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A fluid inclusion study of blueschist-facies lithologies from the Indus suture zone, Ladakh (India): Implications for the exhumation of the subduction related Sapi-Shergol ophiolitic mélange

By

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Abstract

The best occurrence of blueschist-facies lithologies in Himalaya is that of the Shergol Ophiolitic Mélange along the Indus suture zone in Ladakh region of north-western India. These lithologies are characterized by well preserved lawsonite-glaucophane-garnet-quartz assemblages. This paper presents for the first time the results of a detailed fluid inclusion study on these lithologies, in order to understand the fluid P-T evolution and its tectonic implications.

The blueschist rocks from Shergol Ophiolitic Mélange record metamorphic peak conditions at ~ 19 kbar, 470°C. Several types of fluid inclusions are trapped in quartz and garnet, most of them being two-phase at room temperature. Three types of fluid inclusions have been recognised, basing on microtextures and fluid composition: Type-I are primary two-phase carbonic-aqueous fluid inclusions ($V_{CO2} – L_{H2O}$); Type-II are two-phase ($L_{H2O} – V_{H2O}$) aqueous fluid inclusions, either primary (Type–IIa) or secondary (Type–IIb); Type–III are re-equilibrated fluid inclusions. In the Type–I primary carbonic-aqueous inclusions, H_2O is strongly predominant with respect to CO_2; the homogenization temperature of CO_2 range from -7 to -2°C. The clathrate melting temperature in such inclusions varies in between +7.1 to +8.6°C. Type-II two-phase aqueous fluid inclusions show a wide range of salinity, from 7.8–14 wt.% $NaCl_{eq}$ (Type–IIa) to 1.65–6.37 wt.% $NaCl_{eq}$ (Type–IIb) with accuracy ± 0.4wt.% $NaCl_{eq}$.

Type-I and Type-IIa primary fluid inclusions are hosted in peak minerals (garnet and quartz included in garnet), therefore they were likely entrapped at, or near to, peak P-T conditions. The dominantly aqueous fluid of both Type-I and Type-IIa inclusions was most likely produced through metamorphic devolatilization reactions occurring in the subducting slab. Despite their primary nature, the isochores of Type-I and Type-IIa inclusions do not intersect the peak metamorphic conditions of the blueschist mineral assemblage, suggesting that these inclusions stretched or re-equilibrated during nearly isothermal decompression from 19 kbar to 3 kbar or

41 less, at T=290°C. This conclusion is further supported by their large variability in shapes and
42 sizes which range from irregular inclusions ('C'/arc shaped, hook shape and satellite type). This
43 decompression stage was followed by nearly isobaric cooling, testified by the occurrence of
44 dendritic networks of decrepitated and 'imploded' fluid inclusions.
45 Key Words: Fluid inclusions, Blueschist, lawsonite, Ladakh, Himalaya
46

47 **1. Introduction**

48 Fluid inclusions occurring in high pressure – low temperature (HP-LT) metamorphic rocks
49 reveal the nature of fluid as well as the role of fluid in the subduction history of continental or
50 oceanic crust. Fluid inclusions can also be used as a marker point in constraining parts of the
51 *P–T* paths of different metamorphic rocks (e.g. Hollister et al., 1979; Santosh, 1987; Hames et
52 al., 1989; Sisson et al., 1989; Vry and Brown, 1991; Winslow et al., 1994). Additionally,
53 evidences of fluid–rock interactions occurred during blueschist- and eclogite- facies
54 metamorphism can provide significant insights for our understanding of fluid storage and
55 recycling during subduction or continent-continent collision processes (Brunnsman et al., 2000;
56 Scambelluri and Phillipott, 2001).
57 The post-entrapment modification of the size and shape of fluid inclusions in response to
58 deformation, cooling or decompression can help to constrain the P-T paths of their host rocks
59 and hence provide some insights on their exhumation processes (e.g. Scambelluri, 1992;
60 Winslow et al., 1994, Vallis and Scambelluri, 1996; Kuster and Stockhert, 1997; El-Shazly and
61 Sisson, 1999; Bakker and Mamtani, 2000). The detailed observation of the fluid inclusion
62 textures and microthermometric results hold the key to fruitful data elucidation (e.g. Bodnar et
63 al., 1989; Sterner and Bodnar, 1989; Vityk et al., 1994; 1995; Barker, 1995; Sterner et al., 1995;
64 Vityk and Bodnar, 1995; Van den Kerkhof and Hein, 2001).

65 Fluid inclusion studies on subduction-related metamorphic rocks often revealed that the fluid
66 was retained within the subducting slab and that the fluid flow was limited (e.g. El-Shazly and
67 Sisson, 1999; Xiao et al., 2000; Gao and Klemnd, 2001). However, previous fluid inclusion
68 studies focused on HP-LT metamorphic terranes show no general agreement on fluid
69 composition and mechanisms of fluid flow.
70 Blueschists occur in many orogenic belts including the Pacific, Alpine-Himalayan and
71 Precambrian-Phanerozoic belts (Liou et al., 1990). Most of the studies on the blueschist-facies
72 rocks are focused on petrology, mineralogy and geochronology (e.g. Martin et al., 2011; Vitale
73 Brovarone et al., 2011; Chantel et al., 2012; Abers et al., 2013; Cao et al., 2013; Kim et al.,
74 2013; Spandler and Pirard, 2013), whereas the application of fluid inclusion studies to
75 blueschist-facies rocks is limited so far to few examples. Luckscheiter and Morteani (1980)
76 noticed the presence of CO_2-H_2O fluids in glaucophane-bearing rocks from the Tauern
77 Window, while Barr (1990) reported the presence of only aqueous fluid in blueschist rocks
78 from Syros. Invernizzi et al. (1996) studied the fluid inclusions in quartz veins hosted within
79 blueschist-facies rocks, and inferred the re-equilibration history of these fluid inclusions in
80 terms of uplift processes. The re-equilibration of the vast majority of primary fluid inclusions
81 in high *P*-low *T* rocks during exhumation is the main cause for the scarcity of such studies
82 (Touret, 1992). However, Vityk et al. (1994) and Vityk and Bodnar (1995) experimental works
83 on synthetic fluid inclusions presented textural benchmarks for recreating sections of *P*–*T* paths
84 and for enhancing tectonic interpretations.
85 Although the Himalaya is the archetype of collisional orogens, high-pressure metamorphic
86 rocks are rare (e.g. Lombardo and Rolfo, 2000; Guillot et al., 2008). The blueschist-facies rocks
87 in Himalaya are mostly lawsonite-bearing and occur along the Indus Suture Zone in Pakistan
88 (Shangla: Shams, 1972; Frank et al., 1977) and NW India (Sapi–Shergol, Ladakh: Honegger et
89 al., 1989; Groppo et al., 2016; Zildat: Virdi et al., 1977). Few lawsonite- and epidote blueschist-

facies rocks are also reported from the eastern portion of the orogen in the Indo-Burmese Ranges (Nagaland Ophiolite Complex: Ghose and Singh, 1980; Chatterjee and Ghose, 2010; Ao and Bhowmik, 2014). The preservation of lawsonite in the blueschists indicates very low geothermal gradients during subduction as well as during exhumation processes. The best occurrence of such lithologies in the Himalaya is that of Sapi-Shergol: a preliminary petrologic study of these blueschists by Honegger et al. (1989) constrained peak P-T conditions at 9-11 kbar, 350-450°C (conventional thermobarometry). Groppo et al. (2016) estimated significantly higher pressure (19 kbar, 470°C) for the same unit by using *P-T* pseudosection modelling and suggested a very low geothermal gradient (i.e. 8-9°C/km). Sachan and Mukherjee (2001) preliminarily studied the fluids from quartz veins hosted in the blueschist rocks from Sapi-Shergol and provided comments upon their exhumation history. However, these quartz veins are related to a very late stage of the metamorphic evolution, and consequently do not give any direct clue on the blueschist formation and exhumation.

As concerning other subduction-related rocks from the Himalaya, a fluid inclusion study by Mukherjee and Sachan (2009) focusing on the ultra-high pressure eclogites from the Tso Morari, showed the prevalence of high-salinity brine, N_2, CH_4, CO_2 –bearing fluids and low-salinity aqueous fluids. This study shows the complex nature of the fluid involved during deep subduction and exhumation of Indian continental crust. Additionally, Ferrando et al. (2007) reported the presence of pure CO_2 and CO_2-H_2O fluids in granulitized eclogites from eastern Himalaya.

This paper is the first systematic fluid inclusion study on the blueschist rocks from Sapi-Shergol ophiolitic mélange (Ladakh Himalaya, NW India). The fluids associated to peak and retrograde metamorphic stages are characterised, and the source of these fluids is assessed. Moreover, the fluid evolution is discussed in relation to the P–T paths inferred by Groppo et al. (2016) on the base of thermodynamic modelling. The results from this study have important implications for

115 the understanding of the nature and role of fluids during subduction-related metamorphism and
116 the consequent uplift, in the framework of the geodynamic evolution of the Himalayan orogeny.
117

118 ## 2. Geological Setting

119 The blueschists studied in this paper occur along the Indus suture zone in Ladakh, NW India
120 (Fig. 1a). From SE to NW, blueschist-facies lithologies are reported at Zildat, Urtsi, Hinju, and
121 Sapi-Shergol (Honegger et al., 1989). Among these occurrences, the largest one is that of Sapi-
122 Shergol in western Ladakh, south of Kargil. Tectonically, the Sapi–Shergol blueschists belong
123 to a narrow belt and are part of an ophiolitic mélange (Honegger et al., 1989) (Fig. 1b).
124 This ophiolitic mélange is thrusted over the Kargil Molasse to the west as well as to the north,
125 and it is in contact with the Nindam and Lamayuru Formations to the north and to the south,
126 respectively. To the south, the Triassic limestone of the Zanskar Unit is thrusted over the
127 mélange (Frank et al., 1977, Honegger et al., 1982). The ophiolitic mélange is inferred to be a
128 remnant of a paleo-accretionary prism produced by the northward subduction of the Neo-
129 Tethys, originally separating the Ladakh arc to the south from the Asian active margin to the
130 north (Mahéo et al., 2006) (Fig.1a). This paleo-accretionary prism comprises sedimentary
131 bodies including mainly basic lithologies that have been metamorphosed under low-grade to
132 lawsonite blueschist-facies conditions (Frank et al., 1977; Honegger et al., 1989; Reuber et al.,
133 1987; Ahmad et al., 1996; Rolfo et al., 1997; Robertson, 2000; Mahéo et al., 2006). The Sapi–
134 Shergol ophiolitic mélange therefore shows a complex internal architecture, including a number
135 of slivers of the paleo-accretionary prism, intercalated with numerous slivers of other units
136 including the Nindam and Lamayuru Formations, as well as low-grade meta-ophiolitic rocks
137 with serpentinized peridotites intruded by basic dikes (“sheared serpentinites” of Robertson,
138 2000).

The protoliths of the blueschist lithologies cropping out around the Shergol village are mainly basic metavolcanics and volcanoclastites with subordinate amounts of cherts, metasediments and minor carbonatic lithologies. Mahéo et al. (2006) suggested that the blueschists originate from calc-alkaline igneous rocks formed in an intra-oceanic arc environment. Whole-rock and glaucophane K–Ar geochronology gave an age of ca. 100 Ma for the high-pressure metamorphic peak (Honegger et al., 1989). Petrographically, metabasic rocks are mainly fine-grained glaucophane -bearing schists, with lawsonite and minor clinopyroxene and phengite. Metavolcanoclastic rocks show a clastic structure with irregular fragments of metabasic rocks set in a very fine-grained matrix. The clasts generally contain a blue amphibole + lawsonite ± clinopyroxene assemblage. The matrix is very fine-grained and mainly comprises blue amphibole, chlorite and minor aegirine.

Silicic and impure carbonatic metasediments are locally found as intercalations within the metabasic and metavolcanoclastic rocks. The silicic metasediments comprise glaucophane + lawsonite + phengite ± garnet schists, lawsonite + glaucophane + phengite + garnet quartzitic micaschists and glaucophane + garnet + phengite quartzites. Lawsonite and garnet are either fine-grained or porphyroblastic. Lawsonite and garnet porphyroblasts are from few centimeters to few millimeters in size, and generally overgrow the main foliation. The impure carbonatic metasediments are very fine-grained and mainly consist of lawsonite, calcite, glaucophane, and minor phengite ± prehnite.

3. Petrography

A detailed petrographic study of two representative samples of metasediments (samples 14-4B and 14-6F/G) has been carried out. These metapelites are mainly composed of lawsonite, glaucophane, garnet, phengite and quartz, with minor calcite concentrated in late veins.

163 Accessory minerals are titanite and pyrite. Both samples are very fresh, one of them being
164 relatively fine-grained (sample 14-4B) compared to the other (sample 14-16F/G).
165 Sample 14-4B is characterised by alternating layers rich in quartz or lawsonite + glaucophane
166 + phengite, respectively (Fig. 2a, b). Glaucophane occurs in both layers as fine-grained
167 idioblasts; it shows characteristic zoning pattern with light blue core and a dark blue rim. The
168 preferred orientation of glaucophane and phengite defines the main foliation. Garnet is weakly
169 zoned and occurs as small idioblasts (Fig. 2b). Lawsonite occurs as fine-grained idioblasts (Fig.
170 2c); it contains inclusions of quartz and titanite. In the quartz-rich layers, lawsonite locally
171 shows a skeletal habit, being intergrown with quartz.
172 The main assemblage of sample 14-6F/G is lawsonite + glaucophane + phengite + garnet +
173 quartz, with glaucophane occurring as a major constituent (Fig. 2d). Titanite occurs as accessory
174 mineral. The main foliation, defined by the preferred orientation of glaucophane and minor
175 phengite, is overgrown by lawsonite and garnet porphyroblasts and is locally crenulated (Fig.
176 2d, e, f). The growth of lawsonite and garnet porphyroblasts pre-dated the crenulation event.
177 Garnet occurs as medium- to coarse-grained idioblasts (Fig. 2e) locally included within the
178 lawsonite porphyroblasts, and shows strong zoning. Variably sized inclusions of glaucophane,
179 actinolite, phengite, quartz, chlorite and rare omphacite are found in garnet porphyroblasts.
180 Glaucophane occurs as fine-grained nematoblasts in the matrix. Porphyroblastic lawsonite is
181 up to 1 centimetre in size and sub-idioblastic, and overgrows the main foliation. Phengite is
182 fine- to medium-grained and is found in association with glaucophane (Fig. 2d). The main
183 foliation is locally crosscut by quartz ± albite ± chlorite veins (Fig. 2f). It is worth noting that
184 in both samples, garnet was predicted to have grown at (or close to) peak P-T conditions, i.e. in
185 the narrow P-T range of 400-470°C, 17-19 kbar (Groppo et al., 2016).
186
187 ## 4. **Methods**

188 Double polished wafers with a thickness of about 0.4 ± 0.2 mm were prepared from the samples
189 described in chapter 3. A microthermometric study was conducted on a Linkam THMSG 600
190 stage mounted on Olympus microscope at the Wadia Institute Fluid inclusion Laboratory. The
191 equations of Zhang and Frantz (1987) and Brown and Lamb (1989) were used for the estimation
192 of isochores of aqueous fluid inclusions, whereas the equation of Bower and Helgeson (1985)
193 was used for the calculation of isochores of CO_2 –H_2O inclusions in the "Flincor" computer
194 program of Brown (1989).
195
196 **5. Fluid inclusion petrography**
197 We have studied fluid inclusions hosted in garnet and in matrix quartz, as well as in quartz
198 included in garnet porphyroblasts belonging to the peak metamorphic mineral assemblage (i.e.
199 garnet + lawsonite + glaucophane + phengite + quartz ± omphacite; see Groppo et al., 2016).
200 Texturally, two types of fluid inclusions have been found: regular shaped inclusions (Type-I
201 and Type-II) and re-equilibrated inclusions (Type-III). Regular shaped inclusions are either
202 two-phase carbonic-aqueous (Type-I) or two-phase aqueous (Type-II) inclusions. Type-II
203 inclusions are more abundant than Type-I inclusions.
204 Type–I inclusions are two-phase inclusions (V_{CO2} + L_{H2O}) at room temperature (~ 27°C) and
205 contain an aqueous-carbonic fluid (Fig. 3A, B). The size of these inclusions are small (average
206 of 9 μm). The CO_2 vapour has a variable volume proportion ranging from 10 to 30% of the total
207 volume of the inclusion. These inclusions are systematically found as isolated occurrences
208 within garnet and quartz, and are dispersed randomly in the host grains (Fig. 3A, B), therefore
209 they have been interpreted as primary in origin; moreover, because garnet is a peak phase, Type-
210 I inclusions have been entrapped at peak P-T conditions.
211 Type-II inclusions are two-phase aqueous (V_{H2O} + L_{H2O}) inclusions and are further classified as
212 Type-IIa and Type-IIb on the basis of their mode of occurrence. The size of these types of

inclusion ranges between 9 and 15 μm. Type-IIa inclusions are isolated and randomly distributed in the matrix quartz as well as in quartz included in garnet grains, and are therefore interpreted as primary inclusions; they were likely entrapped at, or near to, peak P-T conditions because quartz included in garnet is a prograde-to-peak phase. Type-IIb are concentrated along microfractures of the host minerals and are interpreted as secondary inclusions. The trails of the secondary inclusions often cross-cut the grain margins (i.e. transgranular inclusions) (Fig. 3C, D).

Type-III inclusions are re-equilibrated inclusions showing partially dendridic type network as well as necking phenomena (Fig. 4A-D). The dendritic type network may be originated by the dissolution of inclusion walls, as well as by the closing of inclusion voids which lead to form numerous tails and channels filled by fluids (Invernizzi et al., 1998). Some inclusions show necking phenomena, and few also show "C" type microstructures. Some solitary inclusions surrounded by minute satellite inclusions are observed in matrix quartz.

6. Microthermometric results

Microthermometry reveals compositional variations of the fluid inclusions from aqueous-carbonic to aqueous. The results of microthermometric measurements are shown in histograms (Figs.5, 6 and 7) and summarized in Table 1.The Type-I aqueous-carbonic fluid inclusions hosted in the matrix quartz as well as in quartz included in garnet and in garnet itself, show initial melting temperature (Tim) between -56.9°C and -56.6°C, thus suggesting a nearly pure CO_2 composition for the vapour phase. The homogenization temperature of CO_2 in these inclusions lies between -7 to -2°C (Fig. 5B). The clathrate melting temperature was also observed in some carbonic-aqueous inclusions, which occurred between +7.1 to 8.6°C. The total homogenization of such inclusions took place in between 227° and 265°C (Fig. 5A).

The Type-II two-phase aqueous inclusions mainly occur in matrix quartz as well as in quartz included in garnet. Type-IIa primary inclusions and Type-IIb secondary inclusions show different initial melting temperature (Tim) and final melting temperature (Tfm) values. Type IIa primary H_2O-NaCl inclusions show Tim in the range −21.2°C to −22.6°C, Tfm of −5°C to −10°C (Fig. 6A) and temperature of homogenization (Th) between 250°C and 300°C (Fig. 7B). The corresponding salinity ranges are estimated at 7.8–13.98 wt% NaCl equivalent, with density of 0.863–0.874 g/cm^3. Type-IIb secondary two-phase aqueous inclusions occurring in transgranular trails show a wider range of Tim (−21° to −23 °C), and significantly different Tfm (−1° to -4°C; Fig. 6 B) and Th (145° to 220°C; Fig. 7A) values. The corresponding salinity ranges are estimated at 1.65–6.37 wt% NaCl equivalent, and density is in the range of 0.889–0.938 g/cm^3.

Composition and density of the fluid phase observed in the different generations of inclusions are used to calculate the isochores in the P–T space, and the implications are discussed in a later section. The minimum Th and the maximum Th obtained from the peak value in Figs. 5,6,7 for each inclusion type were considered for construction of isochores, following the criteria given by Touret (2001),.

7. Discussion

The study of fluid inclusions provides important constraints on their trapping conditions (textural evidence in relation to *P–T* estimates), composition and behaviour of metamorphic fluids. This is the first systematic fluid inclusion study on the blueschists from the Sapi-Shergol ophiolitic mélange that correlates P-T data from mineral geothermobarometry and fluid inclusions. Our results reveal the pervasive presence of aqueous-carbonic and aqueous fluids preserved in quartz and garnet.

7.1. Interpretation of fluid inclusions textures

In the studied samples, the isolated nature of Type-I aqueous-carbonic inclusions and of Type-IIa two-phase aqueous inclusions hosted in quartz and garnet reveals that they may have formed during the crystallization of their host minerals, and may thus be considered to be trapped early during the mineral growth (Roedder, 1984). Moreover, because garnet is a peak phase, the primary Type-I and Type-IIa inclusions have been entrapped at, or near to, peak P-T conditions. Type-IIb two-phase aqueous inclusions occur in transgranular trails and are therefore clearly secondary in nature.

Concerning Type-III inclusions, the most commonly observed features are: (i) necking, (ii) inclusions with 'C' or Arc shaped microstructure, (iii) satellite inclusions, (iv) hook shaped or annular inclusions, (v) dendritic network of inclusions (Fig.4 A, B, C & D). The necking textures are interpreted as related to intense dissolution of the inclusion walls to produce highly irregular inclusion morphology. These microtextures were likely formed along a decompression path: fractures initially developed along the inclusion planes in which inclusions primarily necked down, and later on they were stretched (Fig 4C). The inclusions displaying a hook-like morphology or annular textures (Fig.4C) are similar to inclusions formed under isothermal decompression (ITD), as shown in the experimental work of Sterner and Bodnar (1989), Vityk and Bodnar (1995), and Bodnar (2003). Boullier et al. (1991) proposed that annular textures are symptomatic of an anisotropic stress environment. Satellite inclusions around a bigger solitary inclusion are analogous to 'explosion' textures described for natural and synthetic fluid inclusions in quartz that have experienced overpressures caused by 'near isothermal decompression' (e.g., Bodnar et al., 1989; Vityk et al., 1994; Vityk and Bodnar,1995). Opposite to these microtextures which are related to decompression, the development of dendritic networks of imploded inclusions in quartz is indicative of near isobaric cooling (IBC) in the last stages of metamorphic evolution. This type of texture is similar to that described by Vityk

et al. (1994) and Vityk and Bodnar (1995) for their experimental work during isobaric cooling (IBC).

7.2. Nature and Source of Fluids

Two types of fluid (aqueous-carbonic and aqueous) have been found in the blueschist mineral assemblage. Type-I aqueous-carbonic fluids occur as primary inclusions in peak metamorphic garnet as well as in quartz. Type-IIa aqueous fluid with high saline nature is preserved as primary inclusions in matrix quartz, as well as in quartz included in garnet, whereas Type-IIb aqueous fluid with low saline nature is observed as secondary inclusions in quartz.

Type-I aqueous-carbonic inclusions have constant vapour: liquid ratios (i.e. 30:70). The temperature of clathrate melting and the homogenization temperature of the CO_2 phase in these inclusions indicate that they are predominantly aqueous with a small amount of CO_2, consistently with the stability of lawsonite at peak metamorphic conditions. According to Groppo et al. (2016), in fact, the contemporaneous growth of lawsonite and garnet would have been enhanced by a protracted H_2O influx at the relatively high pressure of ca. 17–18 kbar. The predominance of aqueous fluids in high P metamorphic assemblages is also shown in previous studies which call upon substantial fluid release at high P–T conditions (e.g. Clarke et al., 2006; Tsujimori and Ernst, 2014; Ulmer and Trommsdorff, 1995; Scambelluri et al., 2004; Poli and Schmidt, 1995; Poli et al., 2009) through metamorphic devolatilization reactions occurring in the subducting slab (Bebout, 1991, 1995; Jarrard, 2003).

7.3. Reconstruction of the P-T-fluid history

P–T estimates based on conventional thermobarometry proposed peak conditions of 350–420 °C, 9–11 kbar for the Sapi-Shergol blueschists (Honegger et al., 1989). Petrologic modeling performed by our research team (Groppo et al., 2016) inferred peak P–T conditions significantly

312 higher than those previously estimated, i.e. ca. 470 °C, 19 kbar. According to Groppo et al.
313 (2016), the estimated peak metamorphic conditions suggest that the blueschists experienced a
314 cold subduction history along a very low to low thermal gradient ("early" prograde: ca. 5–6
315 °C/km; "late" prograde: ca. 7–8 °C/km). The preservation of lawsonite in the studied lithologies
316 further indicate that the exhumation process must have been coupled with substantial cooling
317 (i.e. without crossing the lawsonite-out boundary; Zack et al., 2004).
318 As concerning the fluid evolution, isochores are plotted in comparison with the *P–T* path
319 inferred on the basis of thermodynamic modelling and thermobarometry (Fig. 8). Isochores of
320 both Type-I aqueous–carbonic inclusions and Type-II aqueous inclusions plot in the lower part
321 of the P–T space. Although these inclusions are texturally primary and are hosted within peak
322 minerals (garnet and quartz included in garnet), they do not plot along the prograde *P–T* path
323 or at peak metamorphic conditions, thus suggesting that these inclusions must have experienced
324 considerable leakage and re-equilibration, resulting in only low-density fluid eventually
325 preserved in the inclusions. Therefore, it is assumed that Type-I CO_2-H_2O and Type-IIa aqueous
326 inclusions most likely formed at or near the peak *P–T* conditions at ~ 19 kbar, ~ 450°C, but
327 they were later re-equilibrated during exhumation. The isochores of secondary aqueous
328 inclusions (Type IIa) plots above the isochores of primary inclusions because the secondary
329 aqueous inclusions originate at lower temperature and their densities are thus higher (Fig.8).
330 Evidences of such a pervasive re-equilibration are very clearly represented by the textural
331 features of Type-III fluid inclusions described in the previous chapters (Fig.4). The observed
332 textural features are indicative either of re-equilibration occurred during isothermal
333 decompression, as well as during isobaric cooling at the final stages of exhumation. Similar
334 high degrees of re-equilibration of fluid inclusions during retrogression is well documented by
335 Sterner and Bodnar (1989) and Kuster and Stockhert (1997) for high-P and low-T rocks. We
336 therefore suggest that these re-equilibrated inclusions might have been trapped during the early

337 stages of subduction, while their re-equilibration took place during the subsequent fast
338 exhumation. Such high internal underpressure at low temperature would develop only when the
339 subduction / burial path is very steep and consequently follows very low geothermal gradients;
340 the preservation of lawsonite in the studied blueschist facies lithologies strongly support the
341 scenario in which very low geothermal gradients were prevailing during initial subduction as
342 well as during exhumation (Fig. 8). The presence of dendritic network type inclusions suggests
343 isobaric cooling (as evidenced by the experimental studies of Bodnar et al., 1989; Vityk et al.,
344 1994; Vityk and Bodnar, 1995). The hook like inclusions were likely trapped at about 3 kbar
345 (Vityk and Bodnar, 1995), whereas the implosion texture formed at higher pressure (Boulier,
346 1999). In our case, the last stage of the P-T evolution is represented by the isobaric cooling after
347 the isothermal decompression, therefore we suggest that the dendritic morphology might have
348 been formed at this last stage (Fig. 8). This interpretation is clearly compatible with the broad
349 geological setting of the area.
350

351 **8. Conclusions**

352 The first detailed fluid inclusions study of Himalayan blueschist-facies lithologies from the
353 Sapi-Shergol ophiolitic mélange in the Indus suture zone, revealed predominantly CO_2–H_2O
354 and H_2O-NaCl inclusions. The primary CO_2-H_2O and H_2O-NaCl inclusions were trapped by
355 host minerals at peak P-T conditions. The fluid was most likely produced through metamorphic
356 devolatilization reactions occurring in the subducting slab. Additional multiple generations of
357 re-equilibrated inclusions show characteristic morphologies developed in response to fast rates
358 of uplift. A nearly isothermal decompression path followed peak metamorphic P-T condition
359 and is characterized by stretching, necking and formation of "C" or hook-like shaped inclusions.
360 A veining stage is associated to late isobaric cooling and is characterized by the local
361 development of a dendritic network of fluid inclusions.

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565 **Figure Captions**

566 **Figure 1.** (A) Regional Geological Map of Ladakh Himalaya (adopted from Thakur and
567 Misra, 1984, and Mahéo et al., 2006); (B) Geological map of Shergol ophiolitic mélange
568 (after Honegger et al., 1989).

569

570 **Figure 2.** Representative microstructures of the studied blueschists. **Sample 14-4B: (a)**
571 Glaucophane + phengite + lawsonite layers alternated to discontinuous quartz-rich layers. Plane
572 Polarized Light (PPL); **(b)** Glaucophane + phengite + lawsonite layer containing variably sized
573 garnet Crosse Polarized Light (XPL); **(c)** Fine-grained idioblasts of lawsonite occur in
574 association with glaucophane (PPL); **Sample 14-6F/G: (d, e)** Glaucophane with minor
575 phengite defines the main foliation. The foliation is overgrown by lawsonite and garnet
576 porphyroblasts (PPL); **(f)** Lawsonite porphyroblast is crosscut by thin quartz veins (XPL).

577

578 **Figure 3.** Representative photomicrographs of Type-I and Type-II fluid inclusions (PPL). **(a)**
579 Type-I primary two-phase carbonic-aqueous inclusions in garnet (sample14-6F/G). **(b)** Type-I
580 primary two-phase aqueous-carbonic inclusions in matrix quartz (sample 14-4B). **(c)** Type-IIa
581 primary two-phase aqueous inclusions in quartz included in garnet (sample14-6F/G). **(d)** Type-
582 IIb secondary two-phase aqueous inclusions and Type-III re-equilibrated inclusions in quartz
583 (sample 14-4B).

584

585 **Figure 4.** Representative photomicrographs of Type-III fluid inclusions (PPL). **(a)** Re-
586 equilibrated inclusions in quartz showing implosion textures (sample14-6F/G). **(b)** Re-
587 equilibrated inclusions in quartz exhibiting necking and dendritic type networks of fluid
588 inclusions (sample14-6F/G). **(c)** Re-equilibrated inclusion in quartz showing different
589 morphologies: (i) hook-shape microstructure; (ii) “C” or “arc” shape, and (iii) necking and

590 stretched phenomena (sample 14-4B). **(d)** Dendritic type network of secondary fluid inclusions
591 in quartz (sample 14-4B).
592
593 **Figure 5.** Histograms showing **(a)** homogenization temperature (Th), and **(b)** final melting
594 temperature of CO_2 for Type-I CO_2-H_2O primary inclusions hosted in quartz and garnet.
595
596 **Figure 6.** Histograms showing final melting temperature of **(a)** Type-IIa primary H_2O-NaCl
597 inclusions in quartz, and **(b)** Type-IIb secondary H_2O-NaCl inclusions in quartz.
598
599 **Figure 7.** Histograms showing homogenization temperature (Th) for **(a)** Type-IIa primary H_2O-
600 NaCl inclusions in quartz, and **(b)** Type-IIb secondary H_2O-NaCl inclusions in quartz.
601
602 **Figure 8.** Comparison between fluid inclusion isochores and the *P–T* path (thick arrow in pink)
603 recorded by blueschist from the Sapi-Shergol ophiolitic mélange (after Groppo et al., 2016).
604 The lawsonite boundary is drawn after Zack et al. (2004). The yellow box shows peak
605 metamorphic conditions. The retrograde path is mainly characterized by nearly isothermal
606 decompression (as evident from hook shaped, annular inclusions and explosion type textures).
607 The last portion of the P-T path is characterized by isobaric cooling (as constrained by the
608 development of dendritic type networks of inclusions).
609

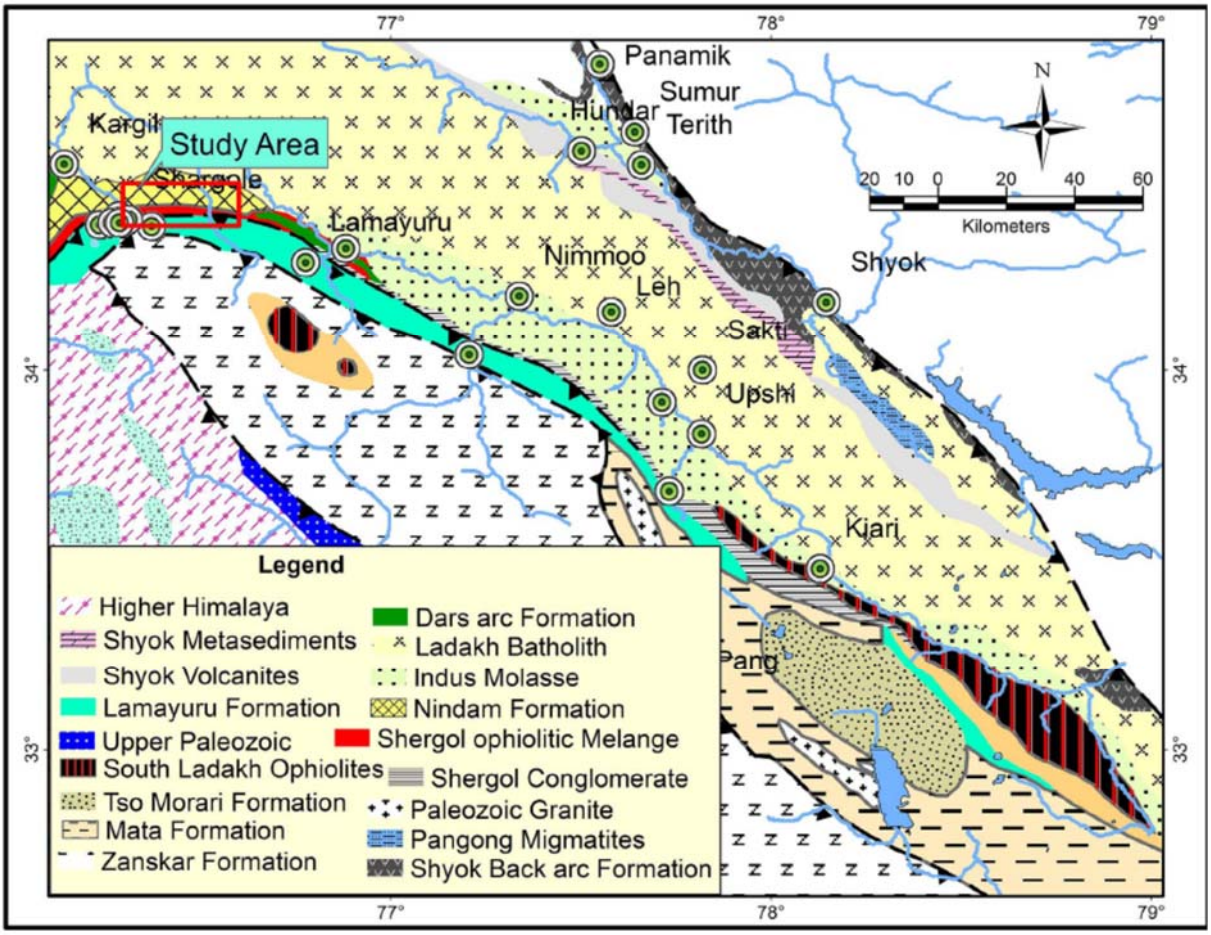
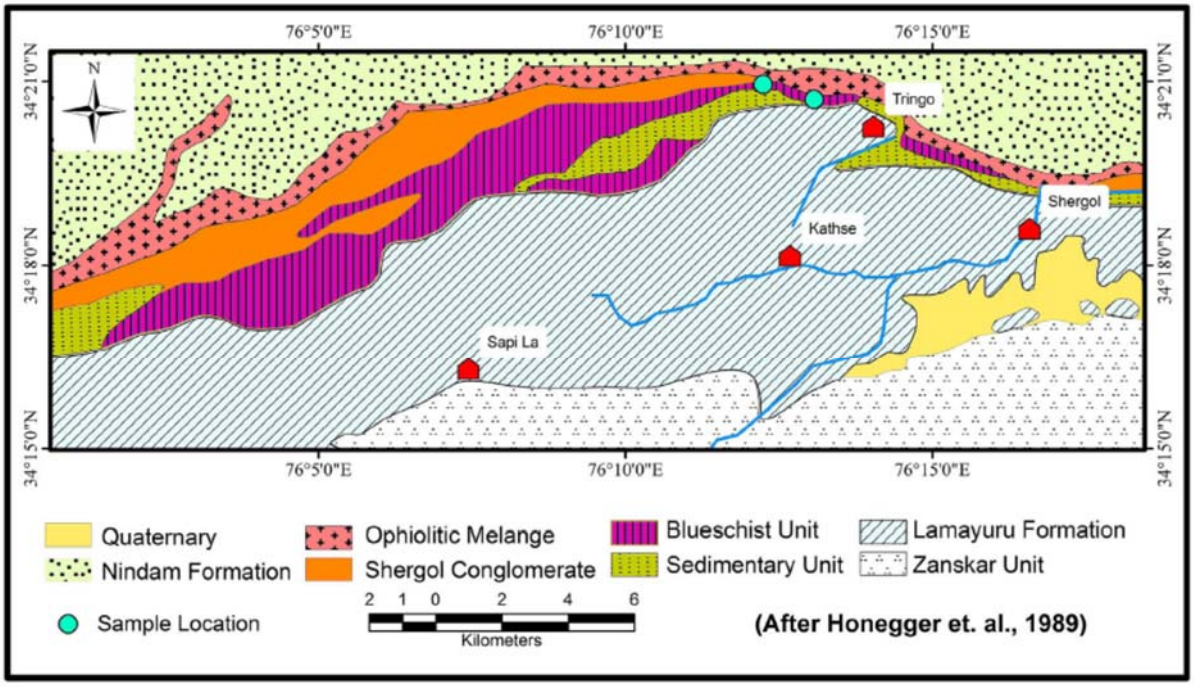


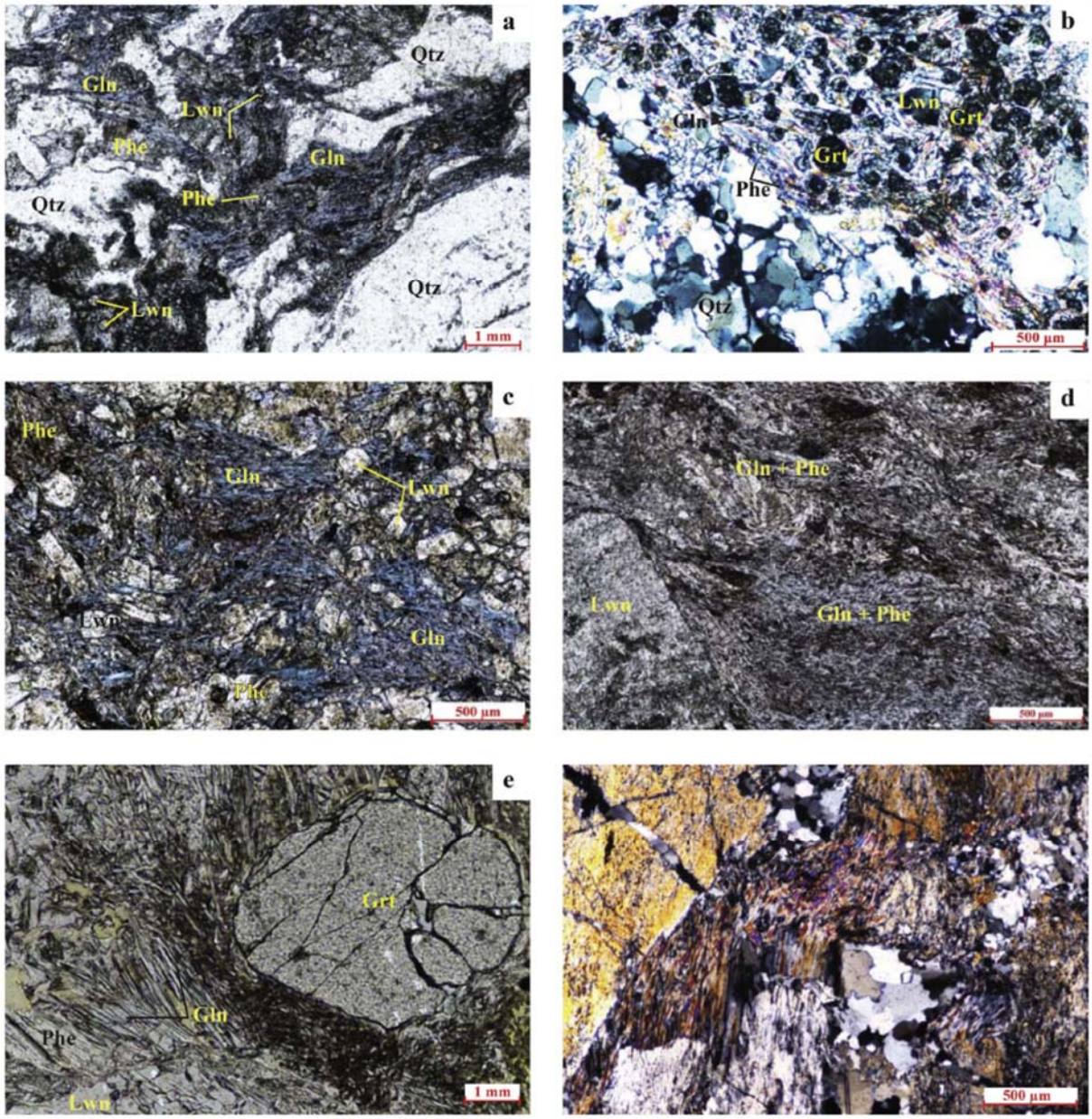
Fig. 1a



610

611 Fig. 1

612


a
Qtz
Gln
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1 mm
b
Gln
Lwn
Grt
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Phe
Qtz
500 µm
c
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Gln
Lwn
Lwn
Gln
Phe
500 µm
d
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e
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Gln
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1 mm
500 µm

Fig. 2

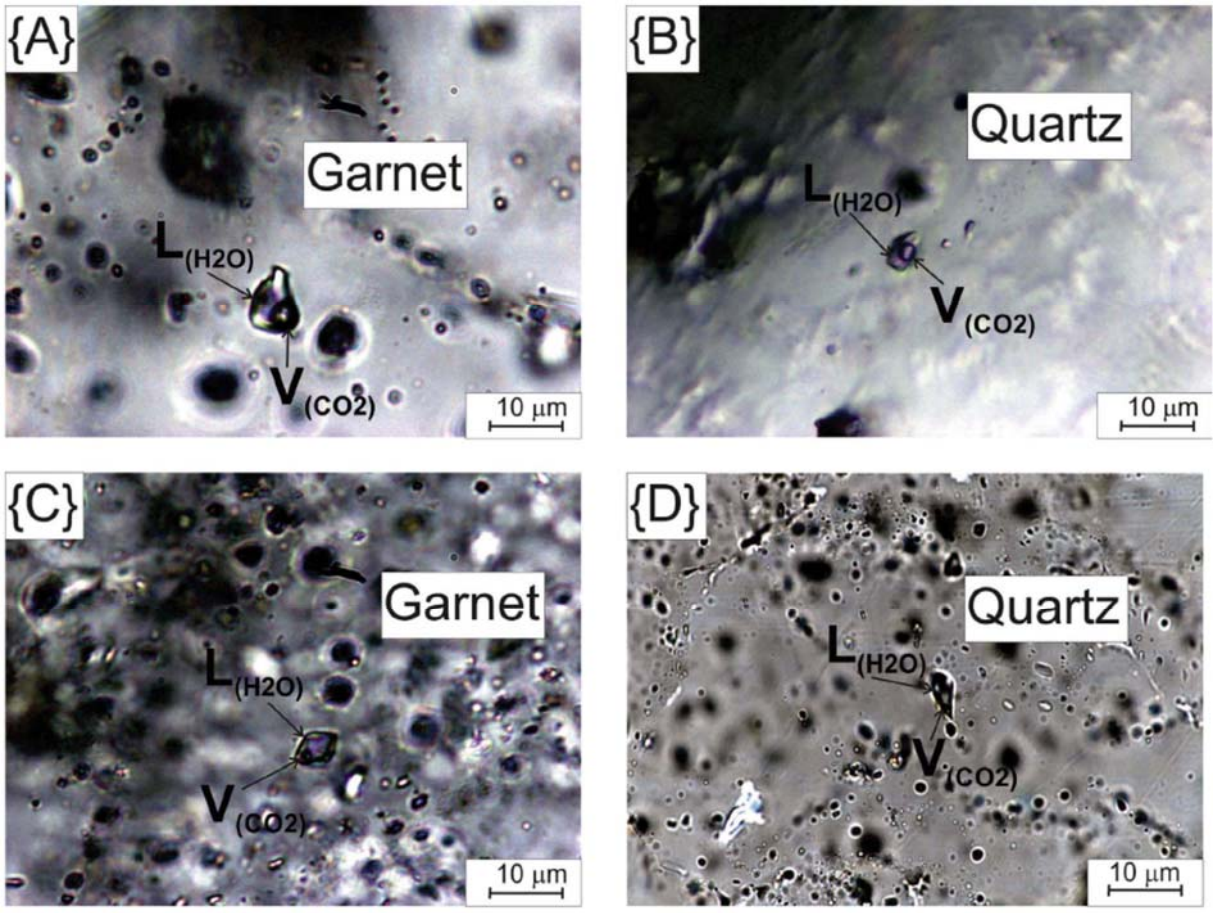

{A}
Garnet
L(H2O)
V(CO2)
10 µm
{B}
Quartz
L(H2O)
V(CO2)
10 µm
{C}
Garnet
L(H2O)
V(CO2)
10 µm
{D}
Quartz
L(H2O)
V(CO2)
10 µm

Fig. 3

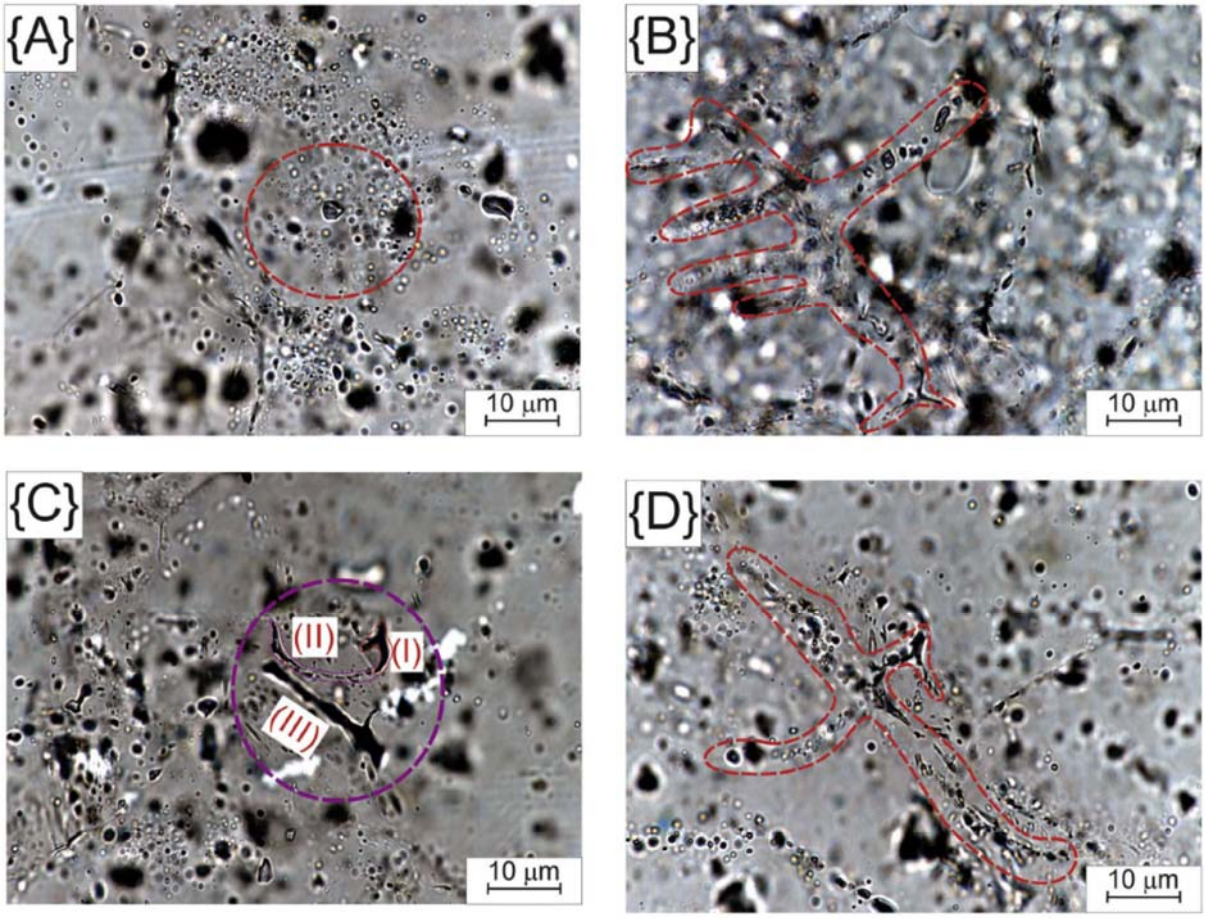
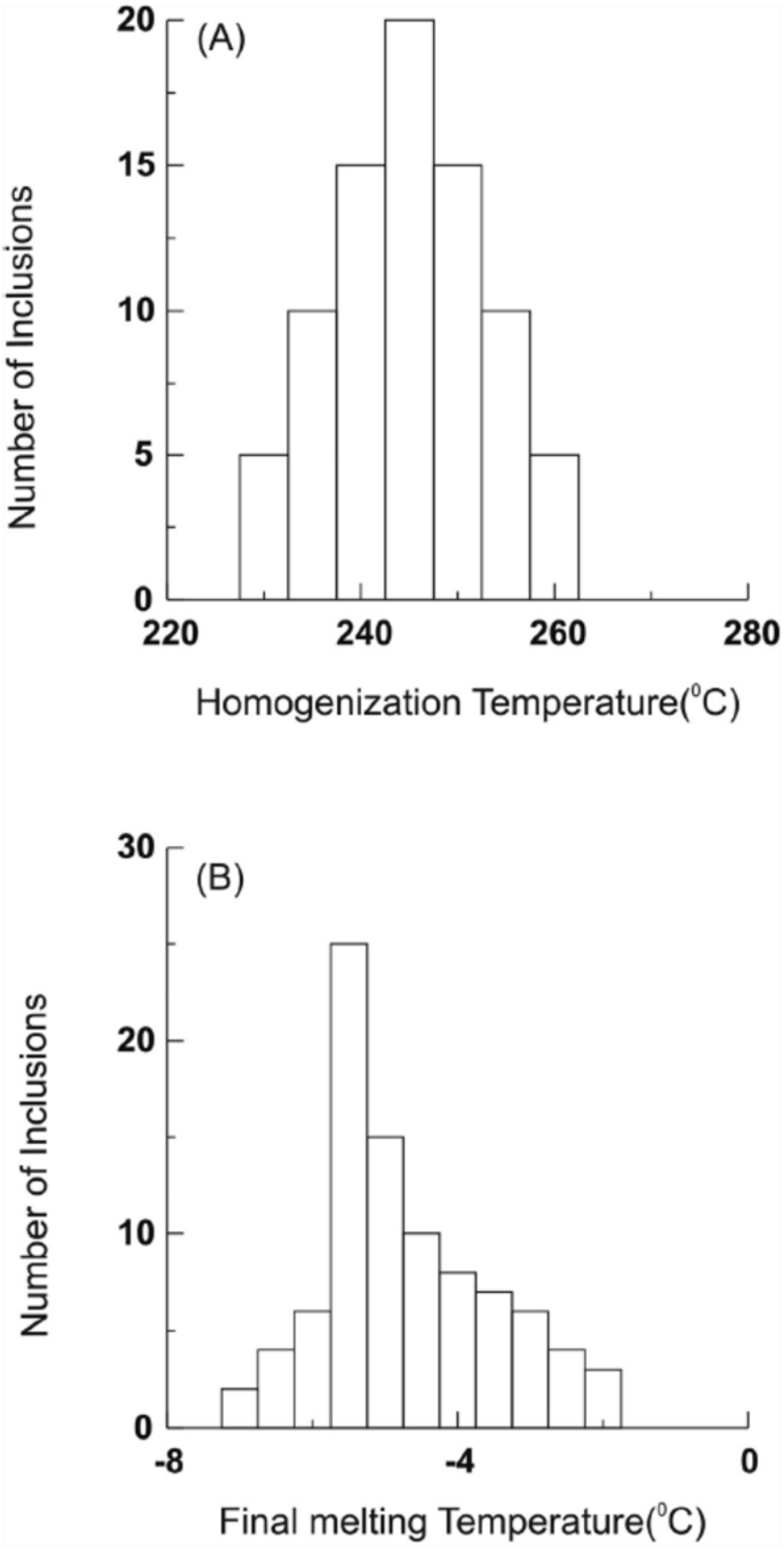


Fig. 4



622

623 Fig. 5

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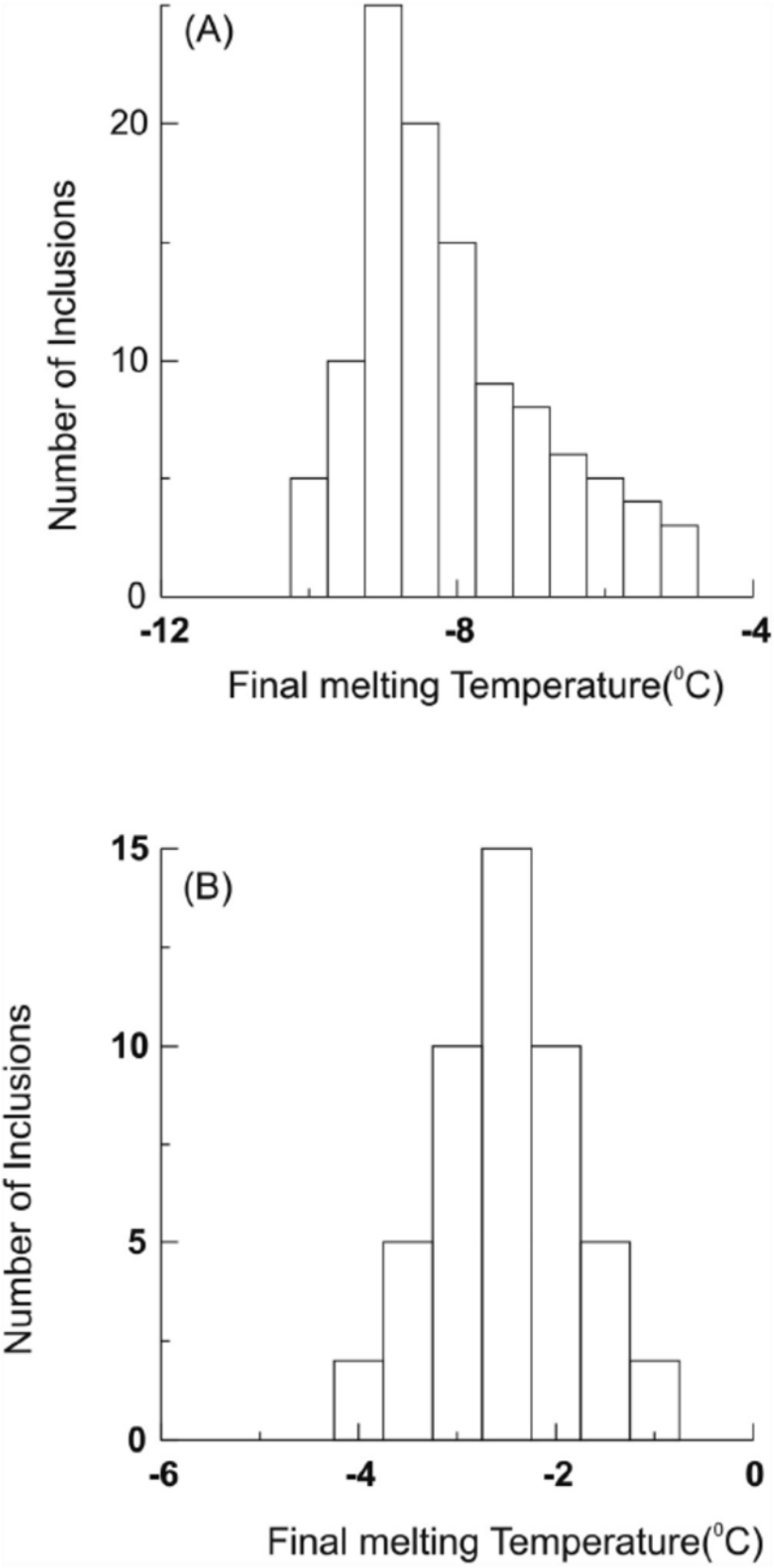

625

626 Fig. 6

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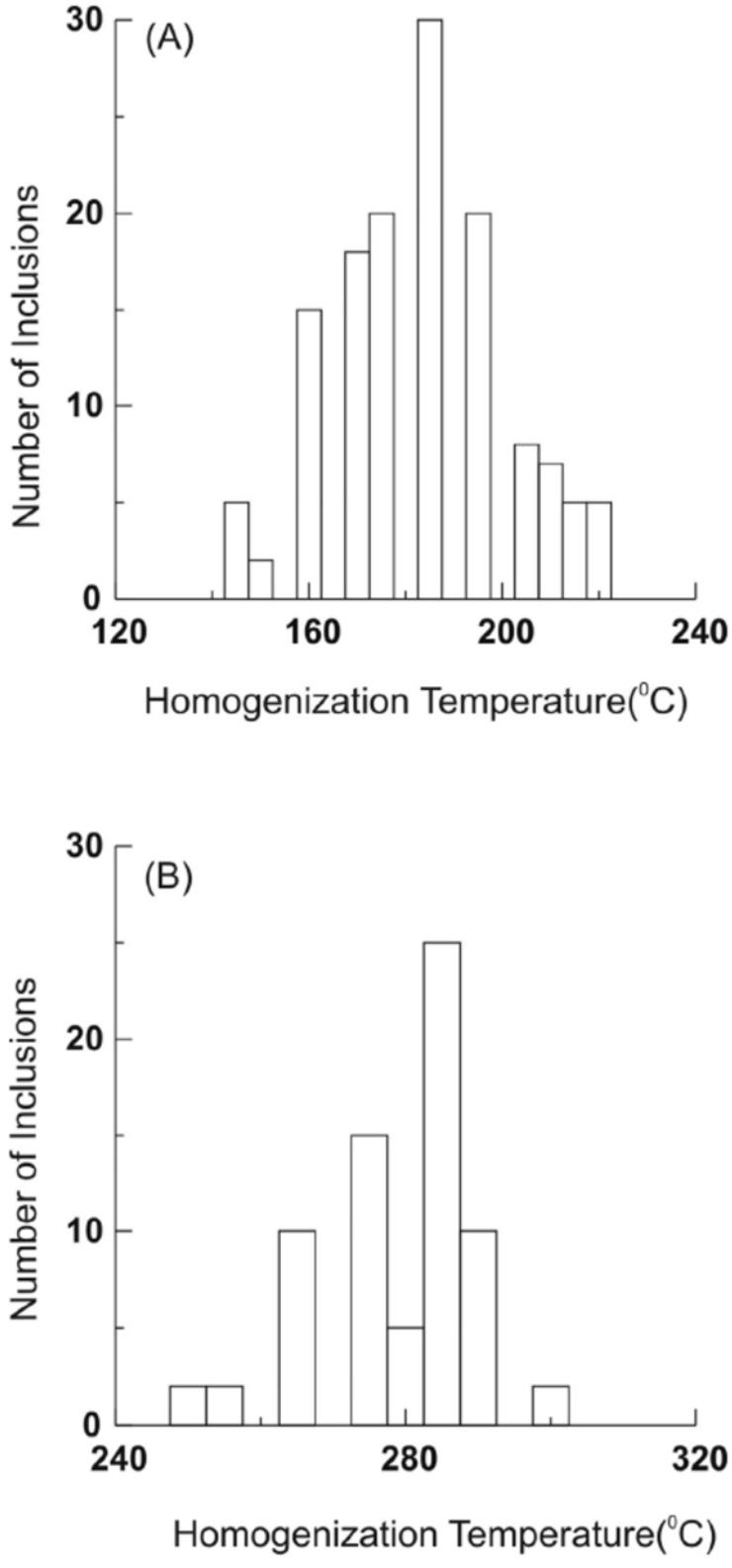

628

629 Fig- 7

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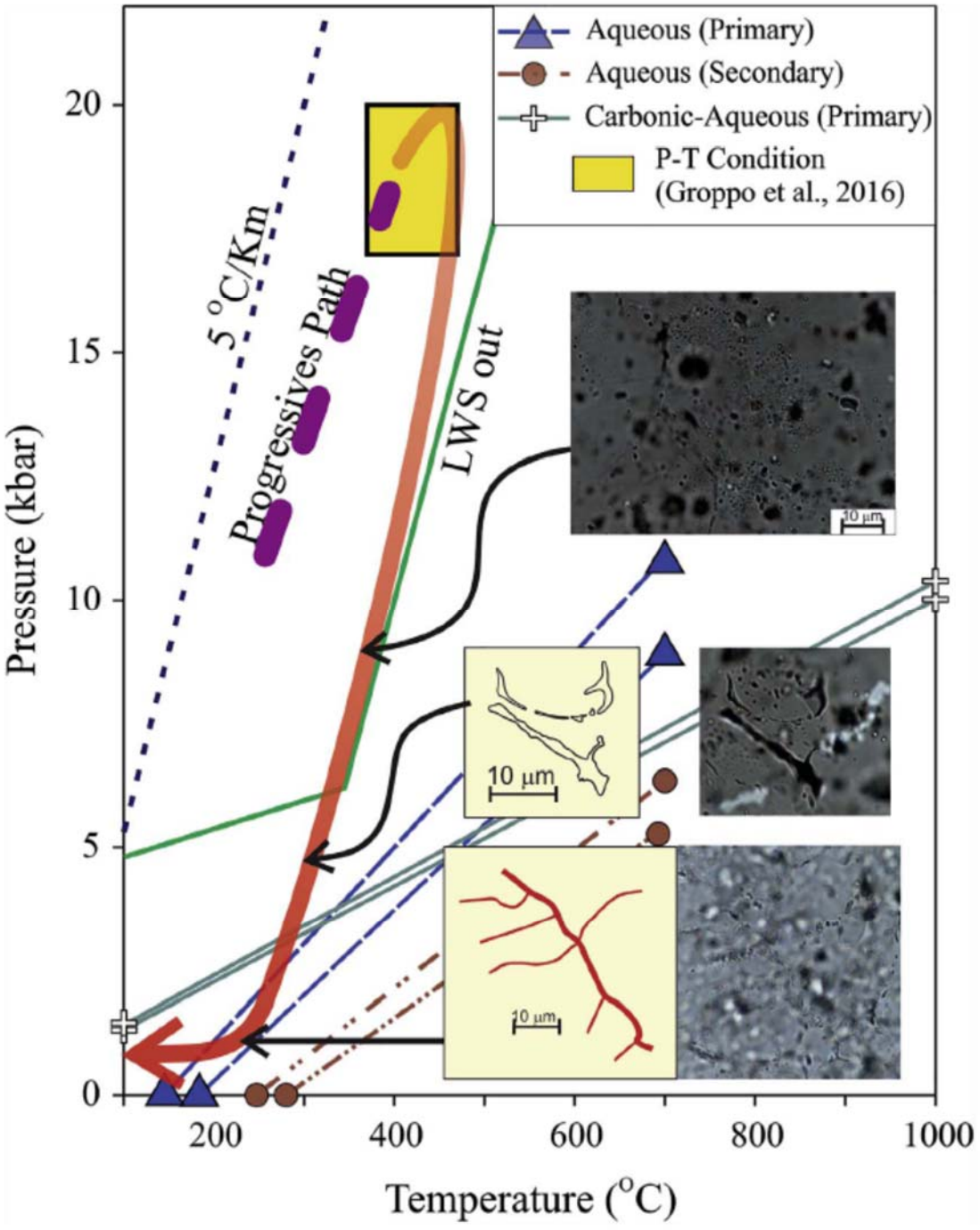
Aqueous (Primary)
Aqueous (Secondary)
Carbonic-Aqueous (Primary)
P-T Condition
(Groppo et al., 2016)
5 °C/Km
Progressives Path
LWS out
10 µm
10 µm
10 µm
Pressure (kbar)
Temperature (°C)
0
5
10
15
20
200
400
600
800
1000

631

632 Fig.8