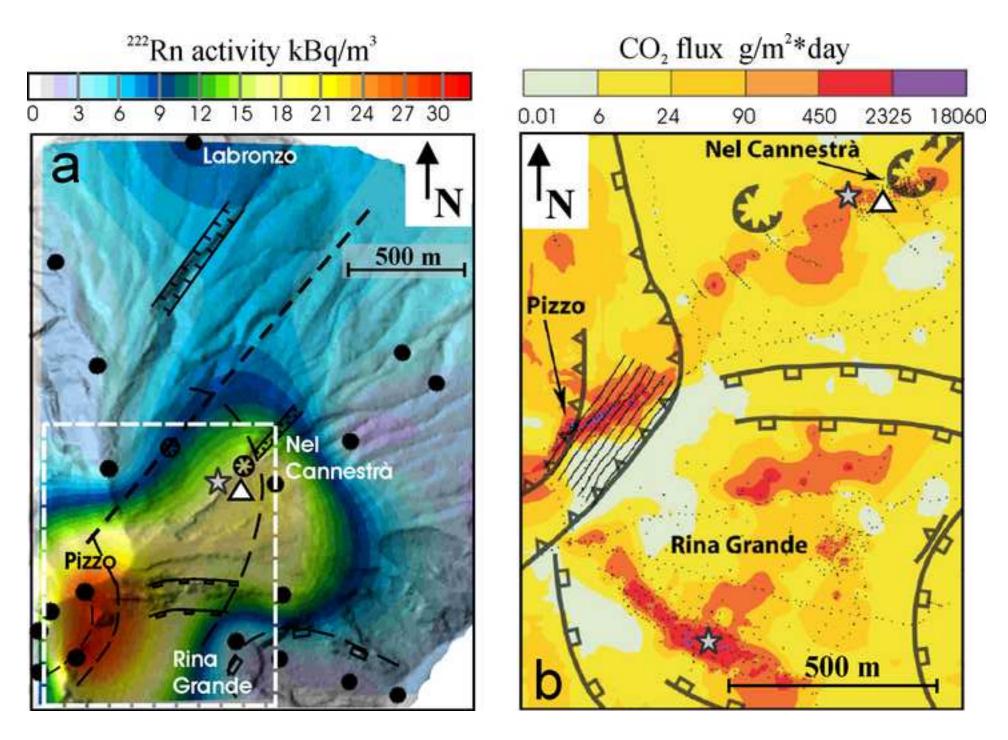
Table 4. Application of the PCR statistical method

A. Environmental variable weights on factors

B. Summary of the Principal Component Regression analysis (PCR) on the acquired dataset

	Coefficient B			t-test	p-level			
Dependent variable = soil CO_2 flux; Data n^o . = 39,412								
Intercept	520.71	2.63		198.05	0.00			
Factor1	296.44	2.63	0.49	112.75	0.00			
Factor2	61.69	2.63	0.10	23.46	0.00			
Factor3	-35.21	2.63	-0.06	-13.39	0.00			
Dependent variable = 222 Rn concentration; Data n ^o . = 35,178								
Intercept	1979.91	4.20		470.97	0.00			
Factor1	779.48	4.20	0.68	185.42	0.00			
Factor2	-230.39	4.20	-0.20	-54.80	0.00			
Factor3	94.85	4.20	0.08	22.56	0.00			





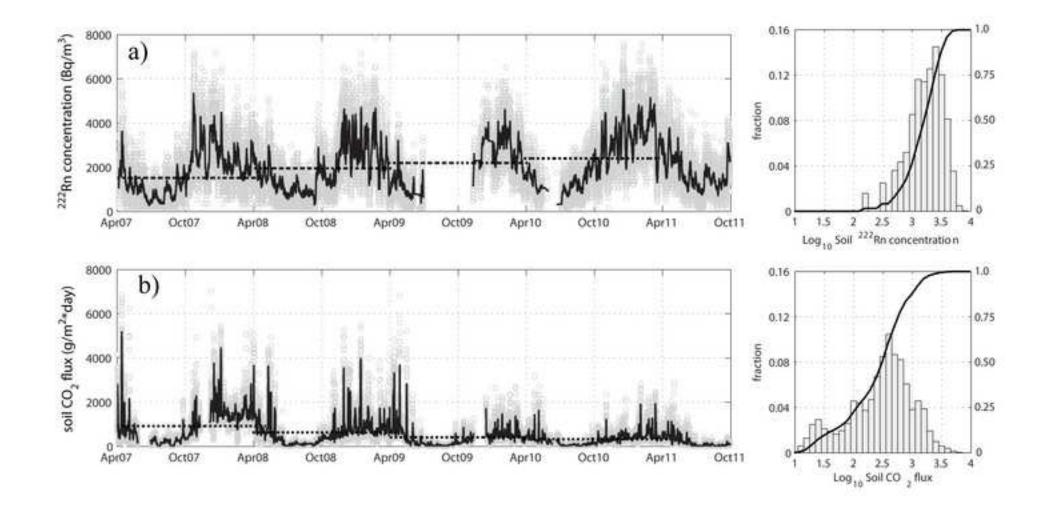
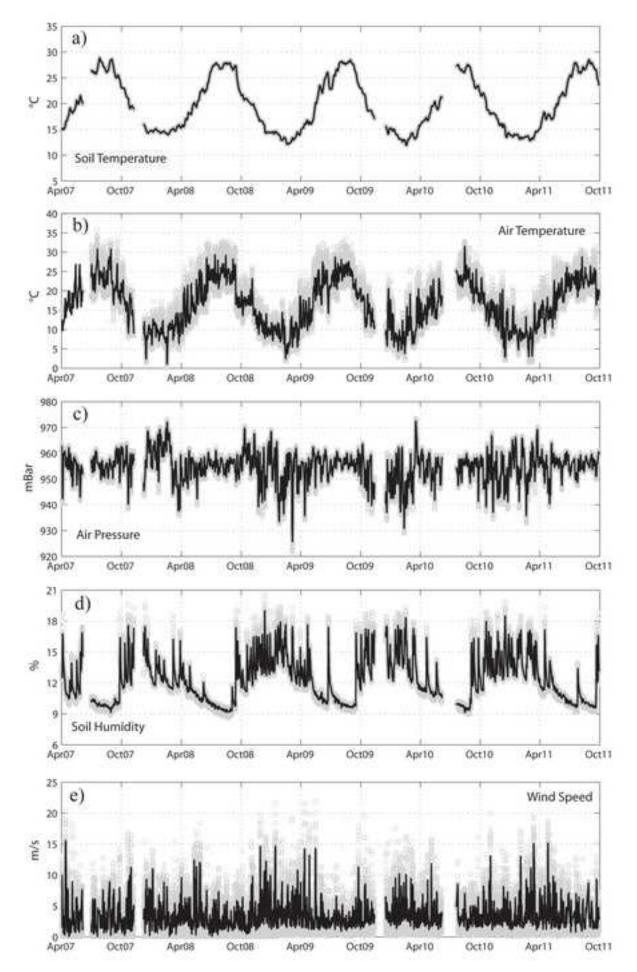
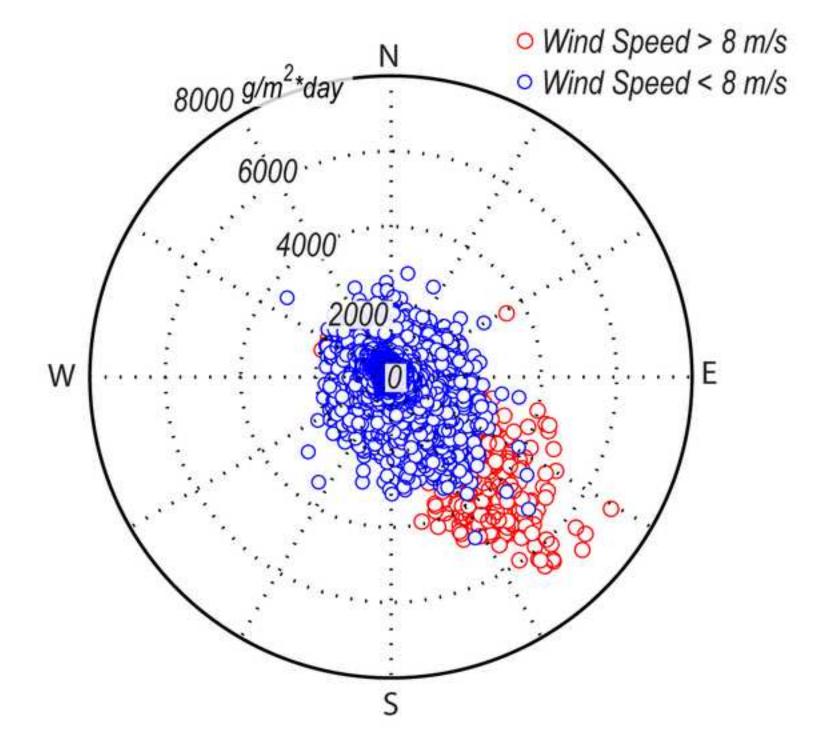
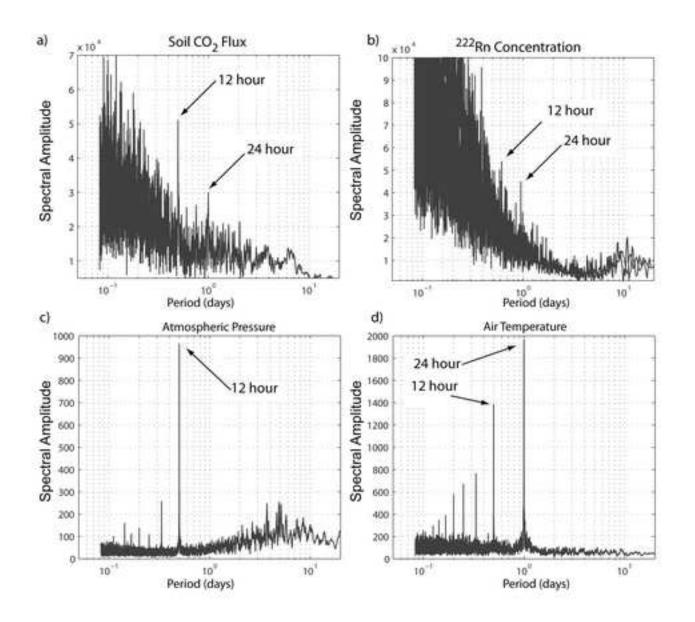
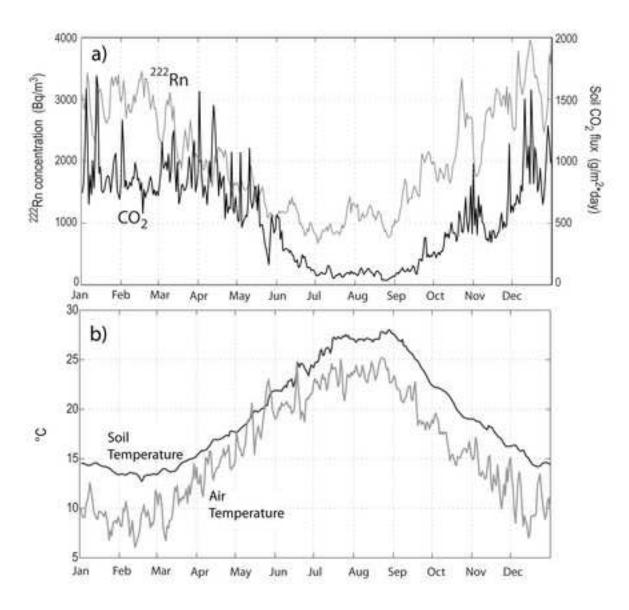


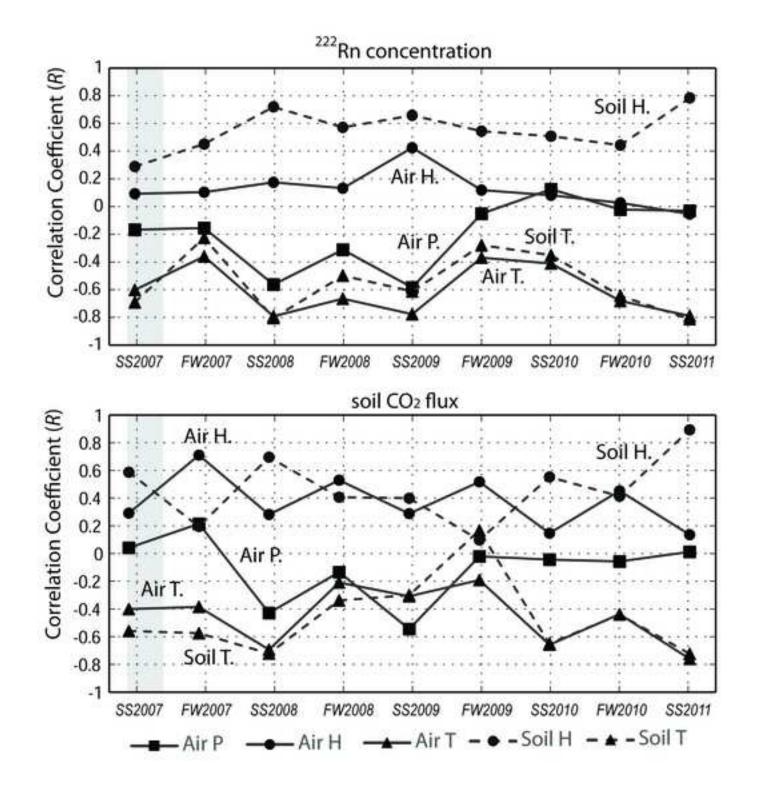
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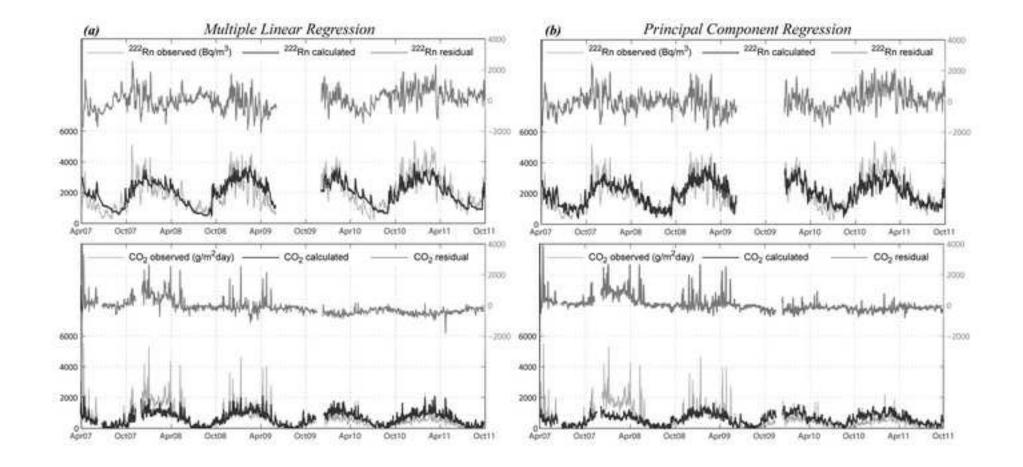


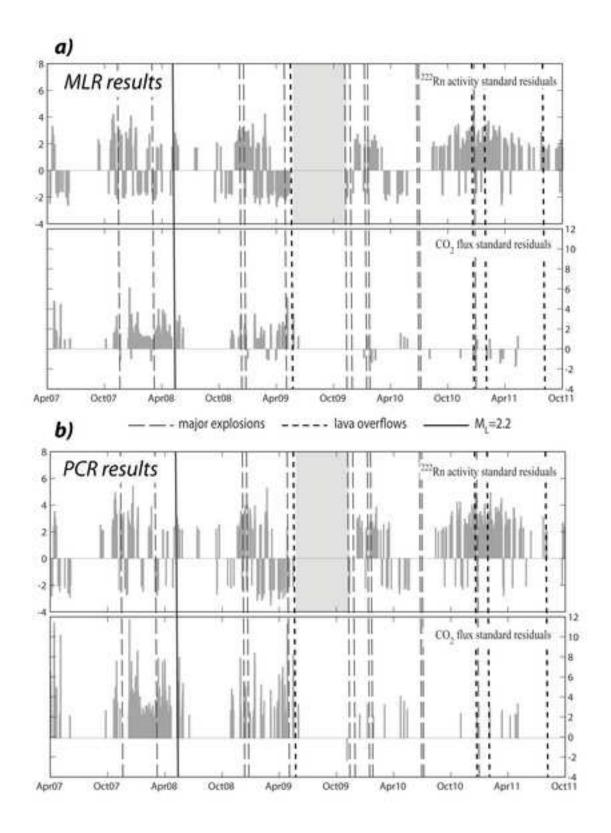












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