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Global occurrence and chemical impact of stratospheric Blue Jets modeled with WACCM4

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Abstract

In this work we present the first parameterizations of the global occurrence rate and chemical influence of Blue Jets, a type of Transient Luminous Event (TLE) taking place in the stratospheric region above thunderclouds. These parameterizations are directly coupled with five different lightning parameterizations implemented in the Whole Atmosphere Community Climate Model (WACCM4). We have obtained a maximum Blue Jet global occurrence rate of about 0.9 BJ per minute. The geographical occurrence of Blue Jets is closely related to the chosen lightning parameterization. Some previously developed local chemical models of Blue Jets predicted an important influence onto the stratospheric concentration of N_2O , NO_x and O_3 . We have used these results together with our global implementations of Blue Jets in WACCM4 to estimate their global chemical influence in the atmosphere. According to our results, Blue Jets can inject about 3.8 Tg N_2O -N yr^{-1} and 0.07 Tg NO -N yr^{-1} near the stratosphere, where N_2O -N and NO -N stand for the mass of nitrogen atoms in N_2O and NO molecules, respectively. These production rates of N_2O and NO_x could have a direct impact on, for example, the acidity of rainwater or the greenhouse effect. We have found that Blue Jets could also slightly contribute to the depletion of stratospheric ozone. In particular, we have estimated that the maximum difference in the concentration of O_3 at 30 km of altitude between simulations with and without Blue Jets can be about -5 % in Equatorial and Polar regions.

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1 Introduction

The electric field produced as a consequence of the separation of electrical charges inside clouds is the origin of lightning in the troposphere. However, as originally proposed by *Wilson* [1925] and later detected by *Franz et al.* [1990], atmospheric electrical discharges can also take place in upper regions of the atmosphere. These types of electrical phenomena are known as Transient Luminous Events (TLEs).

The first detected TLE was a sprite [*Franz et al.*, 1990], an upper atmospheric discharge formed by a complex structure of thousands of streamers and a diffuse non-streamer zone that can extend from 40 km up to 85 km of altitude [*Sentman and Wescott*, 1994, *Lyons*, 1994]. Other types of TLEs, known as halos and elves, can also be produced in the upper atmosphere at altitudes greater than 70 km and 80 km. Both halos and sprites can have a duration of several milliseconds [*Lyons et al.*, 2000, *Barrington-Leigh et al.*, 2001, *Wescott et al.*, 2001a, *Bering et al.*, 2002, *Moudry et al.*, 2003, *Bering et al.*, 2004b,a, *Frey et al.*, 2007, *Sentman et al.*, 2008, *Gordillo-Vázquez*, 2008, *Luque and Gordillo-Vázquez*, 2011], while elves have a duration of less than 1 ms [*Inan et al.*, 1991, 1997, *Taranenko et al.*, 1993, *Moudry et al.*, 2003, *Kuo et al.*, 2007, *Marshall et al.*, 2010, *Gordillo-Vázquez et al.*, 2016, *van der Velde and Montanyà*, 2016, *Pérez-Invernón et al.*, 2018a,b].

In 1995, *Wescott et al.* [1995, 1996] discovered the existence of upward propagating conical-shaped jets in the ranges of altitudes between 15 km and 25 km. Later in year 2002, *Pasko et al.* [2002] reported another type of upward propagating jets that reached the ionosphere. These upward propagating discharges were later called Blue Jets (BJ) and Gigantic Jets (GJ), two types of TLEs that can propagate in the range of altitudes between 15 km and 40 km in the case of Blue Jets, and between 15 km and 90 km in the case of GJs. The upper altitude reached by Blue Jets (about 40 km) corresponds to the level where propagation time equals the relaxation timescale of about 0.2 s [*Sukhorukov et al.*, 1996]. Blue Jets and Gigantic Jets are different events triggered right above the cloud layer [*Pasko et al.*, 2002, *van der Velde et al.*, 2010, *Pasko et al.*, 2012, *Chanrion et al.*, 2017]. According to some evidences [*Krehbiel et al.*, 2008, *Riousset et al.*, 2010, *Pasko et al.*, 2012], Gigantic Jets could be initiated as a cloud lightning discharge propagating upward, while Blue Jets are triggered as a consequence of the electrical breakdown produced between the storm upper charge layer and the screening charge attracted to the cloud top [*Krehbiel et al.*, 2008, *Riousset et al.*, 2010, *Pasko et al.*, 2012].

Since their discovery in 1989, TLEs have been observed from planes, balloons, ground-based detectors and space-based instrumentation. Several campaigns have recorded the spectra of sprites [*Hampton et al.*, 1996, *Kanmae et al.*, 2007, *Passas et al.*, 2016, *Gordillo-Vázquez et al.*, 2018]. Some space-based missions, such as the Space Shuttle [*Boeck et al.*, 1992], the Imager of Sprites and Upper Atmospheric Lightning (ISUAL) of the National Space Organization of Taiwan (NSPO) [*Chern et al.*, 2003, *Chen et al.*, 2008, *Hsu et al.*, 2017] and the Global Lightning and sprItE Measurements (GLIMS) of the Japan Aerospace Exploration Agency (JAXA) [*Sato et al.*, 2015, *Adachi et al.*, 2016] have reported TLE observations from space. Last April 2, 2018 the Atmosphere-Space Interactions Monitor (ASIM) [*Neubert et al.*, 2006] of the European Space Agency (ESA) was successfully launched. ASIM is equipped with the Modular Multi-

Imaging Assembly (MMIA), devoted to the study of TLEs from space. In addition, the Tool for the Analysis of RADIations from lightNING and Sprites (TARANIS) [Blanc *et al.*, 2007] of the Centre National d'Études Spatiales (CNES), will also be devoted to the observation of these events after its expected launch in 2019 or 2020.

Several authors have investigated the local chemical impact of TLEs [Gordillo-Vázquez, 2008, Sentman *et al.*, 2008, Gordillo-Vázquez and Donkó, 2009, Gordillo-Vázquez, 2010, Pasko *et al.*, 2012, Parra-Rojas *et al.*, 2013, 2015, Winkler and Notholt, 2015, Pérez-Invernón *et al.*, 2016, Hoder *et al.*, 2016, Pérez-Invernón *et al.*, 2018a]. Recently, Winkler and Notholt [2015] developed a local chemical model of Blue Jets obtaining an important local enhancement of NO_x , N_2O and O . The global chemical influence of TLEs has been investigated by previous studies. According to previous local models of halos and elves [Pérez-Invernón *et al.*, 2018a], their global chemical impact would be negligible. Arnone *et al.* [2014] estimated the global production of NO_x by sprites using the Whole Atmosphere Community Climate Model version 4 (WACCM4). Arnone *et al.* [2014] found that a perturbation in the tropical concentration of nitrogen oxide by sprites could lie between 0.015 ppbv and 0.15 ppbv. These quantities correspond to a perturbation of the background concentration of NO_x between less than 1 % and up to 20 % at different altitudes. Some observational studies have attempted to measure sprite- NO_x through satellite observations [Arnone *et al.*, 2008, Rodger *et al.*, 2008, Arnone *et al.*, 2009, Arnone and Dinelli, 2016]. However, according to these studies sprite- NO_x is at the edge of current detectability. The predicted significant local chemical influence [Winkler and Notholt, 2015] suggest that Blue Jets could have a non-negligible influence in the chemistry of the atmosphere.

In this work, we have developed the first global parameterization of Blue Jets. We have used the WACCM4 model in order to study the global occurrence rate of Blue Jets and their global chemical impact by developing three different Blue Jet parameterizations. WACCM4 includes a lightning parameterization developed by Price and Rind [1992] based on the cloud top height (CTH). Here we also use other lightning parameterizations based on, respectively, the amount of convective precipitation (CP) [Allen and Pickering, 2002]; the upward mass flux (MFLUX) [Allen and Pickering, 2002]; the precipitation rate and the Convective Available Potential Energy (CPCAPE) [Romps *et al.*, 2014]; and on the upward cloud ice flux (ICEFLUX) [Finney *et al.*, 2014]. The combined use of lightning and Blue Jet parameterizations allow us to predict the geographical and seasonal chemical impact of Blue Jets.

1.1 Physics and chemistry of Blue Jets

Blue Jets are formed by a leader channel surrounded by a large number of streamers. The leader is a highly conductive plasma channel that can heat the air up to thousands of Kelvin. The first interpretations of Blue Jets by a streamer corona of a leader were made by Sukhorukov and Stubbe [1998] and Petrov and Petrova [1999]. Raizer *et al.* [2006, 2007] proposed the development of Blue Jets as a bi-leader channel that propagates upward from the streamer zone of a positive leader. According to Raizer *et al.* [2006, 2007], this leader can transfer the energy contained in the clouds to upper regions of the atmosphere, where the low density allows the development of a streamer corona. Figure 1 shows a photography of a real Blue Jet from Wescott *et al.* [2001b] and

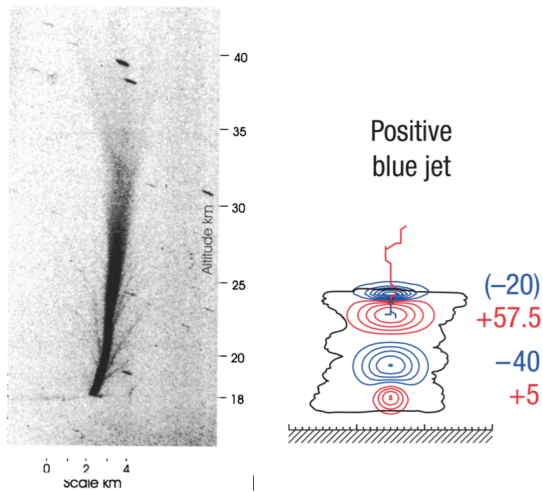


Figure 1: Left panel: Inverted black and white photography of a Blue Jet. The spatial scales of the leader and streamer regions can be appreciated. Image adapted from *Wescott et al.* [2001b]. Right panel: Blue Jet simulated by *Krehbiel et al.* [2008] illustrating the charge structure of clouds. Blue and red lines correspond to positive and negative charges, respectively. Image adapted from *Krehbiel et al.* [2008].

the structure of charges in clouds that trigger the inception of Blue Jets simulated by *Krehbiel et al.* [2008]. *Krehbiel et al.* [2008] developed a model based on quasielectrostatic fields, formed as an imbalance of the electric charge in the cloud tops, to predict Blue Jet and Gigantic Jet inception. After lightning occurs, a charged layer can remain near the storm top layer creating a local electric field. *Krehbiel et al.* [2008] found that conventional electric breakdown near this charged layer could trigger an upward propagating leader, forming a Blue Jet. *Riousset et al.* [2010] upgraded the model proposed by *Krehbiel et al.* [2008], confirming the obtained results. Observations by *Lu et al.* [2011] supported some of the predictions by *Krehbiel et al.* [2008] and *Riousset et al.* [2010]. As hypothesized by *Krehbiel et al.* [2008], there would exist a competition between intra-cloud discharges and Blue Jets in the process of discharging the cloud. The result of this competition would depend on the capability of the convective fluxes to mix the oppositely charged layers located in the cloud top before the inception of a Blue Jet.

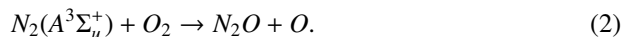
Chanrion et al. [2017] have recently described the observation of a Blue Jet from the International Space Station (ISS). The top height of the thundercloud that initiated the Blue Jet reached the tropopause, an atmospheric region where convection is weak. The reported Blue Jet was preceded in 1.16 s by a strong negative CG lightning with a peak current of -167.5 kA. This CG discharge could possibly be the parent lightning of the Blue Jet. A Blue Jet would then be formed by an upward propagating leader traveling to upper regions of the atmosphere with lower pressure, reaching its maximum altitude between 30 km and 40 km [*van der Velde et al.*, 2010, *Pasko et al.*, 2012, *da*

Silva and Pasko, 2013, Milikh et al., 2014, Chanrion et al., 2017]. Streamers could then emerge from the leader as it passes through the low pressure regions of the atmosphere [*Raizer et al., 2007*].

Mishin [1997] and *Smirnova et al. [2003]* developed the first models to estimate the local chemical impact of Blue Jets. However, these models do not include the latest results on the electrodynamical mechanisms of Blue Jets [*Raizer et al., 2007, Krehbiel et al., 2008, Rioussset et al., 2010*]. *Winkler and Notholt [2015]* developed the most detailed model to date to study the local chemical impact of a Blue Jet including 88 species interacting through more than 1000 reactions. They used their model to estimate the local chemical impact of the leader and streamers of a Blue Jet at several altitudes. According to their estimations, the high-temperature reactions taking place in the Blue Jet leader can enhance by several orders of magnitude the local background concentrations of stratospheric N_2O and NO (due to the high temperature reactions collected in Table 3 of [*Winkler and Notholt, 2015*]) and produce a significant depletion of ozone [*Winkler and Notholt, 2015*]. In addition, the high electric field in the streamer phase would produce an enhancement in the concentration of N_2O by the chemical reactions



and



The injection of NO_x into the stratosphere could also influence the concentration of other species. According to investigations about the chemical influence of lightning-produced NO in the atmosphere, atmospheric electricity phenomena can also contribute to the concentration of OH , HO_2 and CO [*Rohrer and Berresheim, 2006, Murray et al., 2013, Singh et al., 2015*]. In particular, NO interacts with HO_2 producing OH . The production of OH molecules can influence the acidity of rainwater [*Seinfeld and Pandis, 2016*], as they can react with NO_2 molecules producing HNO_3 following the chemical reaction $NO_2 + OH + M \rightarrow HNO_3 + M$ [*Labrador et al., 2005*]. The formation of OH contributes to the loss of CO by the process $CO + OH \rightarrow HO_2 + C$ [*Murray et al., 2013*]. OH molecules can also contribute to the oxidation of SO_2 , leading to the production of H_2SO_4 . In addition, NO_2 molecules contribute to the production of N_2O_5 . The oxidation of N_2O_5 followed by a heterogeneous hydrolysis reaction on aerosol particles contributes to the enhancement of HNO_3 .

1.2 Global budgets of N_2O and NO_x and their relation with atmospheric electricity

According to *Winkler and Notholt [2015]*, Blue Jets could inject an important amount of nitrous oxide (N_2O), nitric oxide (NO) and nitrogen dioxide NO_2 at stratospheric altitudes. These gases play important roles in the chemical balance of stratospheric ozone. In addition, N_2O is one of the most important greenhouse gases.

Natural N_2O sources are estimated to inject about $10.2 \text{ Tg } N_2O\text{-N yr}^{-1}$ in the atmosphere, while anthropogenic sources could produce around $6.3 \text{ Tg } N_2O\text{-N yr}^{-1}$ [*Davidson, 2009, Prather et al., 2015*], where $1 \text{ Tg} = 10^{12} \text{ g}$ and $N_2O\text{-N}$ stands for the mass

of nitrogen atoms in N_2O molecules [Davidson, 2009]. The major natural and anthropogenic sources of N_2O are basically due to nitrification and denitrification produced by microbes at ground level [Davidson, 2009]. However, Sheese *et al.* [2016] have recently proposed an atmospheric source of N_2O based on observations from the satellite instrument “Atmospheric Chemistry Experiment-Fourier Transform Spectrometer” (ACE-FTS) consisting of the chemical reaction described in equation (2) [Arnone and Hauchecorne, 2012, Sheese *et al.*, 2016]. N_2O is the major source of NO in the stratosphere. 90 % of the stratospheric destruction of N_2O is by photolysis ($\text{N}_2\text{O} + h\nu \rightarrow \text{N}_2 + \text{O}$) and 10 % is by reaction with $\text{O}(^1\text{D})$ producing NO, N_2 and O_2 molecules [Seinfeld and Pandis, 2016].

Plieninger *et al.* [2016] compared global-average vertical profiles of N_2O obtained by different instruments. In particular, Plieninger *et al.* [2016] showed the vertical stratospheric concentration of N_2O obtained by the “Michelson Interferometer for Passive Atmospheric Sounding” (MIPAS), the “Atmospheric Chemistry Experiment-Fourier Transform Spectrometer (ACE-FTS), the “Microwave Limb Sounder onboard Aura” (Aura-MLS) and the “Sub-Millimetre Radiometer onboard Odin” (Odin-SMR). It is worth noting that the global-average concentration of N_2O estimated by each of the above mentioned instruments between 20 km and 40 km indicates that the observational uncertainty in the global amount of N_2O is about 10 % [Plieninger *et al.*, 2016].

Lightning is not considered an important source of atmospheric N_2O , as shown in Table 11 of Schumann and Huntrieser [2007] where results from different studies and campaigns conclude that the global lightning-produced emission rate of N_2O is below 5×10^{-4} Tg $\text{N}_2\text{O-N yr}^{-1}$.

Let us now turn to the global budget of NO and NO_2 , which together make up NO_x . Schumann and Huntrieser [2007] presented an extensive study about the global production of NO_x by lightning (or LNO_x) based on satellite and aircraft measurements, laboratory experiments and theoretical studies. Lightning is considered one of the major natural sources of atmospheric NO_x emissions. Different studies estimate the global production of NO_x in thunderstorms in a wide range between 1 and 20 Tg NO-N yr^{-1} [Schumann and Huntrieser, 2007, Huntrieser *et al.*, 2016]. However, the most likely range is 5 ± 3 Tg N yr^{-1} . Lightning would then contribute up to ~10-15% of the total global emissions of NO_x . It is probably something greater than 10% now that anthropogenic emissions have decreased substantially in North America and Europe.

The uncertainties in the contribution of lightning to the global concentration of NO_x is based on theoretical and empirical challenges. Laboratory results of NO_x produced by electrical discharges are difficult to extrapolate to real lightning discharges, as both Cloud-to-Ground (CG) and Intra-Cloud (IC) lightning discharges are different from each other [Price *et al.*, 1997] and cannot be accurately reproduced in the laboratory. Space-based instruments measure NO_2 and cannot accurately measure the concentration of tropospheric NO_x [Schumann and Huntrieser, 2007, Beirle *et al.*, 2010, Bucsela *et al.*, 2010, Pickering *et al.*, 2016]. This concentration has to be usually deduced from the concentration of other species that can react with NO_x molecules. However, some uncertainties in the atmospheric chemical kinetics of NO_x lead to imprecisions in the estimation of NO_x from measurements. These estimations are often based on an assumed upper tropospheric (UT) chemical lifetime of NO_x in a range between 2 and 8 days. Based on reanalysis of the measurements taken by the Deep Convective Clouds

and Chemistry (DC3) atmospheric experiment, *Nault et al.* [2017] recently revised the interaction of atmospheric $\text{CH}_3\text{O}_2\text{NO}_2$ and HNO_3 with NO_x molecules, estimating a new UT NO_x lifetime of about 3 hours in the first few hours downwind of a thunderstorm instead of the previous scale of days. Using this new analysis, *Nault et al.* [2017] estimated a global lightning production of NO_x of about $9 \text{ Tg NO-N yr}^{-1}$. *Nault et al.* [2017] results indicate higher LNO_x in the mid-latitudes than in the tropical regions, in agreement with *Schumann and Huntrieser* [2007]. However, the latest estimations of the global lightning NO_x emissions by new cloud-sliced observations of UT NO_2 in the 6 km - 9 km range from the Ozone Monitoring Instrument (OMI) of the Aura mission combined with the GEOS-Chem model point to a global lightning NO_x source of 5.5 Tg N yr^{-1} [*Marais et al.*, 2018]. *Marais et al.* [2018] reports no significant difference in LNO_x production per flash between the tropics and mid-latitudes. Stratospheric NO can cause ozone depletion through the processes [*Crutzen*, 1979]



Moreover, the oxidation of N_2O is the major source of stratospheric NO, producing 1 Tg N yr^{-1} of NO_x [*Crutzen*, 1979]. Thus, the introduction of Blue Jets in global models as a new possible atmospheric source of N_2O and NO_x could have a non negligible effect in the global budget of ozone.

2 Model

2.1 WACCM4

The Whole Atmosphere Community Climate Model version 4 (WACCM4) [*Marsh et al.*, 2013] is a global circulation model included in the Community Earth Climate System Model version 1 (CESM1). WACCM4 is an extension of the Community Atmosphere Model (CAM4) [*Marsh et al.*, 2013, *Tilmes et al.*, 2015, 2016]. CAM4 couples the troposphere and the stratosphere chemistry, while WACCM4 extends up to the thermosphere. The chemistry of WACCM4 is based on version 4 of the Model for Ozone And Related chemical Tracers (MOZART4) [*Kinnison et al.*, 2007, *Emmons et al.*, 2010, *Lamarque et al.*, 2012, *Tilmes et al.*, 2015], including 183 species and 472 chemical reactions including gas-phase chemistry of neutrals and ions, photolysis and heterogeneous chemistry.

We set the model domain extending from the surface to 140 km of altitude ($5.96 \times 10^{-6} \text{ hPa}$). We divide the vertical domain in 88 levels and set a horizontal resolution of 1.9° in latitude and 2.5° in longitude. We start the numerical experiment with WACCM4 running a complete year (from January 1999 to January 2000) without Blue Jets allowing free dynamics for each considered lightning parameterization. Then, we run the same period of time in the specified dynamics mode (SD-WACCM4) [*Lamarque et al.*, 2012, *Smith et al.*, 2017]. In this study, we use the facility of SD but, instead of nudging to reanalysis fields, we nudge to the meteorological fields from a previous (free-running)

WACCM simulation. The reason for using SD is to ensure that the basic dynamics in the lower and middle atmosphere is identical in simulations in which other changes are made. In this second run, temperature fields and horizontal winds in the troposphere and stratosphere are nudged at each model time step using the output of the first run. We apply the nudging from ground level to 80 km. The nudging is then tapered off in the ranges of altitude between 80 and 90 km, and finally removed at 90 km of altitude [Smith *et al.*, 2017].

Afterward, we use the same specified dynamics in order to run a complete year including all the combinations of lightning and Blue Jet parameterizations. This approach allows us to compare the simulations with and without Blue Jets in order to estimate their global chemical impact in the atmosphere.

As the lifetime of N₂O in the atmosphere is of the order of a century [Prather *et al.*, 2015], we select the most realistic cases and repeat the process for a period of one decade. Following this approach, the obtained results would be closer to the chemical equilibrium. We discuss these cases in section 3.

2.2 Lightning parameterizations

The temporal and geographical occurrence of Blue Jets obtained with WACCM4 will strongly depend on the global occurrence of lightning. In this section, we briefly highlight the particularities of each considered lightning parameterization.

The characteristic size of lightning is some orders of magnitude smaller than the WACCM4 grid size. Therefore, lightning are considered as sub-grid events in the model. WACCM4 includes a lightning parameterization based on the cloud top heights (CTH) [Price and Rind, 1992] that estimates the density of lightning (or flashes) in each domain cell for every time step of 30 minutes. The regional and seasonal flash frequency produced by this parameterization roughly agrees with the observations recorded by the Lightning Imaging Sensor (LIS) and the Optical Transient Detector (OTD) [Christian *et al.*, 2003, Cecil *et al.*, 2014] over a period of two decades. However, the implementation of this lightning parameterization in WACCM4 overestimates the total flashes per second taking place in the globe over a period of one year. According to OTD/LIS, the global lightning occurrence over a year is around 44 flashes per second, while this parameterization produces around 65 flashes per second. In addition, the lightning parameterization by Price and Rind [1992] implemented in WACCM4 also underestimates the lightning occurrence over the oceans. For these reasons, the use of a parameterization for Blue Jets together with the CTH lightning parameterization by Price and Rind [1992] would probably underestimate the occurrence of Blue Jets over the oceans and would overestimate the global occurrence of Blue Jets. However, the spatial correlation between the flash frequency reported by OTD/LIS and the flash frequency estimated by CTH is 0.7602. This is the highest spatial correlation obtained by the use of different lightning parameterizations. Therefore, we choose the CTH lightning parameterization to show the primary results.

Allen and Pickering [2002] developed a lightning parameterization based on the amount of convective precipitation (CP) over USA. We have tested this lightning parameterization in WACCM4, obtaining a good agreement between the predicted flash frequency (51 flashes per second) and the lightning occurrence reported by OTD/LIS.

We obtain a spatial correlation between the flash frequency derived by CP and OTD/LIS of 0.5760. However, the implementation of this parameterization in WACCM4 produces a lightning occurrence that remains almost constant over the four seasons, in disagreement with observations. *Allen and Pickering* [2002] also derived a lightning parameterization based on the upward mass flux (MFLUX) at 440 hPa. This parameterization produces again a good agreement between the predicted flash frequency (43 flashes per second) and the lightning occurrence reported by OTD/LIS. However, it slightly overestimates the occurrence of lightning in the oceans and in South America, while underestimates the flash density in some regions of Africa. In addition, MFLUX produces a low spatial correlation (0.4963) with the flash frequency reported by OTD/LIS.

Apart from these three “classical” lightning parameterizations by *Price and Rind* [1992] and *Allen and Pickering* [2002], we also implement Blue Jet parameterizations in WACCM4 together with two, more recent, lightning parameterizations developed by *Romps et al.* [2014] (CPCAPE) and *Finney et al.* [2014] (ICEFLUX), respectively. The parameterization of *Romps et al.* [2014] is based on the precipitation rate and on the Convective Available Potential Energy (CAPE), while the parameterization by *Finney et al.* [2014] is based on the upward cloud ice flux at 440 hPa. The parameterization by *Romps et al.* [2014] produces global (52 flashes per second), regional and seasonal lightning frequencies that agree with the observation by OTD/LIS but it overestimates the flash occurrence over the oceans. The spatial correlation between the flash frequency derived by CPCAPE and the observations of OTD/LIS is 0.4540. The implementation of the parameterization developed by *Finney et al.* [2014] underestimates by a factor of 2 the global lightning occurrence rate and results in a spatial correlation with the flash frequency reported by OTD/LIS of 0.6739.

2.3 Blue Jet parameterizations in WACCM4

2.3.1 Blue Jet frequency

The characteristic size of Blue Jets is some orders of magnitude smaller than the horizontal size of WACCM4 grids. As in the case of lightning, Blue Jets have to be treated as sub-grid phenomena in WACCM4. Following the basic idea of the previously described global lightning parameterizations, we have developed two different Blue Jet parameterizations to be considered in global models. The first developed parameterization prescribes the estimation of global Blue Jets per minute to predict their chemical influence in the atmosphere. The second proposed parameterization is based on physical assumptions and does not impose the rate of occurrence of Blue Jets.

Parameterization based on ISUAL and the altitude of the tropopause (IS-TROP LOW / IS-TROP UP)

Ignaccolo et al. [2006] proposed a formula to estimate the global rate of sprites based on reports of sprite detections. *Ignaccolo et al.* [2006] obtained a global occurrence rate of sprites about 2.8 per minute. According to ISUAL, the global occurrence rate of TLEs is around 4.13 per minute *Chen et al.* [2008], among which 3.23 are elves, 0.50

are sprites, 0.39 are halos and 0.01 are gigantic jets. As optical emissions from Blue Jets and lightning are difficult to separate, the global occurrence rate of Blue Jets was not derived from ISUAL data. However, we assume that Blue Jets are less frequent than sprites and more frequent than Gigantic Jets. Therefore, as a first approximation we assume that the global occurrence rate of Blue Jet is in the range between 0.01 and 1.0 events per minute. Given that Blue Jets are triggered as a consequence of the remaining imbalance of charge in thunderclouds after lightning occurs [Krehbiel *et al.*, 2008], Blue Jet parameterizations must be spatially and temporally connected with any considered parameterization of lightning. Following these considerations, the total occurrence of Blue Jets at a given time would be the total occurrence of lightning flashes given multiplied by 3.6×10^{-4} (UP) or 3.6×10^{-6} (LOW), respectively.

As we discussed in section 1.1, the model proposed by Krehbiel *et al.* [2008] indicates that the inception of Blue Jets is favored when the two oppositely charged layers located in the upper part of thunderclouds do not mix. Hence, it is reasonable to expect the inception of Blue Jets when the convection near the cloud top is weak. In this regard, the Blue Jet reported by Chanrion *et al.* [2017] was triggered in a thundercloud whose top height was near the tropopause, where the lack of convection keeps the temperature relatively constant. WACCM4 and the most of Global Circulation Models can calculate the altitude of the tropopause and the cloud top height, two atmospheric variables that can be related with the possibility of Blue Jet inception [Krehbiel *et al.*, 2008, Chanrion *et al.*, 2017]. We can then distribute the predicted Blue Jets exclusively in the domain cells where the cloud top height is above the beginning of the tropopause or below tropopause by no more than one kilometer and there is lightning. We also impose as a condition to the existence of Blue Jets that the flash frequency is greater than zero in that domain cell. We restrict the locations where Blue Jets can be distributed to the range of latitudes between 60° S and 60° N. We name this Blue Jet parameterization as “IS-TROP LOW” and “IS-TROP UP”, depending on whether the global occurrence rate of BJ is set to 3.6×10^{-6} (LOW) or 3.6×10^{-4} (UP) Blue Jets per lightning and where IS and TROP refer to ISUAL and to the height of the tropopause, respectively. Although this parameterization could produce a realistic geographical occurrence of Blue Jets, the global occurrence rate is somehow imposed.

Parameterization based on lightning peak currents and the altitude of the tropopause (LPC-TROP LOW / LPC-TROP UP)

According to the model developed by Krehbiel *et al.* [2008], a strong lightning discharge or a set of small amplitude CG lightning discharges occurring within a short time distance would probably precede the inception of a Blue Jet, since Blue Jets are produced by a large imbalance of charge. Chanrion *et al.* [2017] reported the observation of a Blue Jet preceded by a strong lightning discharge. Let us now use this observation to derive a more realistic Blue Jet parameterization based on the peak current value of the lightning possibly preceding a Blue Jet. The Blue Jet reported by Chanrion *et al.* [2017] was preceded by a negative Cloud-to-Ground (CG) lightning discharge with a peak current of -167.5 kA. Elves, the less energetic TLEs, seem to be triggered by lightning discharges with peak currents whose absolute value is above 60 kA [Barrington-Leigh and Inan, 1999]. As an approximation, we can assume that

the peak current threshold of the lightning preceding Blue Jets is between 60 kA and 167.5 kA. We choose two representative values in this range to be the threshold of Blue Jets, such as 100 kA and 150 kA. According to the distribution of global lightning peak current reported by *Said et al.* [2013] using the Vaisala global lightning data set GLD360, approximately 1 % of lightning have a peak current above 100 kA and only 0.1 % have a peak current above 150 kA. We can then develop a Blue Jet parameterization in WACCM4 where the spatial occurrence of Blue Jets is again restricted to domain cells where the cloud top height is higher than one kilometer below the tropopause. However, instead of imposing the global occurrence of Blue Jets, we can now assume that the Blue Jet frequency in such domain cells is given by the amount of lightning in the domain cell with peak currents above 100 kA or 150 kA. Hence, we define the Blue Jet frequency in each cell where the cloud top height is near the tropopause as 0.01 (UP) or 0.001 (LOW) times the frequency of lightning. We refer to these two Blue Jet parameterizations as “LPC-TROP UP” and “LPC-TROP LOW”, where LPC and TROP recall to lightning peak current and to the height of the tropopause, respectively. The maximum peak current of lightning is not homogeneously distributed over land and ocean [*Said et al.*, 2013]. However, we do not include in this parameterization any parameter to take into account this inhomogeneity. This simplification is justified because the scope of this paper is to describe the global chemical influence of Blue Jets rather than the regional influence.

2.3.2 Chemical impact of Blue Jets

Winkler and Notholt [2015] developed the most detailed zero-dimensional model until now to predict the local chemical impact of a Blue Jet in the center of Blue Jet leader and streamers. *Winkler and Notholt* [2015] estimated the chemical impact of an upward propagating Blue Jet at different altitudes between 18 km and 38 km, obtaining a significant enhancement in the densities of some chemical species such as N_2O , NO or O and a decrease in the density of O_3 in the center of Blue Jets. We use the local chemical impact of a single Blue Jet obtained by *Winkler and Notholt* [2015] together with the previously derived Blue Jet parameterizations to estimate the global chemical impact of Blue Jets using WACCM4. We assume that each Blue Jet would produce an enhancement in the concentrations of N_2O , NO or O. However, as Blue Jets are considered as sub-grid phenomena in WACCM4, we take into account the following considerations:

1. As the area of a WACCM4 cell is larger than the horizontal cross-section of the Blue Jet, we have to estimate the total number of species produced by all Blue Jets at each altitude and distribute them over the area of the grid at each altitude level. Hence, we need to estimate the electrodynamical radius of the Blue Jet where the chemical reactions are produced. *Winkler and Notholt* [2015] noted that while the optical radius of a Blue Jet is a few hundreds of meters [*Wescott et al.*, 2001b], the electrodynamical radius could be between 20 and 100 times smaller [*Shneider et al.*, 2012, *Milikh et al.*, 2014]. Based on optical observations by *Wescott et al.* [2001b], we have assumed that Blue Jets have an optical radius of 250 m at their base. Hence, the electrodynamical radius

R_e would be in the range between $R_1 = 2.5$ m and $R_2 = 12.5$ m. According to *Winkler and Notholt* [2015], the production of N_2O and NO is dominated by the leader for altitudes ranging between 18 km and 28 km and by streamers between 28 km and 38 km. We have then assumed that the electrodynamical radius of the Blue Jet is completely filled by a leader between 18 km and 28 km of altitude and by streamers between 28 km and 38 km of altitude.

The global production of N_2O by Blue Jets can then be estimated using the density variations reported by *Winkler and Notholt* [2015] (figure 19 of *Winkler and Notholt* [2015]). As a first approximation, we can consider a Blue Jet as a 20 km long cylinder with a constant radii of 2.5 m or 12.5 m formed by a leader and a streamer region. A single Blue Jet would produce between 2×10^{28} and 6×10^{29} molecules of N_2O for electrodynamical radii of 2.5 m and 12.5 m, respectively. Assuming that the global occurrence rate of Blue Jets is between 0.01 and 1 per minute [*Ignaccolo et al.*, 2006, *Chen et al.*, 2008], we find that Blue Jets with a radii of 2.5 m would produce between 6×10^{-3} Tg N_2O-N yr^{-1} and 0.6 Tg N_2O-N yr^{-1} , while Blue Jets with a radii of 12.5 m would produce between 0.15 Tg N_2O-N yr^{-1} and 15 Tg N_2O-N yr^{-1} .

2. We assume that the production of species decays parabolically across the radial coordinate r from the center of the Blue Jet up to the limit of the electrodynamical radius as

$$N(r) = N_{max} \left(1 - \frac{r^2}{R_e^2} \right), \quad (5)$$

where R_e is the electrodynamical radius and N_{max} is the enhancement in the density of species in the symmetry axis of the Blue Jet as predicted by *Winkler and Notholt* [2015].

3. Recorded optical emissions from Blue Jets indicate an increase of its radius with altitude [*Wescott et al.*, 2001b]. We assume that the Blue Jet radius increases with altitude following the simple scale law

$$P_i R_i = P_j R_j, \quad (6)$$

where P and R are the atmospheric pressure and Blue Jet radius at two different altitude levels denoted as i and j .

Following these considerations, the Blue Jet region at 30 km of altitude filled by streamers would have a radius between 16 m and 80 m. Given the assumed altitude-dependence of the leader and streamer-region radius and the production profile by *Winkler and Notholt* [2015], the total production of N_2O and NO by a Blue Jet is dominated by the leader phase.

3 Results

We have implemented in WACCM4 the Blue Jet parameterizations derived in section 2.3 using five different lightning parameterizations. We present the obtained global occurrence rate of Blue Jets in subsection 3.1. We have coupled these Blue Jet frequencies with the chemical impact of a single Blue Jet predicted by *Winkler and Notholt [2015]*. In subsection 3.2 we present the predicted global impact of Blue Jets for each global parameterization.

3.1 Global occurrence and seasonal cycle of Blue Jets

Let us firstly focus on the global occurrence of Blue Jets derived for each combination of lightning and Blue Jet parameterizations. The first column of figure 2 shows the obtained lightning flash frequency using different lightning parameterizations in WACCM4. The second column of figure 2 shows the Blue Jet frequency for the Blue Jet parameterization denoted as “IS-TROP LOW” using different lightning parameterizations. As we have previously detailed, this Blue Jet parameterization take as input the global occurrence rate of Blue Jet estimated from the TLE occurrence reported by ISUAL. The Blue Jet frequency obtained with the BJ parameterizations “LPC-TROP LOW” and “LPC-TROP UP” using different lightning parameterizations is plotted in figure 3. In the “LPC-TROP” parameterization, the Blue Jet occurrence rate is not fixed to any given value. Instead, it is based on the physical assumptions previously described in section 2.3.

The Blue Jet frequencies obtained with all the considered Blue Jet parameterizations and with the CTH lightning parameterization are collected in table 1. The results obtained with the other tested lightning parameterizations are shown in the supplementary material.

The Blue Jet parameterizations “IS-TROP” and the most of the Blue Jet parameterizations “LPC-TROP LOW” produce a global occurrence rate of Blue Jets lower than 1 BJ per minute, as estimated from the TLE frequency reported by ISUAL. However, the Blue Jet parameterizations “LPC-TROP UP” and “CP LPC-TROP LOW” significantly overestimate the Blue Jet frequency.

The comparison of the spatial distribution of lightning flashes and Blue Jets indicates that the relative occurrence of Blue Jets in Asia with respect to other regions is larger than the relative occurrence of lightning flashes. In addition, most of the considered parameterizations produce a maximum in the lightning flash frequency and in the Blue Jet frequency over Africa and in the north of South America.

The monthly global average occurrences of Blue Jets obtained with different lightning parameterization schemes are presented in figure 4. The maximum occurrence of Blue Jets takes place between June and August, coinciding with the maximum occurrence of lightning. As the occurrence of Blue Jets is related with a high lightning activity on clouds [*Krehbiel et al., 2008*] (see figure 1), we conclude that the obtained coincidence of the seasonal cycle of lightning and Blue Jets can be considered as realistic.

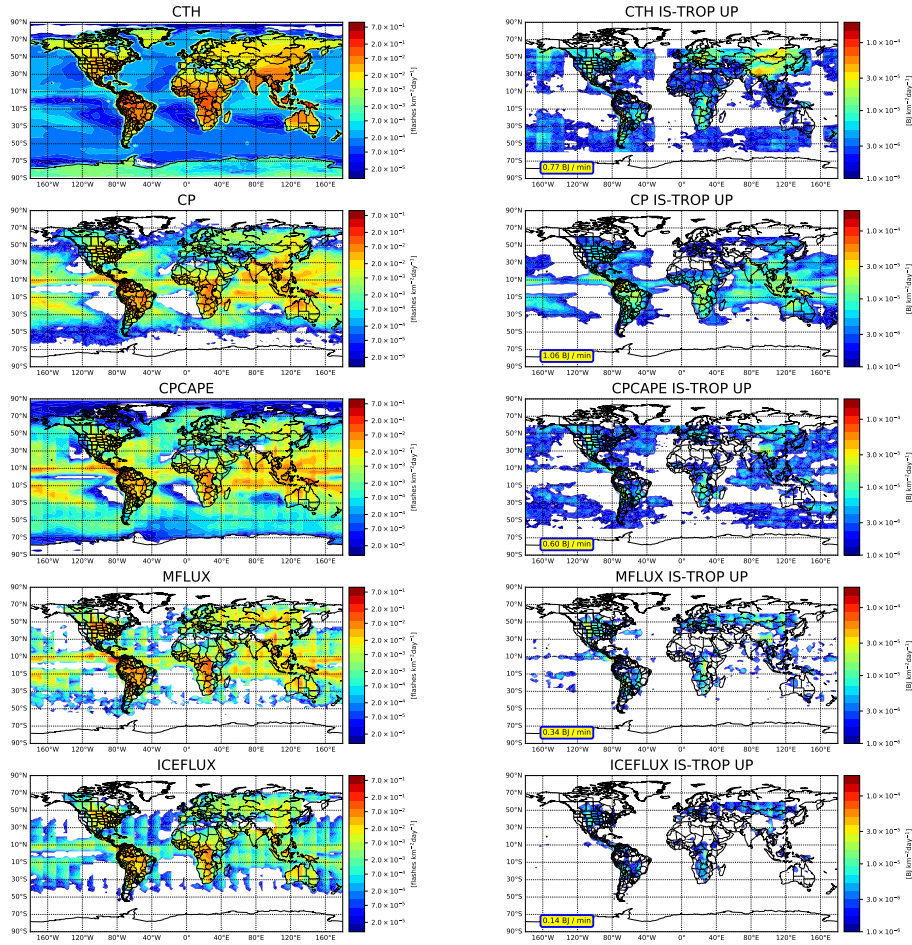


Figure 2: (First column) annual average lightning flash frequencies in flashes $\text{km}^{-2}\text{day}^{-1}$ and (second column) annual average Blue Jet frequencies in BJ $\text{km}^{-2}\text{day}^{-1}$ using the Blue Jet parameterization denoted as “IS-TROP LOW” and different lightning parameterizations. We have used different lightning parameterizations denoted as CTH [Price and Rind, 1992] based on the cloud top height, CP [Allen and Pickering, 2002] based on the precipitation rate, CPCAPE [Romps et al., 2014] based on the precipitation rate and convective available potential energy (CAPE), MFLUX Allen and Pickering [2002] based on the updraft mass flux and ICEFLUX [Finney et al., 2014] based on the upward cloud ice flux. We annotate in boxes the total annual Blue Jets per minute.

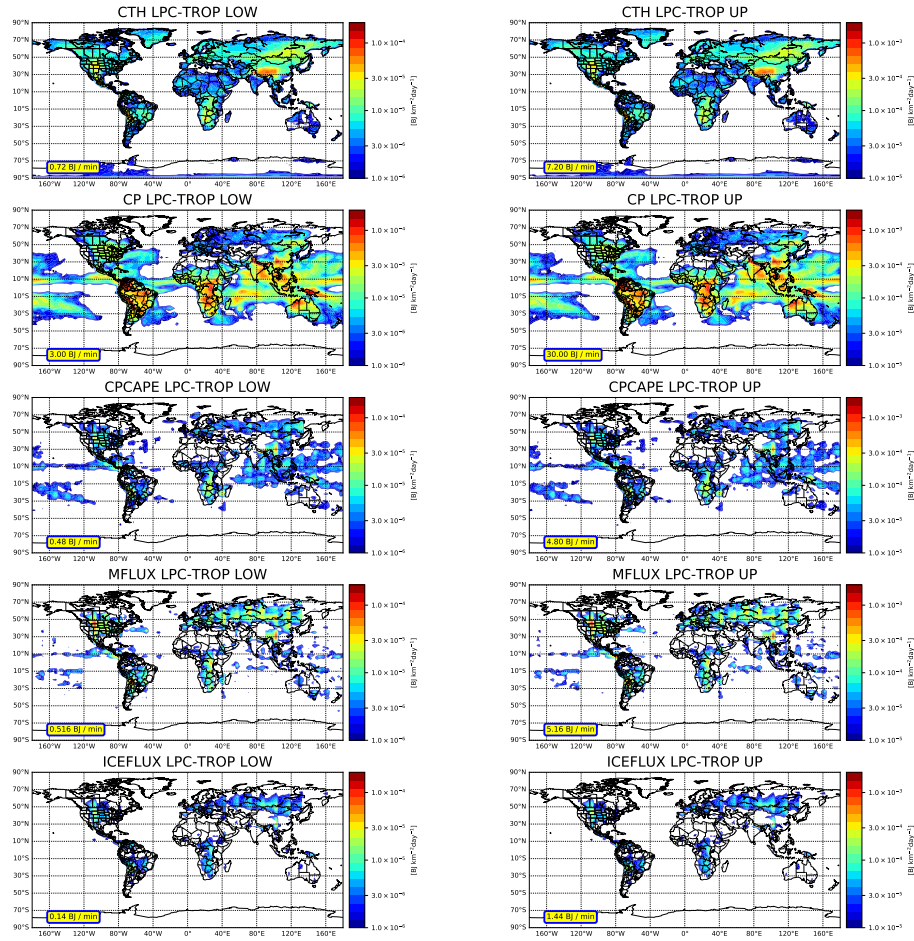


Figure 3: Annual average Blue Jet frequencies $\text{BJ km}^{-2}\text{day}^{-1}$ using the parameterizations denoted as “LPC-TROP LOW” and “LPC-TROP UP”. The shown Blue Jet frequencies have been calculated using different lightning parameterizations denoted as CTH [Price and Rind, 1992] based on the cloud top height, CP [Allen and Pickering, 2002] based on the precipitation rate, CCAPE [Romps et al., 2014] based on the precipitation rate and convective available potential energy (CAPE), MFLUX [Allen and Pickering, 2002] based on the updraft mass flux and ICEFLUX [Finney et al., 2014] based on the upward cloud ice flux. We annotate in boxes the total annual Blue Jets per minute.

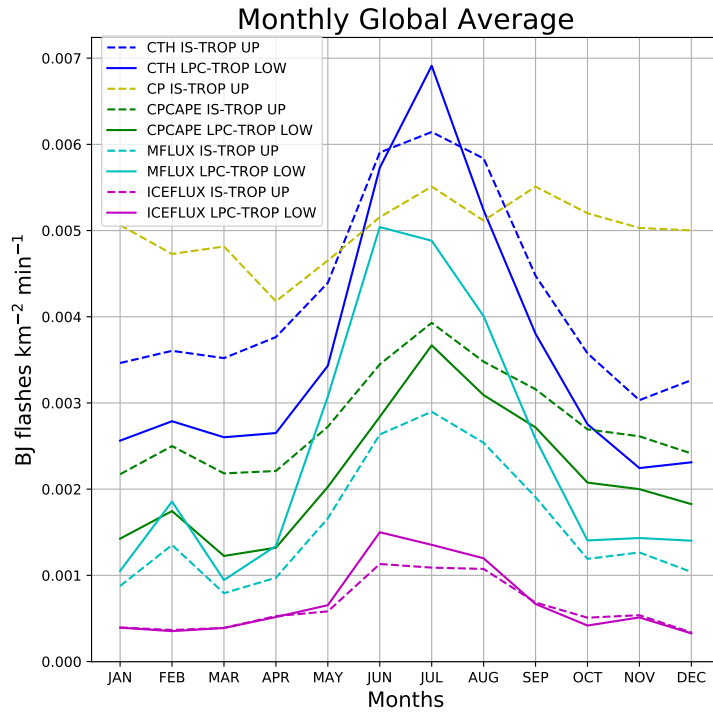


Figure 4: Monthly global average Blue Jet frequencies $\text{BJ km}^{-2}\text{day}^{-1}$ using the developed Blue Jets parameterization. The shown Blue Jet frequencies have been calculated using different lightning parameterizations denoted as CTH [Price and Rind, 1992] based on the cloud top height, CP [Allen and Pickering, 2002] based on the precipitation rate, CPCAPE [Romps et al., 2014] based on the precipitation rate and convective available potential energy (CAPE), MFLUX [Allen and Pickering, 2002] based on the updraft mass flux and ICEFLUX [Finney et al., 2014] based on the upward cloud ice flux.

3.2 Global chemical impact of Blue Jets

We follow the simulation scheme proposed in section 2.1 in order to predict the global chemical impact of Blue Jets in the atmosphere. We use the same specified dynamics to simulate the atmosphere with and without Blue Jets.

First, we run simulations of one year for all the considered Blue Jet parameterizations (subsection 3.2.1). Then, we choose some of the most representative cases and extend the simulations up to ten years in order to obtain the chemical influence of Blue Jets when equilibrium is reached (subsection 3.2.2). We do not include in this discussion the chemical impact of Blue Jets using the parameterizations that overestimate the Blue Jet frequency (“LPC-TROP UP”).

3.2.1 Transient response

In this section we present and discuss the global impact of Blue Jets over one year for all the considered cases. It is important to note that simulations of one year including the chemical perturbation of Blue Jets do not allow the atmosphere to reach the equilibrium. However, these short simulations are useful to choose the most realistic cases before extending the simulation to five and ten years. We collect in table 1 the total annual production, i.e., the total number of NO, N₂O and O molecules injected in the atmosphere by Blue Jets for all the considered cases of the CTH lightning parameterization, while we show in the supplementary material the results corresponding to other lightning parameterizations. The cases in which the production of N₂O is larger than the natural source of atmospheric N₂O (10.2 Tg N₂O-N yr⁻¹) [Davidson, 2009, Prather *et al.*, 2015] and the total occurrence rate of Blue Jets is higher than 1 BJ per minute will be considered as unrealistic scenarios. Therefore, the realistic scenarios would be most of the “IS-TROP UP”, all “IS-TROP LOW” and most of “LPC-TROP LOW”. The lower realistic scenario (3.6 × 10⁻⁶ BJ per lightning flash and R₁ = 2.5 m) produces a Blue Jet frequency of 1.4 × 10⁻³ BJ per minute and 6.6 × 10⁻⁴ Tg N₂O-N yr⁻¹ (ICEFLUX IS-TROP LOW R₁), while the higher realistic case (3.6 × 10⁻⁴ BJ per lightning flash and R₂ = 12.5 m) produces a Blue Jet frequency of 0.72 BJ per minute and 7.6 Tg N₂O-N yr⁻¹ (CTH LPC-TROP LOW R₂). The predicted production of NO in the so-called realistic cases is about two orders of magnitude lower than the production of NO by lightning. The global production of NO by Blue Jets is then negligible. The global production of O is also negligible.

Let us now estimate the transient chemical impact of Blue Jets in the atmosphere over one year. For this purpose, we calculate the global annual average vertical profile of some chemical species obtained from the simulations of Blue Jets and compare them with the profiles produced in the simulations without Blue Jets. We plot the obtained results with the cases whose predicted Blue Jet frequency is close to the maximum value estimated from the TLE occurrence reported by ISUAL (IS-TROP UP R₁ and R₂; and LPC-TROP LOW R₁ and R₂, respectively) in figure 5, together with the percentage of change at each altitude between simulations with and without Blue Jets (relative enhancement). The last three plots of figure 5 correspond to the vertical production rate of NO, N₂O and O by both Blue Jets and lightning. Figure 5 shows results for the CTH lightning parameterization, while the supplementary material shows figures collecting

results with other lightning parameterizations (CP, CPCAPE, MFLUX, ICEFLUX).

It is worth analyzing the obtained chemical impact for each considered species. The most remarkable chemical impact of Blue Jets are the enhancements in the densities of NO_x and N_2O at altitudes between 10 km and 30 km. Most of the simulations producing a realistic Blue Jet frequency and imposing $R_2 = 12.5$ m (IS-TROP UP R_2 and LPC-TROP LOW R_2) predict maximum density increases of NO_x and N_2O of 30 % and 5 %, respectively. Simulations that produce the lowest possible Blue Jet frequency (IS-TROP LOW R_1 and R_2) and simulations producing a realistic Blue Jet frequency (IS-TROP UP R_1 and LPC-TROP LOW R_1) and imposing $R_1 = 2.5$ m predict a negligible influence of Blue Jets in the global amount of NO_x and N_2O .

Vertical profiles of other species can also be influenced by the inclusion of Blue Jets in the global atmospheric chemistry. Blue Jets could produce a decrease in the upper tropospheric density of OH and HO_2 of about 5 % and 20 %, respectively. The injected NO molecules would lead to an increase in OH and a reduction of HO_2 by the process $\text{NO} + \text{HO}_2 \rightarrow \text{NO}_2 + \text{OH}$ [Murray *et al.*, 2013]. However, the conversion of NO into NO_2 can also contribute to a decrease in the concentration of OH, specially at lower altitudes by the process $\text{NO}_2 + \text{OH} + \text{M} \rightarrow \text{HNO}_3 + \text{M}$, where M represents air molecules ($\text{N}_2 + \text{O}_2$). According to our results, the concentration of HNO_3 and SO_2 could increase about 20 %. HNO_3 and SO_2 molecules can directly contribute to the production of acid rain [Seinfeld and Pandis, 2016]. The density profile of CO can exhibit both relative increases and decreases at different altitudes, as its gains and losses mechanisms depend on the concentration of OH according to the process $\text{CO} + \text{OH} \rightarrow \text{CO}_2 + \text{H}$ [Murray *et al.*, 2013]. The global density profiles of other species, such as O_3 and O, are not significantly influenced by Blue Jets.

The first panel of figure 6 shows the annual average total column density of N_2O after a WACCM4 simulation of 1 year using the lightning parameterization by Price and Rind [1992] without Blue Jets. The rest of the panels in figure 6 show the annual average total column density difference of N_2O between two simulations of 1 year with and without Blue Jets using different lightning parameterizations. The geographical distribution of N_2O changes is directly linked to the adopted lightning parameterization (see figure 3), with strong increases in N_2O in the tropics and/or mid-latitudes in relation to a local stronger Blue Jet occurrence in those regions. Interestingly, all lightning parameterizations produce an enhancement in the concentration of N_2O near the northern high latitude and polar regions. Given the limited amount of Blue Jet simulated to occur at high latitude, N_2O is likely increased by wave-driven transport and mixing from lower latitudes in the extratropical upper troposphere-lowermost stratosphere on relatively fast timescales (see e.g. Holton *et al.* [1995]). On longer timescales (see e.g. the 5 and 10 years cases presented below), increases in N_2O at high latitude can occur through poleward and downward adiabatic transport of tropical air. Similar effects are produced also in simulations of LNO_x by Grewe [2009] with high impact of LNO_x to changes at high latitude.

3.2.2 Response close to equilibrium

The analysis of the global chemical impact of Blue Jets presented in the previous section is based on a simulation of one year. As we pointed out in section 2.1, the lifetime

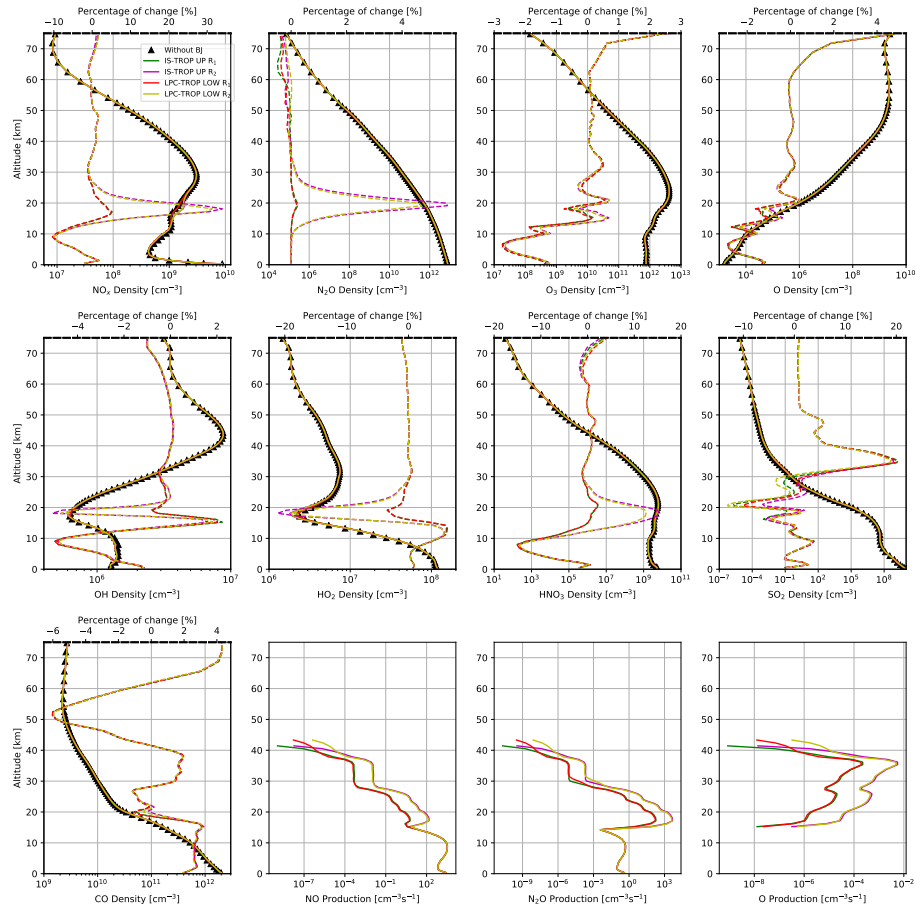


Figure 5: Solid lines correspond to annual global average density of some species after a WACCM4 simulation of 1 year including Blue Jets and using the lightning parameterization CTH [Price and Rind, 1992]. Triangles correspond to the same simulation with lightning but without Blue Jets. Dashed lines represent the percentage difference when Blue Jets are included. The last three subplots in the lower row show the total production rate of NO, N₂O and O, respectively by Blue Jets and lightning.

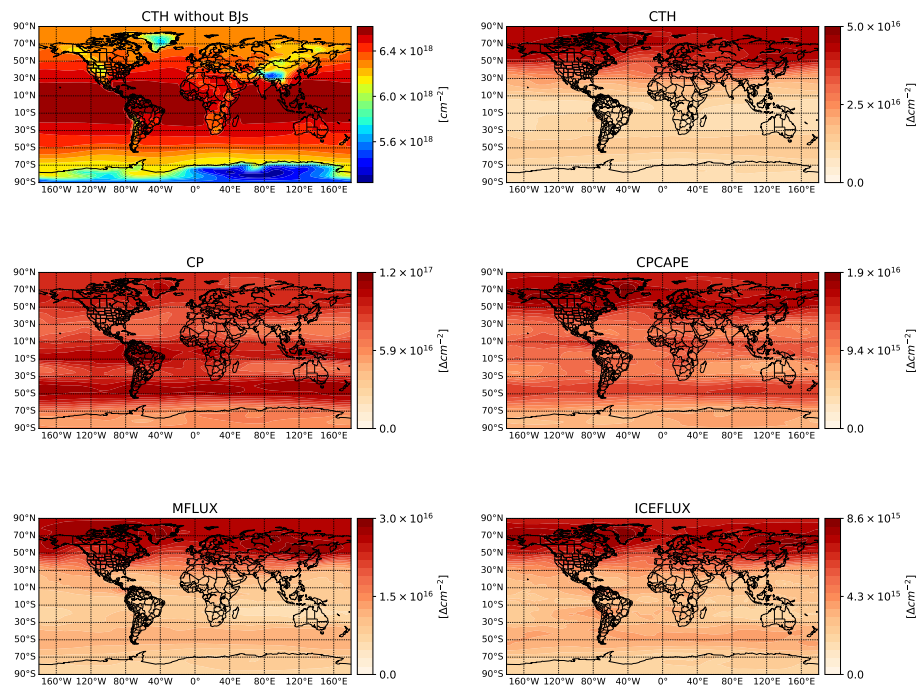


Figure 6: The top left panel shows the annual average total column density of N_2O after a WACCM4 simulation of 1 year using the lightning parameterization by *Price and Rind* [1992] without Blue Jets. The other panels correspond to the variation in the annual average total column density of N_2O between two simulations of 1 year with and without Blue Jets using different lightning parameterizations and the realistic Blue Jet parameterization “LPC-TROP UP R₁”.

of N_2O is of the order of a century, while the time scale of the overturning circulation is about 5 years.

We select two of the previously identified realistic cases in terms of Blue Jet frequency (LPC-TROP LOW R_1 and R_2) and extend the CTH-based simulations up to five and ten years. We also extend the control case without Blue Jets up to five and ten years. This approach allows us to see the global distribution after the injected species have been transported. It is important to emphasize that a simulation of more than 100 years would be necessary to reach the complete chemical equilibrium. However, such a long simulation is out of the scope of this paper. We plot on figures 7 and 8 the atmospheric chemical influence of Blue Jets annual averaged for the fifth and the tenth year of simulation, respectively. The chemical influence of Blue Jets is a factor of two larger in the five year simulation (see figure 7) than in the simulation of one year (figure 5). Hence, we cannot assume that the atmosphere has already reached an equilibrium after including Blue Jets. However, the chemical influence of Blue Jets as shown in the 10 year simulation (see figure 8) is quite similar to the one obtained in the simulation of five years, indicating that a simulation of ten years may be sufficient to estimate the global chemical impact of Blue Jets despite the 100 year lifetime of N_2O . After a simulation of ten years, the density enhancements and decreases obtained in the previous section (one year simulations) are increased by a factor of two. The increase in the tropospheric density of HNO_3 suggests that Blue Jets could also have a direct influence in the acidity of rainwater.

Let us now investigate the geographical chemical impact of Blue Jets resulting from a 10 year simulation. We plot in figure 9 the annual average total column density of some species after simulating a decade using the realistic BJ parameterization “LPC-TROP LOW R_2 ”. The differences with respect to a simulation without Blue Jets are plotted in figure 10 (total column density) and in figure 11 (longitudinally averaged vertical profile of N_2O and O_3). The maximum influence in the density of N_2O and HNO_3 is concentrated near the North Pole, as can be seen in figures 10 and 11. Figure 11 also shows a slight depletion of about 5 % in the column density of O_3 above 30 km near the Equator. Although the total column density of O_3 in polar regions is not significantly affected by Blue Jets (see figure 10). Figure 11 shows that there is an increase of O_3 below 18 km of altitude at all latitudes and a decrease above 20 km of altitude of about 5 %. Some other species show differences that are distributed through mid latitudes, especially around points of maximum Blue Jet occurrence rates.

The obtained density profile of N_2O can be compared with measurements of Aura-MLS and MIPAS [Plieninger *et al.*, 2016] to determine how well the obtained response corresponds with observations. As detailed by Plieninger *et al.* [2016], Aura-MLS and MIPAS (operating at reduced resolution mode) measured the density profile of N_2O over a wide range of latitudes. Figure 7 of Plieninger *et al.* [2016] shows the reported N_2O and their error bars from Aura-MLS and MIPAS. We plot these profiles together with the N_2O obtained after 10 year simulations with Blue Jets in figure 12. It can be seen that the equilibrium response of WACCM4 including Blue Jets produce a global average N_2O profile that falls in the range reported by Aura-MLS and MIPAS for the cases “LPC-TROP LOW R_1 and R_2 ”.

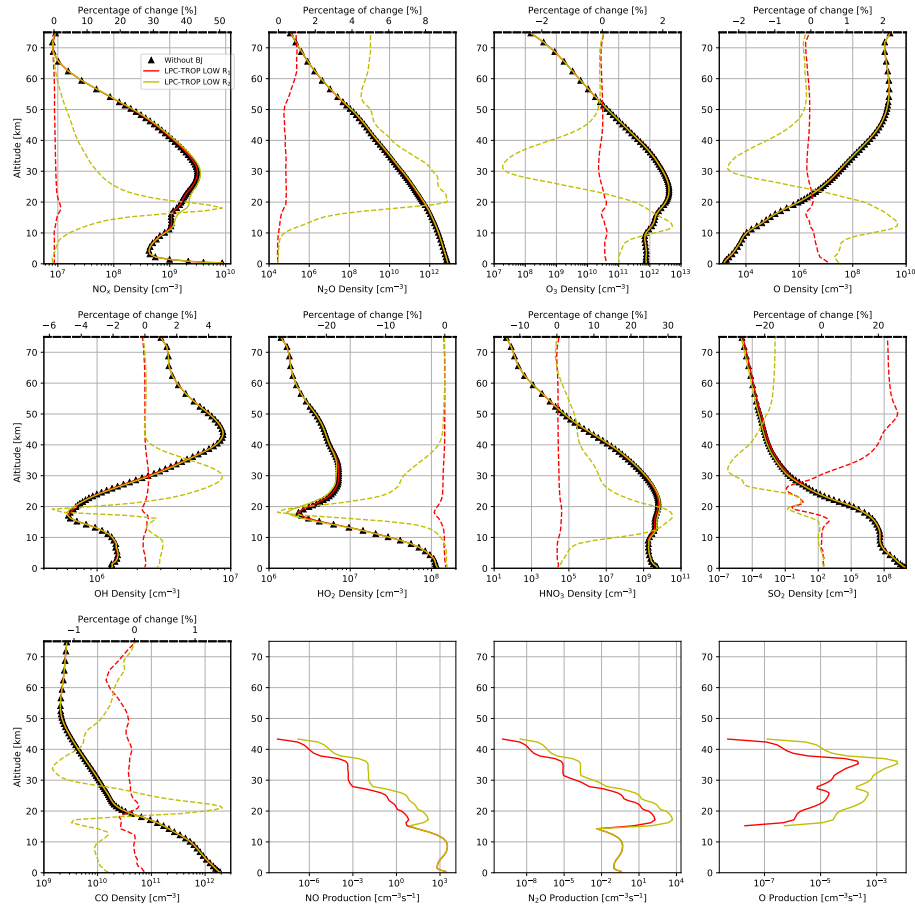


Figure 7: Solid lines correspond to annual global average density of some species after a WACCM4 simulation of 5 years including Blue Jets and using the lightning parameterization CTH [Price and Rind, 1992]. Triangles correspond to the same simulation with lightning but without Blue Jets. Dashed lines represent the percentage variation with respect to a similar simulation without Blue Jets. The last three subplots in the lower row show the total production rate of NO, N₂O and O by lightning and Blue Jets. Note that the horizontal upper scale of figures 5, 7 and 8 are different.

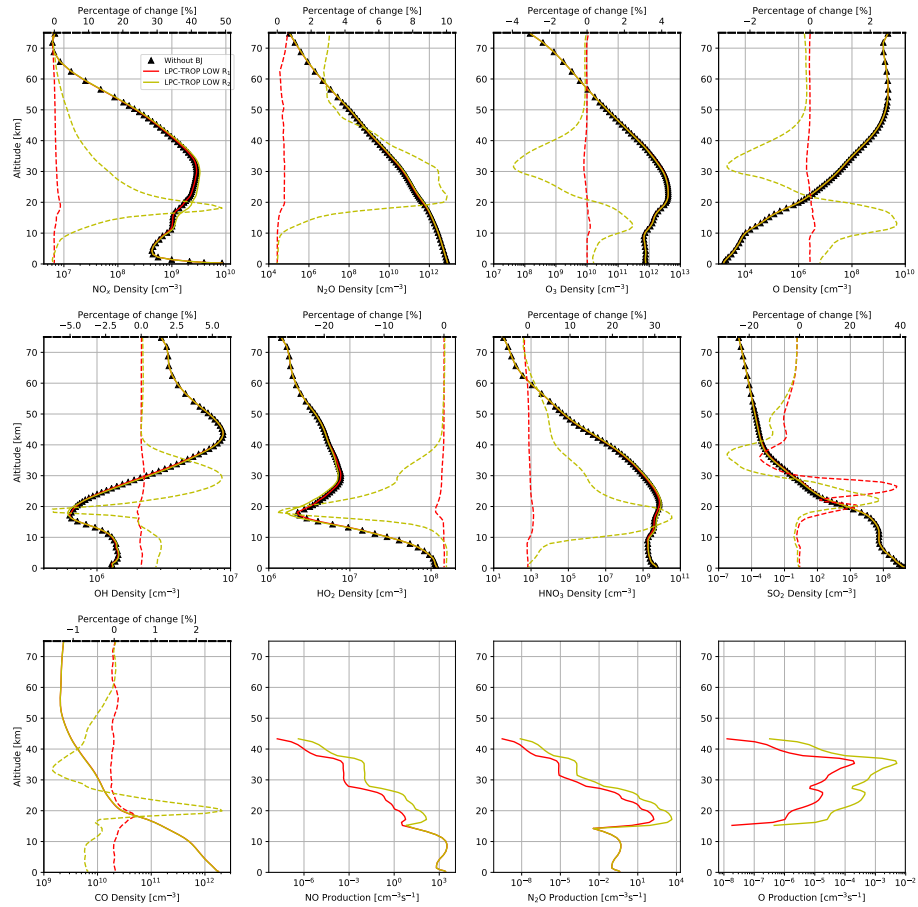


Figure 8: Solid lines correspond to annual global average density of some species after a WACCM4 simulation of 10 year including Blue Jets and using the lightning parameterization CTH [Price and Rind, 1992]. Triangles correspond to the same simulation with lightning but without Blue Jets. Dashed lines represent the percentage variation with respect to a similar simulation without Blue Jets. The last three subplots in the lower row show the total production rate of NO, N₂O and O by lightning and Blue Jets. Note that the horizontal upper scale of figures 5, 7 and 8 are different.

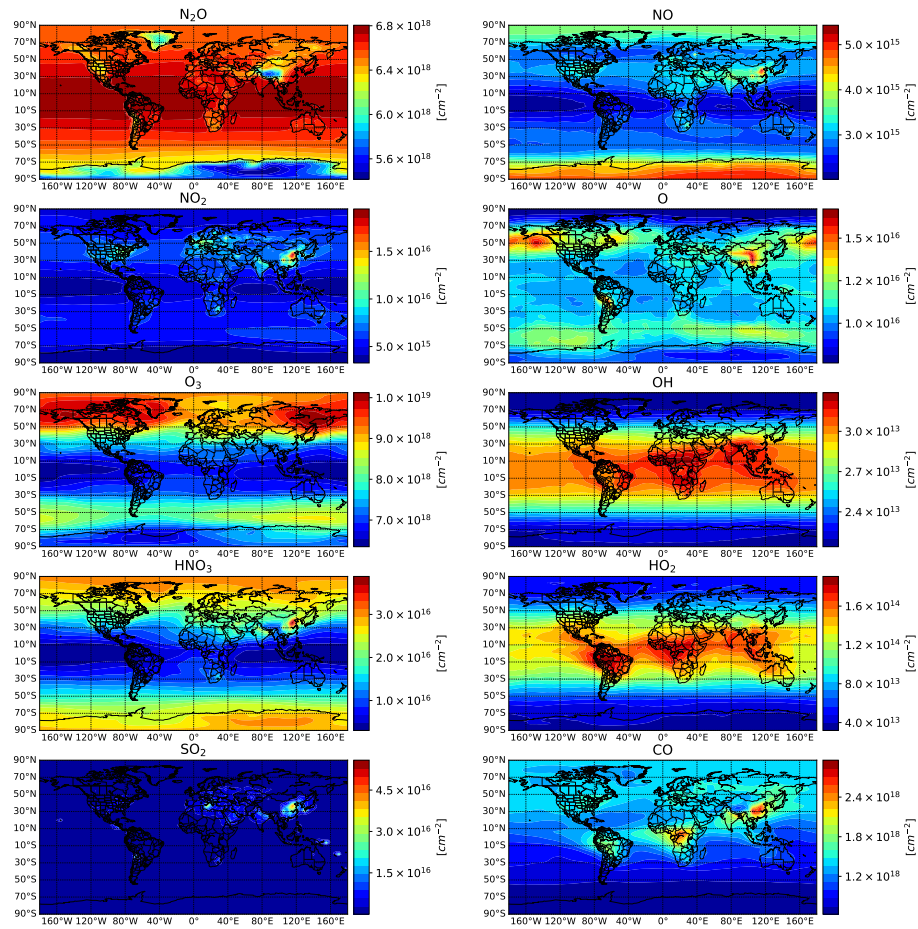


Figure 9: Annual average total column density of some chemical species after a WACCM4 simulation of 10 years including Blue Jets. These subplots have been calculated using the lightning parameterization based on the cloud-top height CTH [Price and Rind, 1992] and the Blue Jets parameterization denoted as “LPC-TROP LOW R₂”.

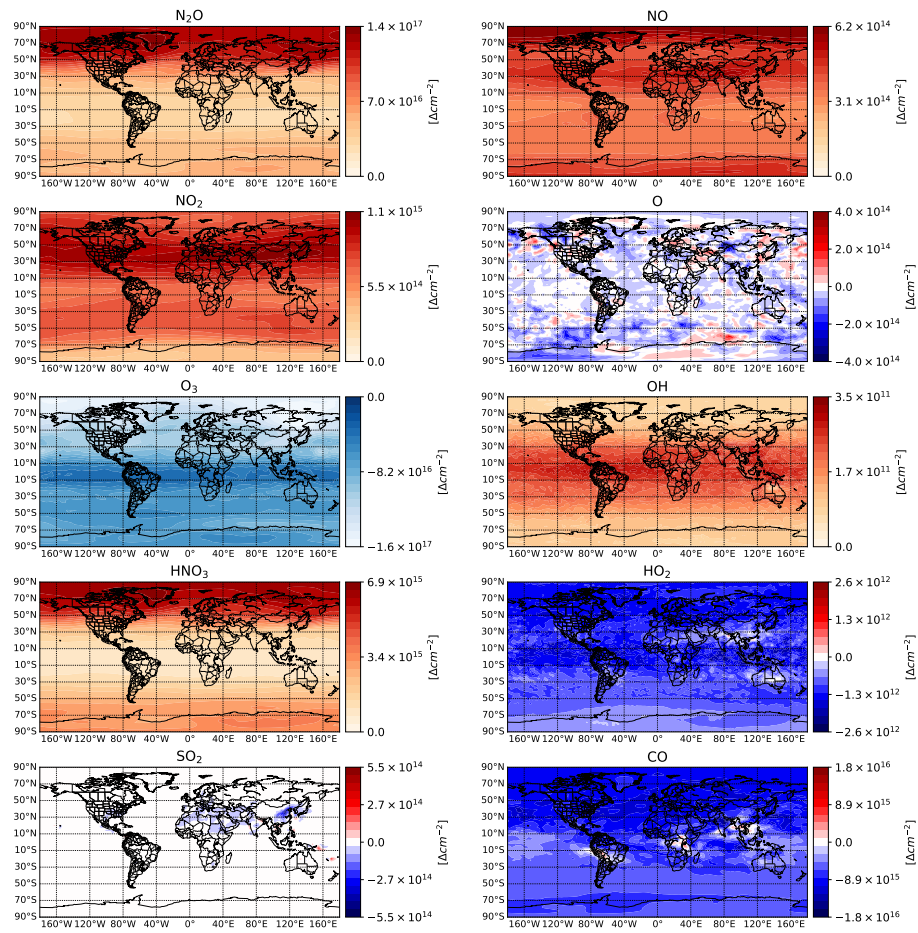


Figure 10: Differences in the annual average total column density of some chemical species between two simulations of 10 years with (as in figure 9) and without Blue Jets. Positive values correspond to enhancement in densities due to Blue Jets, while negative variations represent density decrease produced by Blue Jets.

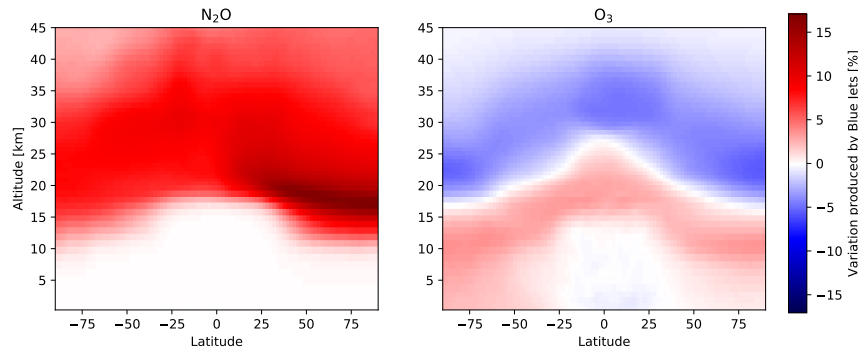


Figure 11: Latitude-altitude distribution of the differences in the annual average density profile of N_2O and O_3 between two simulations of 10 years with (as in figure 9) and without Blue Jets. These variations are longitudinally average.

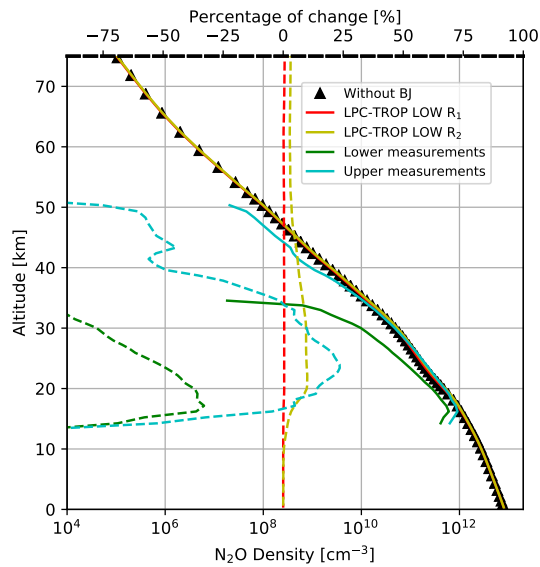


Figure 12: Comparison between the global average profile of N_2O obtained by Aura-MLS and MIPAS [Plieninger et al., 2016] and the global average profile of N_2O after 10 year simulation with and without Blue Jets in WACCM4 (second subplot of the first row of figure 8) shown here with red, purple and superimposed green solid lines (absolute values) and dashed lines (percentage of change). Blue and yellow solid lines correspond to the lower and upper total N_2O density reported by Aura-MLS and MIPAS according to the error bars shown in figure 7 of Plieninger et al. [2016]. Blue and yellow dashed lines are the percentage of difference between the lower and upper total N_2O density reported by Aura-MLS and MIPAS and the N_2O profile of a WACCM4 simulation without Blue Jets.

4 Summary and conclusions

We have introduced for the first time Blue Jets in an atmospheric global circulation model. The Blue Jet parameterization presented here is a step further in the coupling between local and global models of atmospheric electricity phenomena. Previous local models of Blue Jets predicted an important local enhancement of N_2O and NO_x molecules between 18 km and 38 km of altitude, as well as a depletion of O_3 [Winkler and Notholt, 2015]. The significant local chemical influence of Blue Jets suggests that their global chemical influence could be non-negligible.

In this work, we have developed two different global parameterizations of Blue Jets. The first parameterizations (IS-TROP LOW and UP) is based on the ratio between lightning and TLE occurrence rate as reported by ISUAL and introduces a physical-based geographical dependence for the occurrence of Blue Jets. These parameterizations link the occurrence of TLEs with the altitude of the cloud top. It imposes the condition that the top of the thunderclouds must be near the tropopause in order to favor the inception of Blue Jets. Finally, the second Blue Jet parameterizations (LCP-TROP LOW and UP) are based on the observational evidences pointing to a close relationship between strong lightning discharges and Blue Jets in thunderstorms. We have obtained a good agreement between the TLE occurrence rate reported by ISUAL and the predicted ones by the Blue Jet parameterizations introduced in WACCM4 except with LCP-TROP UP.

The implementation of these Blue Jet parameterizations in WACCM4 has allowed us to estimate their global chemical influence in the atmosphere. We have made several assumptions about the geometry of single Blue Jets in order to couple the local chemical model of Winkler and Notholt [2015] with the global chemistry implemented in WACCM4. Depending on the differences between the obtained N_2O profile and the profiles reported by Aura-MLS and MIPAS [Plieninger *et al.*, 2016], we have distinguished between realistic and extreme cases. According to the most realistic cases, Blue Jets would inject between 6.6×10^{-4} Tg $\text{N}_2\text{O-N yr}^{-1}$ and 7.6 Tg $\text{N}_2\text{O-N yr}^{-1}$ near 20 km of altitude. The average value 3.8 Tg $\text{N}_2\text{O-N yr}^{-1}$ corresponds about 38 % of natural N_2O sources. In addition, we have obtained that the global production of NO_x by Blue Jets is between 10^{-5} Tg NO-N yr^{-1} and 0.14 Tg NO-N yr^{-1} . The average value 0.07 Tg NO-N yr^{-1} is about 1 % of natural NO sources, two orders of magnitude below the production of NO_x by lightning on the troposphere. WACCM4 has allowed us to estimate the influence of Blue Jets in other chemical species apart from NO_x and N_2O . In particular, we have found that the stratospheric (between 20 km and 40 km) concentration of some species such as OH, HO_2 , SO_2 and HNO_3 could also be influenced by Blue Jets. Finally, we have also found that the inclusion of Blue Jets in WACCM4 can account for a maximum decrease of O_3 by about 5 % between 20 km and 40 km of altitude.

There are several reasons behind the high uncertainty in what we call realistic results. First, there is not a clear convenient global parameterization of lightning to be combined with the proposed Blue Jet parameterizations [Tost *et al.*, 2007]. Second, the detailed mechanisms behind the production of Blue Jets are still poorly described, which makes it difficult to build global parameterization for Blue Jets. Finally, the complex chemistry taking place in the high temperature leader-phase of Blue Jets together

with the uncertainties in their electrodynamical radius imply an important uncertainty of the local chemical influence of Blue Jets. All in all, we consider this work as a first approximation to the understanding of the influence of Blue Jets on the global atmospheric chemistry.

L-BJ parameterizations	BJ frequency [min^{-1}]	Tg NO-N yr^{-1}	Tg N ₂ O-N yr^{-1}	Tg O yr^{-1}
CTH IS-TROP UP R ₁	0.9	6×10^{-3}	0.42	3×10^{-7}
CTH IS-TROP UP R ₂	0.9	0.16	10.4	7×10^{-6}
CTH IS-TROP LOW R ₁	9×10^{-3}	6×10^{-6}	4×10^{-3}	3×10^{-9}
CTH IS-TROP LOW R ₂	9×10^{-3}	1.5×10^{-3}	0.1	8×10^{-8}
CTH LPC-TROP UP R ₁	7.2	5×10^{-2}	3.0	3×10^{-6}
CTH LPC-TROP UP R ₂	7.2	1.36	76.0	7×10^{-5}
CTH LPC-TROP LOW R ₁	0.72	5×10^{-3}	0.3	3×10^{-7}
CTH LPC-TROP LOW R ₂	0.72	0.14	7.6	7×10^{-6}

Table 1: BJ frequency and production of NO, N₂O and O obtained for different one year simulations using BJ parameterizations and the CTH lightning parameterization.

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